

Projection of Global Future Lightning Occurrence using only Large-Scale Environmental Variables in CAM5

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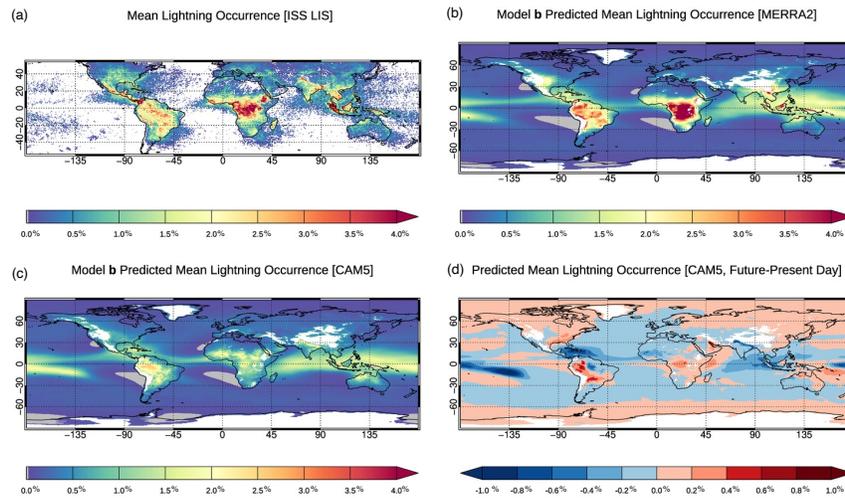
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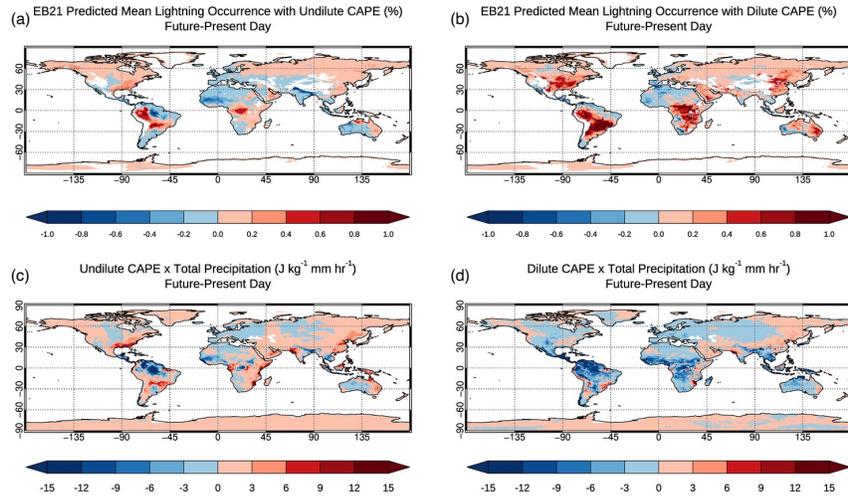
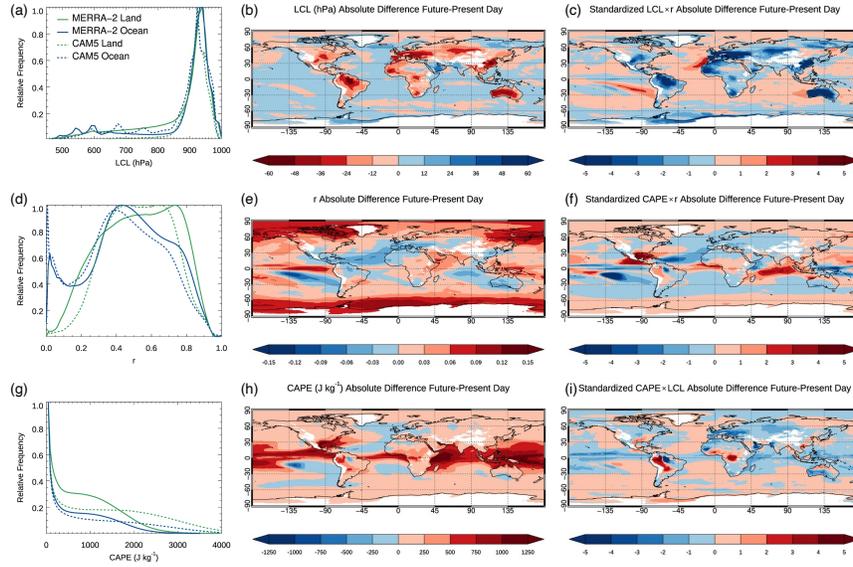
³TAMU

March 15, 2024

Abstract

This study evaluates a lightning parameterization that utilizes only large-scale environmental variables (i.e., convective available potential energy (CAPE), column moisture, and lifting condensation level (LCL)) for present-day (2017-19) and end-of-century (2098-2100) RCP8.5 climate scenarios in the Community Atmosphere Model version 5 (CAM5). Using a single equation, the present-day prediction can produce a reasonable land/ocean ratio in lightning occurrence. The end-of-century prediction shows relative increases of about 50% over higher-latitude land, but much more variable increases and decreases across mid-latitude ocean and the tropics such that the overall global lightning occurrence is expected to slightly decrease. Lightning occurrence over land predicted from present-day CAM5 is less than that using MERRA-2 reanalysis because of differences in the basic-state variables used as predictors. In addition, the choice of dilute or undilute CAPE will impact future lightning predictions over land, but the environment-only parameterization results are more consistent than a CAPE x precipitation parameterization.





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2 **only Large-Scale Environmental Variables in CAM5**

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6 **Key Points:**

- 7 • A single-equation based only on environmental variables provides reasonable land
8 and ocean lightning occurrence predictions in CAM5.
9 • Lightning occurrence is projected to increase at higher latitudes by the end of cen-
10 tury, but the long-term trend varies across the tropics.
11 • Basic state biases and the type of CAPE used can impact current and future pre-
12 dictions of lightning patterns and magnitudes.

Abstract

This study evaluates a lightning parameterization that utilizes only large-scale environmental variables (i.e., convective available potential energy (CAPE), column moisture, and lifting condensation level (LCL)) for present-day (2017-19) and end-of-century (2098-2100) RCP8.5 climate scenarios in the Community Atmosphere Model version 5 (CAM5). Using a single equation, the present-day prediction can produce a reasonable land/ocean ratio in lightning occurrence. The end-of-century prediction shows relative increases of about 50% over higher-latitude land, but much more variable increases and decreases across mid-latitude ocean and the tropics such that the overall global lightning occurrence is expected to slightly decrease. Lightning occurrence over land predicted from present-day CAM5 is less than that using MERRA-2 reanalysis because of differences in the basic-state variables used as predictors. In addition, the choice of dilute or undilute CAPE will impact future lightning predictions over land, but the environment-only parameterization results are more consistent than a CAPE \times precipitation parameterization.

Plain Language Summary

Lightning parameterizations currently being used in climate model studies use output from other physical parameterizations (i.e., cloud ice, precipitation, etc.). These variables have large uncertainties that propagate into the lightning prediction and can vary strongly amongst models, thus requiring scaling factors to produce realistic and consistent lightning predictions. In addition, almost all existing parameterizations require separate land and ocean equations to produce reasonable global lightning patterns, and many still produce unrealistic ratios with too much oceanic lightning. We show here that we can produce a reasonable global lightning occurrence distribution in a climate model using only three large-scale environmental variables derived from temperature and humidity profiles and a single equation applicable to both land and ocean components. While these variables can still have uncertainties and biases amongst models, they are less than the cloud and precipitation outputs, thus providing a more stable framework for assessing lightning changes. Our end-of-century projection under a high-emissions scenario shows relatively large increases in lightning occurrence over land at mid- and high-latitudes in the Northern Hemisphere, but a varying pattern of increases and decreases across the tropics such that the global mean lightning occurrence is expected to slightly decrease by the end of the century.

1 Introduction

Understanding lightning and its relationship with the large-scale environment is important in simulating lightning in global climate models (GCMs) in order to predict how lightning will vary with climate change, and how upper-tropospheric chemistry and wildfires associated with lightning will be impacted (e.g., Krause et al., 2014; Whaley et al., 2024). The large-scale environment plays a key role in storm dynamics, and therefore lightning development. Most previous studies have investigated lightning's relationship with cloud features and precipitation, but few have isolated the role of the large-scale environment for the prediction of lightning.

One of the earliest parameterizations predicted lightning flash rates using only convective cloud-top height (Price & Rind, 1992, hereafter PR92) and has been used as the basis of many other parameterizations (Boccippio, 2002; Luhar et al., 2021; Michalon et al., 1999; Zhang et al., 2021). In a warming climate, parameterizations that use PR92 typically predict a global increase in lightning (Clark et al., 2017; Finney et al., 2018; Krause et al., 2014; Price & Rind, 1994). However, PR92 uses separate land and ocean equations to predict lightning and requires a scaling of these equations to the observed

62 global mean lightning. In addition, convective cloud-top height, especially when output
63 as a grid-scale value from a coarse-resolution GCM, is a highly-derived variable and re-
64 mains an indirect measure of convective intensity and lightning.

65 Different cloud and rain variables have since been utilized in lightning parameter-
66 izations from GCM output such as convective mass flux (Allen & Pickering, 2002; Grewe
67 et al., 2001; Magi, 2015), upward cloud ice flux (Deierling et al., 2008; Finney et al., 2014;
68 Romps, 2019), convective precipitation (Allen & Pickering, 2002; Magi, 2015; Meijer et
69 al., 2001), cloud droplet concentration (Michalon et al., 1999), graupel mixing ratio and
70 updraft velocities (McCaul et al., 2009; Williams, 2005; Zipser & Lutz, 1994), cold cloud
71 depth (Yoshida et al., 2009), and cloud base height (Lopez, 2016). Most of these param-
72 eterizations produce general increases in lightning flash rates for warming climates (Clark
73 et al., 2017; Finney et al., 2016, 2020), except when using ice-based parameterizations
74 (Finney et al., 2018; Romps, 2019). However, these frameworks still require separate land
75 and ocean equations and often need to be scaled to the current global mean lightning
76 to provide a realistic prediction. In addition, Charn and Parishani (2021) found that the
77 ice-based lightning parameterizations may be sensitive to the microphysics scheme used,
78 not necessarily to the variables used to predict lightning, which adds motivation to avoid
79 highly uncertain storm-scale variables as inputs for lightning parameterizations in GCMs.

80 The inclusion of large-scale environmental variables in predicting lightning in GCMs
81 has become more prevalent in recent years (Romps et al., 2014; Stolz et al., 2015, 2017;
82 Wang et al., 2018; Etten-Bohm et al., 2021) and could help reduce the large uncertainty
83 that is carried when using cloud and convection variables as predictors. Utilizing large-
84 scale variables like convective available potential energy (CAPE) can be beneficial be-
85 cause of how closely it relates to a storm’s thermodynamics. Romps et al. (2014) (here-
86 after R14) used CAPE and precipitation (CAPE \times P) over the continental United States
87 (CONUS) to predict lightning flash rate. Evaluating the parameterization in multiple
88 GCMs, R14 found that CAPE increased over CONUS between the current climate and
89 late 21st century in all the models, therefore also increasing the lightning flash rate. It
90 is worthwhile noting that future projection of precipitation sometimes increased and some-
91 times decreased depending on the GCM and did not constrain the lightning prediction
92 nearly as much as CAPE.

93 Although the R14 parameterization performed well over CONUS, it did not trans-
94 late well on a global scale because it could not distinguish between land and ocean (Romps
95 et al., 2018). Cheng et al. (2021) had better success using a different equation over ocean,
96 but a similar issue as discussed previously occurs with an arbitrary separation of land
97 and ocean equations to predict lightning. Stolz et al. (2015, 2017) were better able to
98 differentiate land and ocean lightning environments by using a combination of cloud and
99 environmental parameters in a multiple linear regression model, but still did not com-
100 pletely capture the spatial pattern of global lightning, overpredicting over the ocean and
101 underpredicting over land (Stolz et al., 2021).

102 Etten-Bohm et al. (2021) (hereafter EB21) presented a lightning parameterization
103 based solely on large-scale environmental variables, with the goal of limiting the issues
104 and uncertainty in other parameterizations mentioned previously. EB21 evaluated a num-
105 ber of covariate sets from reanalysis output and each prediction represented the spatial
106 pattern of lightning occurrence well, including a distinction between land and ocean us-
107 ing just one equation. They found that the use of three environmental variables (CAPE,
108 lifting condensation level [LCL], and column saturation fraction [r]) and their interac-
109 tions provided the best basis for a GCM parameterization in terms of performance and
110 simplicity.

111 The main goals of this study are to implement and evaluate this EB21 environment-
112 only lightning parameterization in the high-resolution (25 km) Community Atmosphere
113 Model version 5 (CAM5), project end-of-century global lightning occurrence changes,

114 and determine the environmental factors most important to the changes. Additionally,
 115 we will assess how the EB21 parameterization performs compared to the CAPE \times P pa-
 116 rameterization over land, including sensitivity tests using different CAPE calculations
 117 (i.e., dilute and undilute) since there aren't standard definitions of CAPE in GCMs.

118 2 Data and Methods

119 EB21 utilized a logistic regression trained on Tropical Rainfall Measuring Mission
 120 (TRMM) Lightning Imaging Sensor (LIS) observations (Kummerow et al., 1998) and Modern-
 121 Era Retrospective Analysis for Research and Application Version 2 (MERRA-2) reanal-
 122 ysis data (Gelaro et al., 2017) to predict lightning occurrence based on 3-hourly, 0.5° in-
 123 put. EB21 tested three predictor sets increasing in complexity from model **a** to **c**. Only
 124 model **b** (with predictors CAPE, LCL, r , and their interactions) will be evaluated in this
 125 study since it provided the best balance between simplicity and performance amongst
 126 the three models. The parameterization outputs the predicted probability of lightning
 127 occurrence at each grid point from zero (0% chance) to one (100% chance).

128 The GCM environmental predictors for this study were obtained from a 0.25° res-
 129 olution, free-running version of CAM5 (Meehl et al., 2019; Neale et al., 2012). Three-
 130 hourly temperature, LCL, and specific humidity fields were interpolated to a 0.5° grid
 131 to match the LIS and MERRA-2 datasets. CAPE and r were then computed from the
 132 temperature and specific humidity profiles, and all variables were standardized to have
 133 a mean of zero and a standard deviation of 1. The CAM5 predictors for present day (2017-
 134 19) were input into the EB21 parameterization, which was further applied to the end-
 135 of-century (2098-2100) simulation under the Representative Concentration Pathway (RCP)
 136 8.5 scenario to assess the relative impact of a warming climate on lightning production.
 137 Note that since CAM5 is free running, the years chosen may not specifically correspond
 138 to those years, but using three years should still provide a reasonable mean representa-
 139 tion of the present-day and future climates.

140 CAPE can be obtained directly from CAM5 output, but CAM5 uses a dilute-plume
 141 model where entrainment of environmental air is incorporated (Neale et al., 2008). Us-
 142 ing the CAPE \times P parameterization, Charn and Parishani (2021) found that lightning
 143 predictions varied depending on how CAPE was calculated, with undilute CAPE pro-
 144 jecting a $\sim 7\%/K$ increase in lightning and dilute CAPE only projecting a $\sim 1\%/K$ in-
 145 crease. The authors noted that neither case is completely correct, and flash rates pre-
 146 dicted using CAPE \times P are likely somewhere between the two cases. Only undilute CAPE
 147 will be used in Sections 3.1 and 3.2, with a caveat that greater decreases could be pro-
 148 jected as a result. Sensitivity tests using dilute CAPE will be presented in Section 3.3.

149 3 Results

150 3.1 CAM5 Lightning Projection

151 CAM5 fields were input into the logistic regression from EB21 to compute a pre-
 152 dicted mean lightning occurrence and compared to the International Space Station (ISS)
 153 LIS (Blakeslee et al., 2020, Figure 1a) for present day (2017-19). Although the lightning
 154 parameterization was trained with TRMM LIS observations, the ISS expands on the lat-
 155 itudinal extent of TRMM (from 35° to 54°) and allows for greater comparison with CAM5's
 156 global output. Following EB21, elevation over 1500 meters is removed because of the in-
 157 accurate predictions from the logistic regression, likely due to the LCL term. Whaley
 158 et al. (2024) found improvements by disregarding the LCL term over high elevation in
 159 version 5.1 of the Canadian Earth System Model (CanESM). Figure 1b is similar to Fig-
 160 ure 9c in EB21 except for using years 2017-19 and all latitudes. The overall magnitudes
 161 increase in Figure 1b as a result of the standardization of the predictors to have a mean

162 of zero and a standard deviation of 1. The fields change when extending to higher latitudes,
 163 resulting in different standardized variables, and therefore predictions.

164 The MERRA-2 lightning predictions in Figure 1b match the LIS observations well
 165 (as expected since the parameterization was trained using MERRA-2 data), albeit with
 166 some overprediction over the ocean. For example, the land/ocean lightning occurrence
 167 ratio observed by ISS LIS is 5.1, while the MERRA-2 ratio for the same latitude range
 168 is 2.2. However, these ratios are much closer to one another than the land/ocean flash
 169 rate ratios reported by Charn and Parishani (2021) between observations and five other
 170 lightning parameterizations, some of which had land/ocean lightning ratios less than 1.

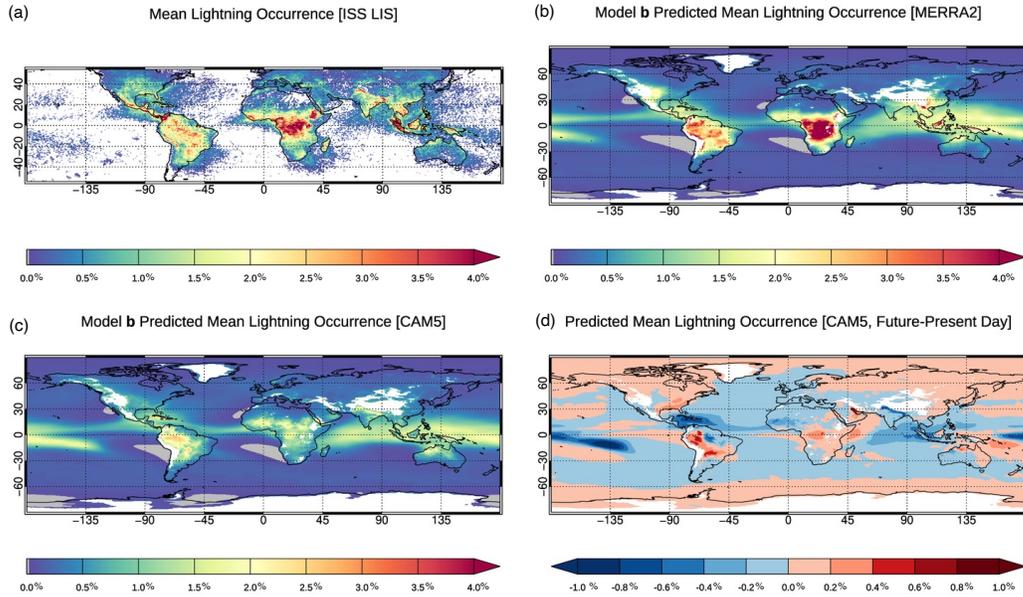


Figure 1. Present-day (2017-2019) lightning occurrence (in %) from (a) ISS LIS observations and (b) MERRA-2 and (c) CAM5 predictions using the EB21 parameterization. (d) CAM5 lightning occurrence difference between end-of-century (2098-2100) and present-day.

171 When applied to CAM5 environmental variables, the EB21 lightning parameterization
 172 produces a large underprediction over land (Figure 1c). However, expected regional
 173 variations still exist, including more lightning over the Amazon and central Africa
 174 compared to other land regions and greater overall lightning occurrence over land
 175 compared to ocean with a land/ocean ratio of 1.5. This result is promising considering that
 176 the parameterization does not have separate equations for land and ocean and does not
 177 scale the prediction to match the global mean lightning observations, which most previous
 178 lightning parameterizations have done (e.g., Clark et al., 2017). An environment-only
 179 lightning parameterization would also be expected to be more consistent between
 180 different GCMs, since cloud and precipitation variables, highly parameterized in GCM
 181 themselves, can vary much more widely compared to environmental variables (e.g., Charn
 182 & Parishani, 2021; Romps et al., 2014). However, discrepancies between the basic-state
 183 input parameters must exist between MERRA-2 and CAM5 to account for the difference
 184 in the lightning predictions in Figures 1b and c, which will be addressed in Section
 185 3.2.

186 The EB21 parameterization was further applied to output from a CAM5 end-of-
 187 century high-emissions climate run. Figure 1d indicates varied future changes in light-

188 ning occurrence over both land and ocean with increases (decreases) shown in red (blue).
 189 While many land regions indicate increasing lightning occurrence, including most higher
 190 latitude land in the Northern Hemisphere, the southeastern US, western Amazon, cen-
 191 tral Africa, and eastern Australia, other land regions show decreases, such as the cen-
 192 tral US, northeastern Amazon, Sahel, Indian subcontinent, and western Australia. The
 193 ocean shows large absolute decreases over regions that tend to have more lightning in
 194 present-day CAM5, like the South Pacific convergence zone, Caribbean Sea, Atlantic ITCZ,
 195 and Indian Ocean. Lightning is projected to increase over the ocean near the edges of
 196 these higher lightning occurrence regions. Despite many regions of increases, including
 197 higher-latitude land regions that show a relative increase of $\sim 50\%$, the global mean light-
 198 ning occurrence is predicted to decrease by about 5%. These results are generally con-
 199 sistent with end-of-century predictions using the EB21 parameterization on output from
 200 CanESM5.1 (Whaley et al., 2024).

201 Figure 1d contrasts with many previous studies that have shown more widespread
 202 increases (Clark et al., 2017; Finney et al., 2016; Romps et al., 2014; Schumann & Huntrieser,
 203 2007; Williams, 2005) or decreases (Jacobson & Streets, 2009) in global tropical light-
 204 ning flash rates in a warming climate. However, lightning parameterizations are not only
 205 sensitive to the parameters used (Finney et al., 2018; Romps, 2019), but also the method-
 206 ologies used to train the parameterization and the models in which they are implemented.
 207 For example, Finney et al. (2020) used a high-resolution, convection-permitting model
 208 and the McCaul et al. (2009) ice-based lightning parameterization to investigate light-
 209 ning day changes (similar to lightning occurrence) regionally and found a similar, albeit
 210 opposite, varied pattern to the one presented in Figure 1d over Africa.

211 3.2 Basic-state Variable Analysis

212 To evaluate environmental factors driving differences between the MERRA-2 and
 213 CAM5 present-day lightning predictions and changes in the projected mean lightning
 214 occurrence between present-day and end-of-century climate scenarios in CAM5, the three
 215 predictors (LCL, r , and undilute CAPE) are investigated separately. Figure 2 (left col-
 216 umn) shows histograms of each variable over land (green) and ocean (blue) from MERRA-
 217 2 (solid) and CAM5 (dashed). While the MERRA-2 and CAM5 environmental variable
 218 distributions show general similarities, there are some notable differences that help ex-
 219 plain the discrepancies between their lightning predictions in Figure 1. For example, while
 220 LCLs maximize around 900 hPa in both datasets (Figure 2a), offsets occur as LCLs get
 221 higher. For MERRA-2, land has relatively more LCLs between 850 and 650 hPa com-
 222 pared to ocean, while the opposite is true for CAM5 where the ocean has higher LCLs
 223 than land. Higher LCLs (more convective environment) would increase lightning occur-
 224 rence (as shown in EB21), providing one reason why lightning occurrence is underpre-
 225 dicted over land and overpredicted over ocean in CAM5.

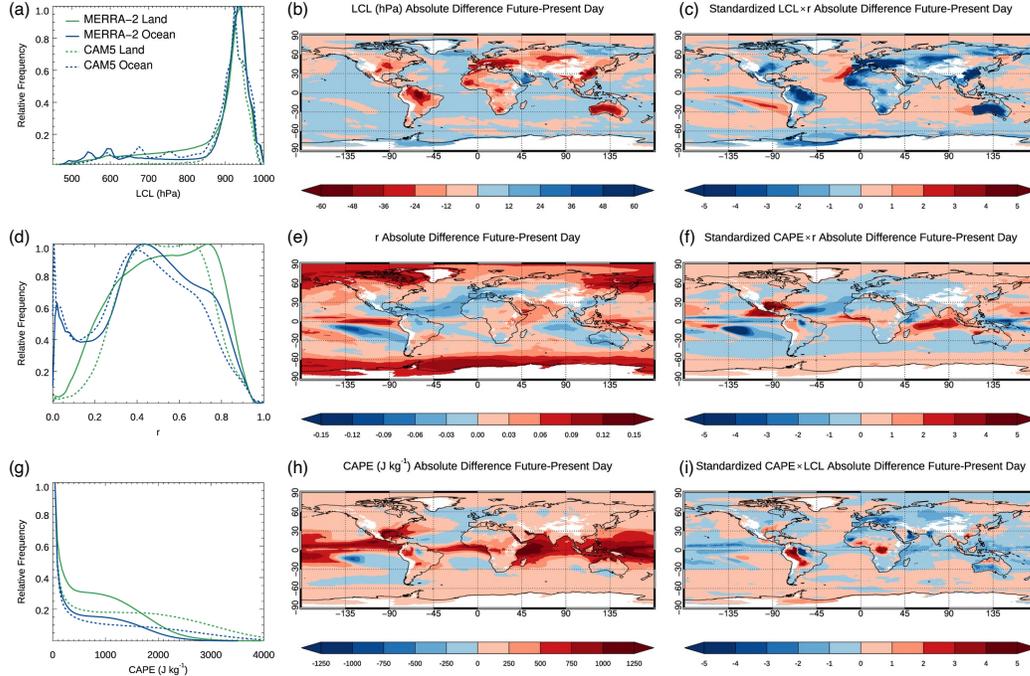


Figure 2. Histograms of land and ocean environmental variables for MERRA-2 and CAM5 for (a) LCL, (d) r , and (g) undilute CAPE for present day. Absolute differences between CAM5 end-of-century and present-day climates for (b) LCL, (e) r , and (h) CAPE and standardized interactions (c) LCL and r , (f) CAPE and r , and (i) CAPE and LCL.

226 In addition, Figure 2d shows that CAM5 r is shifted left (indicating a drier envi-
 227 ronment) compared to MERRA-2 over both land and ocean at r values where lightning
 228 is most likely to occur (i.e., $r > 0.7$, EB21). This shift also helps explain why large light-
 229 ning underpredictions happen over land in CAM5, while the drier ocean environments
 230 likely offset the higher LCLs making the CAM5 ocean lightning prediction more simi-
 231 lar to MERRA-2.

232 The CAPE distribution comparisons are more nuanced. Figure 2g indicates that
 233 MERRA-2 has a higher occurrence of moderate CAPE (up to 1800 J kg^{-1}) compared
 234 to CAM5, but that CAM5 produces more CAPE values $> 1800 \text{ J kg}^{-1}$. EB21 showed
 235 that essentially any positive CAPE would enhance lightning occurrence so it is unclear
 236 how these distribution differences would contribute to MERRA-2 and CAM5 lightning
 237 prediction differences.

238 To evaluate the spatial variability of the environmental variables and their poten-
 239 tial contribution to end-of-century lightning changes, the middle column of Figure 2 shows
 240 the absolute change between the future and present-day for each of the individual predic-
 241 tors from CAM5. Red represents changes that would be expected to enhance light-
 242 ning occurrence, and blue is the opposite. Note that we standardize individual predic-
 243 tors around their mean values before they are input into the logistic regression such that
 244 the standardized inputs (not shown) will shift to be more negative (blue) for r and CAPE
 245 because their mean individual change at the end of the century is greater than zero, while
 246 the mean LCL change is around zero.

247 Figure 2b shows that LCL decreases up to 60 hPa almost everywhere over land (i.e.,
 248 attains higher heights) by 2100, except for a handful of regions like Saudi Arabia and

249 the Indian subcontinent where LCLs increase by 15-30 hPa (i.e., become lower in height).
 250 The opposite is true almost everywhere over the ocean, where LCL values are projected
 251 to increase and thus lower in height by the end of the century, although the magnitude
 252 of change is much smaller than over land. The LCL changes in Figure 2b only partially
 253 align with the lightning changes in Figure 1d (i.e., the LCL pattern suggests large light-
 254 ning increases over land and smaller decreases over ocean globally) so other variables and
 255 their interactions remain at play.

256 Future r shows large increases in CAM5 pole-ward of 45°N and 45°S and more varied
 257 changes over land and ocean in the tropics and subtropics (Figure 2e). Changes in
 258 r often offset the influence of LCL on end-of-century lightning occurrence. For example,
 259 decreases in lightning over the eastern Amazon, West Africa, Siberia, and western Aus-
 260 tralia are more consistent with the r pattern. However, changes in lightning over the west-
 261 ern Amazon, Congo, Indian subcontinent, and China remain more consistent with the
 262 LCL pattern. Alaska is one of the few land regions where the sign change is consistent
 263 between LCL, r , and lightning occurrence. Over ocean, r appears to play an important
 264 role in the lightning decreases over the Southeast Pacific, Caribbean, tropical North At-
 265 lantic, and near-equatorial Indian Ocean.

266 Lastly, CAPE shows end-of-century absolute increases almost everywhere, espe-
 267 cially across the rainy regions of the tropical oceans with most areas increasing 500 to
 268 1250 J kg^{-1} (Figure 2h). These increases are consistent with Romps (2016) who found
 269 that CAPE should increase in a warming climate following the Clausius-Clapeyron re-
 270 lation, and J. Chen et al. (2020) who showed similar CAPE differences globally between
 271 1980-99 and 2081-2100. There are only a few areas in which notable decreases in CAPE
 272 occur: the Southeast Pacific, central Amazon, and Atlantic Ocean along 20°N . While the
 273 largest absolute CAPE changes are projected to occur over the ocean, the oceanic pat-
 274 tern is generally not consistent with the end-of-century lightning changes in Figure 1d,
 275 whereas the relatively smaller CAPE changes over land appear to be more relevant, es-
 276 pecially over the Southeast US, South America, central Africa, and eastern Australia.

277 The difference in standardized interactions between future and present day are plot-
 278 ted in the right column of Figure 2. Note that the interaction terms account for 19% of
 279 the relative importance in the logistic regression, while the individual predictors account
 280 for the other 81% (EB21). Also, all columns are multiplied by -1 since all interactions
 281 have negative coefficients and we still want to represent conditions likely to lead to in-
 282 creases in lightning in red, and decreases in blue. The $\text{LCL} \times r$ interaction results in light-
 283 ning decreasing almost everywhere over land, offsetting the large LCL height increases.
 284 However, most places over oceanic locations would result in a net increase in lightning
 285 from this interaction. $\text{CAPE} \times r$ shows a more variable global signature, while the $\text{CAPE} \times \text{LCL}$
 286 interaction appears to best align with the future lightning changes in CAM5, which is
 287 consistent with EB21 as the $\text{CAPE} \times \text{LCL}$ term is the most important of the three in-
 288 teractions.

289 Figure 2 shows that CAPE, LCL, and r all play an important role in predicting light-
 290 ning in present and future climate scenarios, but large regional variability exists. For ex-
 291 ample, r and CAPE are the most relevant variables over South America (i.e., their end-
 292 of-century predictions are most similar to the overall prediction in Figure 1), while LCL
 293 is the only variable that predicts an increase in lightning over Australia (albeit overly
 294 intense such that the negative predictions from the other variables appear to mute this
 295 overprediction). The interactions improve the predictive potential of the logistic regres-
 296 sion, including helping mitigate some of the overprediction over the ocean that plagues
 297 other lightning parameterizations.

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3.3 Dilute vs Undilute CAPE

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While LCL and r are either direct outputs or found by a straightforward calculation from GCM environmental variables, CAPE has numerous formulations. Undilute CAPE is about an order of magnitude larger than dilute CAPE, so we consider them spanning the range of possible CAPE values. Recall that dilute CAPE is output by CAM5, while undilute CAPE must be calculated but is closer to the CAPE used in previous parameterization studies, including EB21 and R14. We use total precipitation in the following CAPE \times P calculations, but note that two of the four precipitation data sets in Romps et al. (2018) were convective-only. However, the use of convective precipitation doesn't qualitatively change our results.

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Figure 3 shows the change in end-of-century lightning occurrence for the EB21 parameterization and flash rate for CAPE \times P using undilute and dilute CAPE. Similar to Charn and Parishani (2021), we scaled each present-day prediction to match the mean land ISS LIS lightning observations to more fairly compare future changes. The EB21 and R14 parameterizations produce very similar patterns of lightning increases and decreases using undilute CAPE (Figures 3a and c). EB21 produces larger increases in lightning occurrence when using dilute CAPE (Figure 3b), but the pattern of negative and positive changes still strongly resembles the undilute CAPE result in Figure 3a.

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The largest difference occurs when dilute CAPE is used in R14 (Figure 3d). Almost all land regions show end-of-century decreases in flash rate, especially in the tropics. Charn and Parishani (2021) also showed larger decreases in flash rate using dilute CAPE in various formulations of CAPE \times P in a +4 K sea-surface temperature (SST) simulation of a superparameterized version of CAM, although the decreases were not as dramatic as seen here. The sign of change between the EB21 undilute and dilute CAPE results (Figures 3a and b) is more consistent because the predictors are normalized about their mean before being used in the parameterization. The inclusion and interactions with the other environmental inputs also limits large changes due to only one variable.

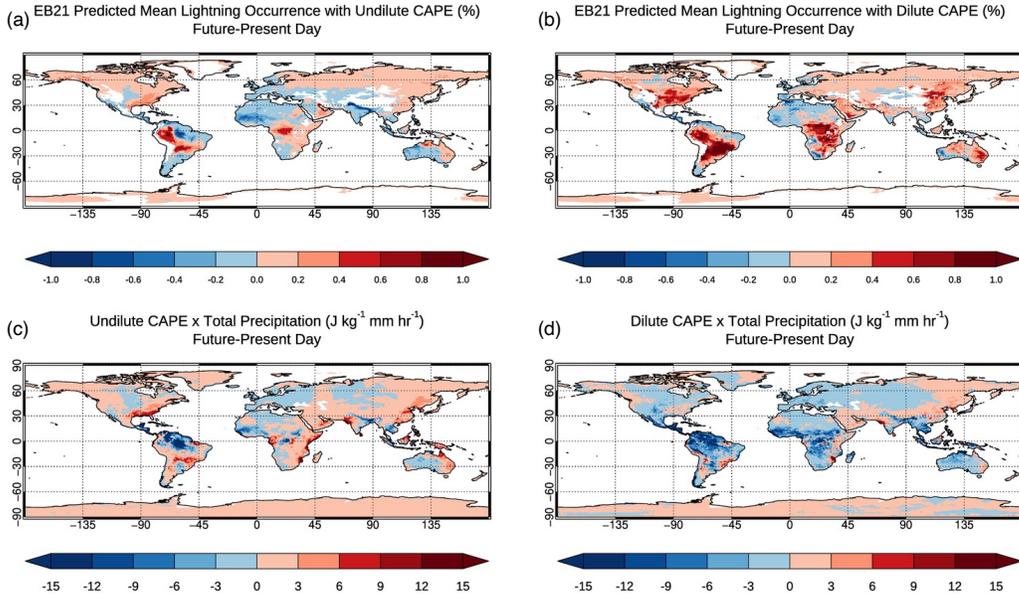


Figure 3. Predictions after present-day scaling to ISS LIS land values of CAM5 end-of-century land-only lightning occurrence (in %) using the EB21 parameterization with (a) undilute and (b) dilute CAPE and flash rate (in $\text{J kg}^{-1} \text{mm hr}^{-1}$) using the R14 CAPE \times P parameterization with (c) undilute and (d) dilute CAPE.

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4 Conclusions

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The EB21 lightning parameterization, which utilizes LCL, CAPE, r , and their interactions, was implemented in CAM5 for present-day (2017-19) and end-of-century (2098-2100) RCP8.5 climate scenarios. Compared to observations from ISS LIS, the CAM5 present-day prediction generally captures the global lightning occurrence pattern but underpredicts lightning over land and overpredicts over the ocean. This is a perennial problem with almost all GCM lightning parameterizations (e.g., Charn & Parishani, 2021; Clark et al., 2017), but the EB21 parameterization produces a better land/ocean lightning ratio than most other schemes when applied to CAM5 fields and does so with a single equation not separated by land and ocean. The land/ocean ratio improves even further when the EB21 parameterization is applied to MERRA2 fields, which can be explained by differences in the individual basic-state predictors. For example, LCLs are higher over land in MERRA-2 compared to CAM5, while the opposite is true over ocean, causing relatively higher lightning occurrence over land for MERRA-2 and over ocean for CAM5. In addition, land and ocean environments are drier in CAM5 for moist environments compared to MERRA-2, causing even further underpredictions of lightning occurrence over land for CAM5, although the drier ocean environments offset the overly high oceanic LCLs to some extent in the EB21 logistic regression formulation.

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The end-of-century lightning projection from CAM5 shows variable increases and decreases over both land and ocean, although higher latitude land regions show across-the-board increases in frequency, which has implications for increased wildfires in locations that typically don't experience much lightning (Y. Chen et al., 2021; Whaley et al., 2024). The large regional variability in positive and negative lightning changes, especially in the tropics, is of significance as many previous studies (e.g., Finney et al., 2018) have found either widespread increases or decreases for tropical lightning activity in a warming climate. The resulting global mean lightning occurrence is projected to slightly

351 decrease by the end of the century, which is consistent with the lower end of the range
 352 of flash rate changes found in Clark et al. (2017) based on results from eight lightning
 353 parameterizations using CAM5 output. When the EB21 parameterization is run with
 354 dilute CAPE instead of undilute CAPE, it provides a more consistent future lightning
 355 prediction than a CAPE \times P parameterization. The EB21 parameterization is simple and
 356 stable to moderate variations in input parameters, providing an attractive alternative
 357 to lightning parameterizations that rely on variables output from convective, cloud, and
 358 microphysics schemes.

359 Acknowledgments

360 This work is supported by NASA Grant NNX17AH66 G S003.

361 5 Open Research

362 5.1 Data Availability Statement

363 ISS LIS data were obtained from NASA GHRC (https://ghrc.nsstc.nasa.gov/lightning/data/data_lis_iss.html) and MERRA-2 data were obtained from NASA GMAO (<https://gmao.gsfc.nasa.gov/reanalysis/MERRA-2/>). The processed CAM5 data is available on the Texas Data Repository (<https://dataverse.tdl.org/dataset.xhtml?persistentId=doi:10.18738/T8/58NOQU>).

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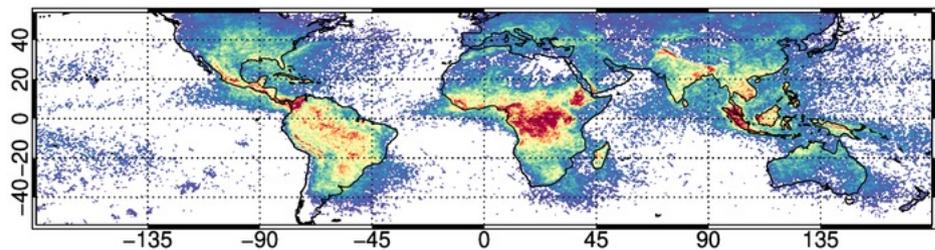
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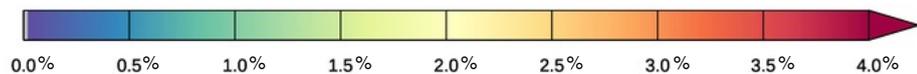
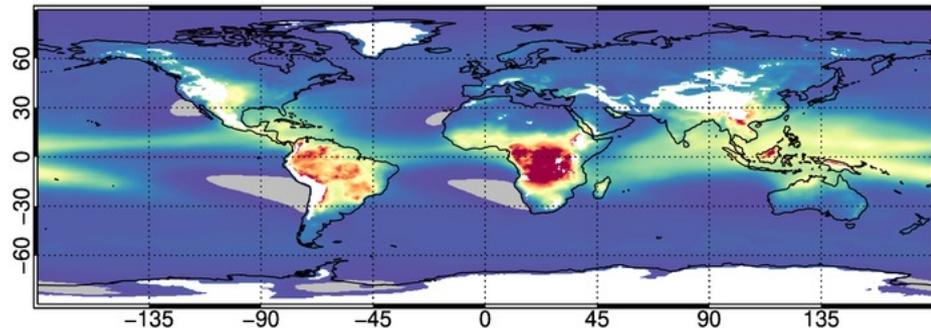
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Figure 1.

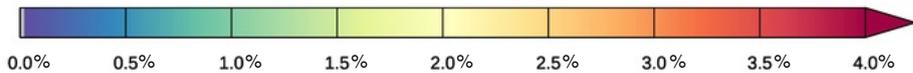
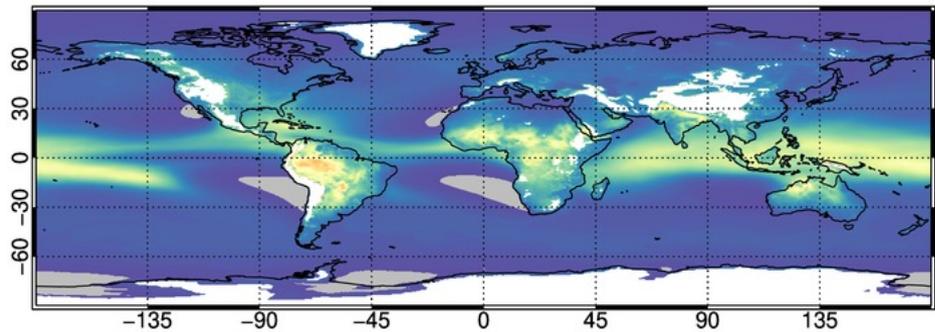
(a) Mean Lightning Occurrence [ISS LIS]



(b) Model **b** Predicted Mean Lightning Occurrence [MERRA2]



(c) Model **b** Predicted Mean Lightning Occurrence [CAM5]



(d) Predicted Mean Lightning Occurrence [CAM5, Future-Present Day]

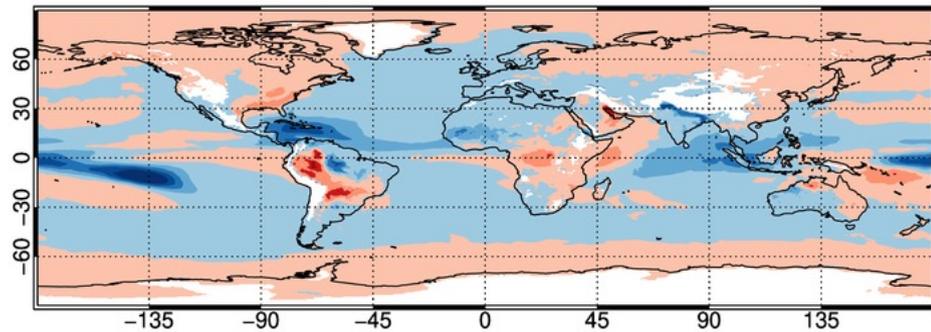


Figure 2.

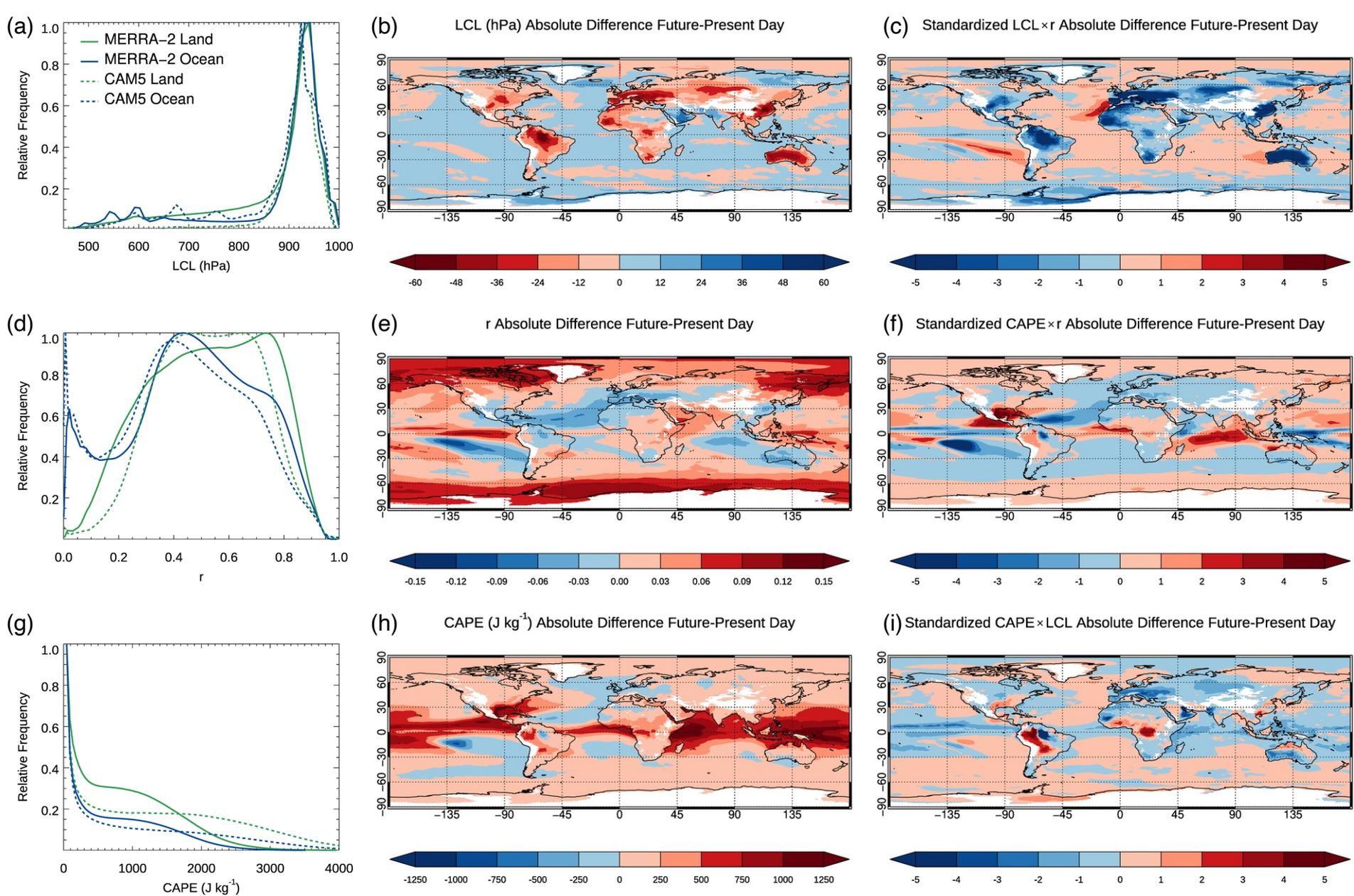
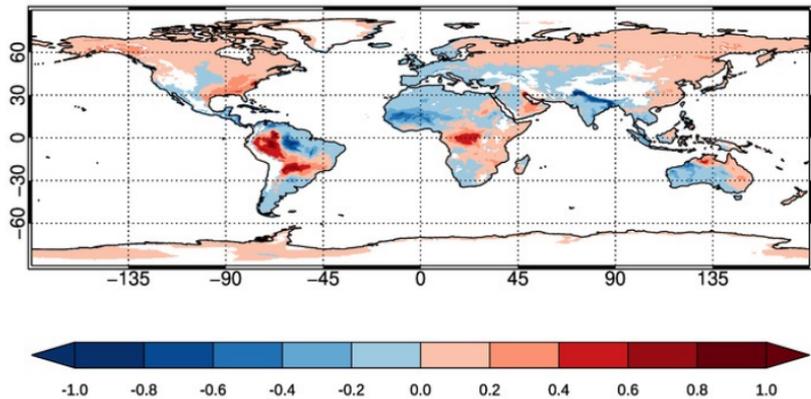
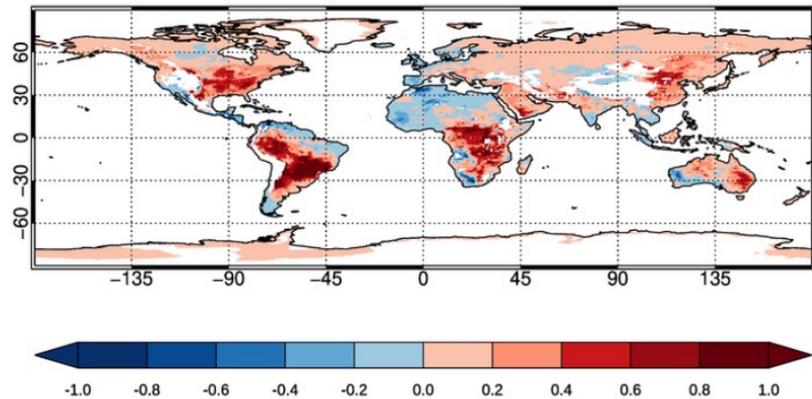


Figure 3.

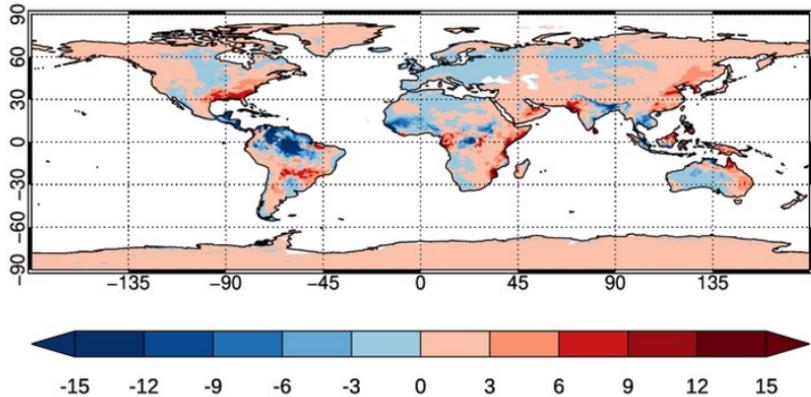
(a) EB21 Predicted Mean Lightning Occurrence with Undilute CAPE (%)
Future-Present Day



(b) EB21 Predicted Mean Lightning Occurrence with Dilute CAPE (%)
Future-Present Day



(c) Undilute CAPE x Total Precipitation ($\text{J kg}^{-1} \text{mm hr}^{-1}$)
Future-Present Day



(d) Dilute CAPE x Total Precipitation ($\text{J kg}^{-1} \text{mm hr}^{-1}$)
Future-Present Day

