

# Upscaling of solute plumes in periodic porous media through a trajectory based spatial Markov model

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## Abstract

We propose an approach to upscale solute transport in spatially periodic porous media. Our methodology relies on pore scale information to predict large scale transport features, including explicit reconstruction of the solute plume, breakthrough curves at fixed distances, and spatial spreading transverse to the main flow direction. The proposed approach is grounded on the recently proposed trajectory-based Spatial Markov model (tSMM), which up-scales transport based on information collected from advective-diffusive particle trajectories across one periodic element. In previous works, this model has been applied solely to one-dimensional transport in a single periodic pore geometry. In this work we extend the tSMM to the prediction of multi-dimensional solute plumes. This is obtained by analyzing the joint space-time probability distribution associated with discrete particles, as yielded by the tSMM. By comparing numerical results from fully resolved simulations and predictions obtained with the tSMM over a wide range of Péclet numbers, we demonstrate that the proposed approach is suitable for modeling transport

of conservative and linearly decaying solute species in a realistic pore space and showcase the applicability of the model to predict steady state solute plumes. Additionally, we evaluate the model performance as a function of numerical parameters employed in the tSMM parameterization.

10 *Keywords:* Upscaling, Porous Media, Transport, Spatial Markov Model,  
11 Transverse dispersion

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## 12 **1. Introduction**

13 Solute transport in porous media is a fundamental problem across many  
14 disciplines, including subsurface geological systems and the performance op-  
15 timization of engineered materials such as filtration membranes. A key chal-  
16 lenge in this context is to obtain accurate predictions at spatial scales much  
17 larger than the ones associated with individual pores without having to re-  
18 solve the physical and chemical processes taking place within complex pore  
19 spaces. This is achieved by upscaled formulations that embed pore scale  
20 features into effective parameters and therefore can be employed to predict  
21 large scale behaviors. To this end, a classical approach is resorting to a  
22 continuum-scale advection-dispersion equation (ADE) [1, 2]. In such a for-  
23 mulation mechanical dispersion induced by pore scale velocity gradients is  
24 modelled through a Fickian-like dispersion term, parameterized via a fourth-  
25 rank dispersivity tensor. The definition of the dispersivity tensor purely  
26 based on pore scale properties presents significant challenges. From a theo-  
27 retical perspective, the solution of three closure problems is required to fully  
28 parameterize solute transport based on pore scale information through vol-  
29 ume averaging [3]. These separate closures are necessary to isolate and char-

acterize the separate effects of diffusion and advection on transport. However, even such a detailed approach may not yield reliable predictions due to a lack of separation of scales, violating the assumptions required by the volume averaging method. In such a case, non-Fickian transport features emerge, particularly at relatively short times and distances [4, 5]. Formally, these effects can still be represented with Eulerian nonlocal (integro-differential) models. In principle these models can be derived by applying upscaling approaches, such as volume averaging, that can relate pore scale geometry and fluid velocities with the emerging transport dynamics through a set of closure differential equations [6]. However, it is often found that resorting to such approaches leads to formidable mathematical and numerical complexity [7, 8], which is associated with *i*) the numerical resolution of various closure problems and *ii*) the approximation of integro-differential equations to obtain the desired large scale outputs.

A specific problem in the context of solute transport upscaling is posed by the modeling of solute plumes, which correspond to the explicit spatial reconstruction of the solute spatial spread at a given time, or at steady state (i.e., under steady state boundary conditions, such as continuous injection). For instance, the analysis of transverse spreading and mixing of steady state solute plumes has great practical relevance in bioremediation and reactive transport scenarios at field and laboratory scales [9, 10, 11]. In these applications the target process is the spreading and mixing of a solute in the direction transverse to a steady flow field characterized by a prevalent direction. Following classical ADE-based descriptions, transport in the transverse direction is typically modeled by introducing a dispersivity parameter. This

55 standard definition typically considers dispersion to be uniquely proportional  
 56 to advective velocity [1]. This formulation was successfully employed, for ex-  
 57 ample to interpret transport and mixing in microfluidic systems characterized  
 58 by relatively simple geometries [12]. However, studies performed in the last  
 59 decade have demonstrated the impact of molecular diffusion on transverse  
 60 dispersion through experiments and numerical simulations [10, 13, 14, 15].  
 61 Such results can be qualitatively linked with analytical and numerical stud-  
 62 ies showing that the dispersion tensor becomes asymmetric in advection-  
 63 dominated scenarios [16, 17]. These studies show that full parameterization  
 64 of the dispersion tensor can become a troublesome task, particularly in media  
 65 characterized by a complex and multi-scale pore structure. Additional levels  
 66 of complexity are introduced when reactive processes are also considered on  
 67 top of pore scale advective-diffusive transport.

68 Over the last decade it has been recognized that pore-to-continuum up-  
 69 scaling of solute transport can often be conveniently obtained by considering  
 70 solute velocities (or associated travel times) over fixed spatial increments by  
 71 means of a Markov chain. This led to the formulation of various flavours  
 72 of so-called Spatial Markov models [e.g. 18, 19, 20, 21]. The SMM is based  
 73 on the calculation of the travel time across a fixed distance and a one step  
 74 correlation existing between successive travel times. By including correla-  
 75 tion the SMM is able to employ information available on a limited portion  
 76 of the system to predict transport across much larger distances. Notably  
 77 such an approach is effective in the presence of advection-dominated scenar-  
 78 ios that become challenging to upscale with classical Eulerian approaches.  
 79 The advantages of employing a spatial Markov approach to obtain the solute

80 breakthrough curve (or first passage time) at a given longitudinal distance has  
 81 been demonstrated in a number of previous works, relying on both numerical  
 82 and laboratory scale experimental datasets [e.g. 18, 21, 22, 23]. Several re-  
 83 cent works have discussed methodologies that employ Lagrangian SMM-like  
 84 approaches to predict solute particles' space-time locations at various scales  
 85 of observations [24, 25]. Yet, to the best of our knowledge, this approach has  
 86 not been applied to the explicit space-time reconstruction of solute plumes  
 87 starting from pore scale properties. In this work we present a methodology to  
 88 upscale transport of solute plumes in the longitudinal and transverse direc-  
 89 tion via a SMM. We consider periodic media, which are routinely considered  
 90 as model porous media in theoretical approaches and are employed in engi-  
 91 neering systems across a wide range of applications [26, 27, 28, 29]. Our work  
 92 starts from a recently proposed trajectory-based SMM (here labelled tSMM)  
 93 to upscale transport, mixing and surface reactions across porous media made  
 94 up of periodic elements [30, 31]. To date, the tSMM has been constrained to  
 95 a highly idealized setting, a periodic wavy channel, which while it displays  
 96 some characteristics of real porous media cannot represent their full com-  
 97 plexity, such as transverse flow or complex pore size distributions. Building  
 98 on this, we analyze longitudinal and transverse transport within a realis-  
 99 tic two-dimensional porous domain. Briefly, the specific objectives of this  
 100 contribution are to *i*) extend the tSMM to the case of a multi-dimensional  
 101 unsteady solute transport and *ii*) yield an efficient and accurate representa-  
 102 tion of transient and steady state solute plumes in porous media based on the  
 103 multi-dimensional tSMM. This second goal entails a specific methodological  
 104 challenge, as steady state plumes are typically computationally expensive to

105 simulate with Lagrangian particle-based approaches. Our objective is the  
 106 development of an upscaled, parsimonious and computationally affordable  
 107 particle-based model for this specific configuration.

## 108 2. Methodology

### 109 2.1. Pore scale setting

110 We consider transport of a passive solute in a periodic two-dimensional  
 111 porous medium. The medium is composed of a collection of periodic unit  
 112 cells, whose geometry is represented in Figure 1. The unit cell is artificially  
 113 created using a stochastic generation procedure as detailed in [21, 32, 33].  
 The cell properties are reported in Table 1. We assume here to deal with

Porosity	$L_x$ [m]	$L_y$ [m]	$\Delta$ [m]	$\ell_S$ [m]
0.631	$4.096 \times 10^{-3}$	$4.096 \times 10^{-3}$	$2 \times 10^{-6}$	$1 \times 10^{-4}$

Table 1: Geometrical characteristics of the unit cell

114  
 115 a semi-infinite periodic porous domain, i.e.  $x \in [0, \infty)$  and  $y \in (-\infty, +\infty)$ .  
 116 Because the medium is composed by an infinite number of identical cells, for  
 117 convenience we define

$$\hat{x} = x - L_x \text{floor} \left( \frac{x}{L_x} \right), \quad \hat{y} = y + \frac{L_y}{2} - L_y \text{floor} \left( \frac{y}{L_y} + \frac{1}{2} \right) \quad (1)$$

118 as a coordinate system referenced to the unit cell,  $\hat{x} \in [0, L_x]$  and  $\hat{y} \in [0, L_y]$ ,  
 119 where  $L_x$  and  $L_y$  define the dimensions of the unit cell in the  $x$ - and  $y$ -  
 120 directions, respectively. The cell is discretized into square pixels of side  
 121  $\Delta = 2 \times 10^{-6}$  m, which for our example results in a unit cell composed of  
 122  $2048 \times 2048$  pixels. The solid and fluid phases are identified by an indicator

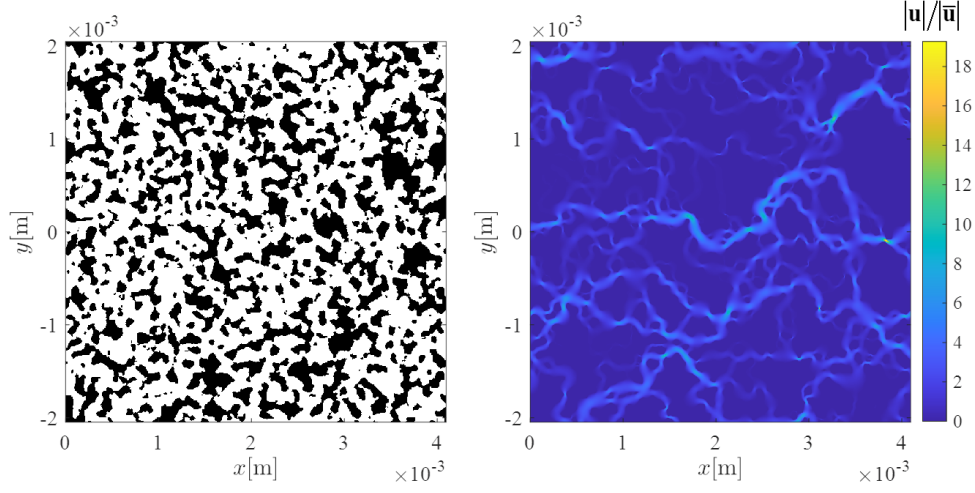


Figure 1: Geometry (left) and normalized velocity magnitude  $|\mathbf{u}|/|\overline{\mathbf{u}}|$  (right) in the unit cell.

field  $I$ , with  $I = 1$  associated with pore space and  $I = 0$  with solid pixels. The correlation length of the indicator field  $I$  is taken as a representative length scale for the pore space and is denoted as  $\ell_S$ . Transport is described by the standard advection-diffusion equation with no flux boundary conditions at the solid-fluid interface

$$\begin{aligned}
 \frac{\partial C(\mathbf{x}, t)}{\partial t} + \nabla \cdot [\mathbf{u}(\mathbf{x})C(\mathbf{x}, t)] &= \nabla \cdot [D\nabla C(\mathbf{x}, t)] \quad \forall \mathbf{x} \in \Gamma_{fluid}, t > t_0 \\
 D \frac{\partial C(\mathbf{x}, t)}{\partial n} &= 0 \quad \forall \mathbf{x} \in \Sigma_{surface}, t > t_0 \\
 C(\mathbf{x}, t_0) &= C_0
 \end{aligned} \tag{2}$$

where  $\mathbf{u} = [u, v]$  is the fluid velocity,  $C$  is solute concentration,  $D$  is the diffusion coefficient,  $C_0$  is the initial concentration distribution. The velocity  $\mathbf{u}$  is obtained by numerically solving the Navier-Stokes equations with

131 OPENFOAM®, release v1712 [34] and diffusion is assumed to be known and  
 132 constant. The velocity is computed assuming periodic boundary conditions  
 133 on the cell boundaries and the no slip condition on the fluid-solid interface.  
 134 We impose a uniform pressure gradient along the  $x$ -direction, labeling  $x, y$   
 135 as longitudinal and transverse directions, respectively. The Péclet number  
 136 associated with transport is calculated as  $Pe = U\ell_s/D$ , where  $U$  is the av-  
 137 erage fluid velocity. In our simulation we set  $D = 10^{-9}$  m<sup>2</sup>/s and we adjust  
 138  $Pe$  by setting  $U$  to the desired value. Note that this is acceptable as our  
 139 simulations are in a Stokes regime, where inertial effects are negligible with  
 140 respect to viscous ones. Transport is solved numerically using a Lagrangian  
 141 particle based random walk method, where the solute plume is discretized  
 142 into a finite number of  $N$  particles. Each particle displaces according to

$$\begin{aligned}
 x_i^{n+1} &= x_i^n + u_i dt + \xi_i \sqrt{2Ddt} \\
 y_i^{n+1} &= y_i^n + v_i dt + \eta_i \sqrt{2Ddt}
 \end{aligned}
 \quad i = 1, \dots, N, \quad (3)$$

143 where  $dt$  is a time step that is constant,  $\xi_i, \eta_i$  are independent identically  
 144 distributed random numbers drawn from normal distributions with zero mean  
 145 and unit variance. We define a reference time step  $dt^*$  according to the  
 146 following criterion  $|\mathbf{d}_{max}| \leq 0.5\Delta$  where

$$|\mathbf{d}_{max}| = \max(|\mathbf{u}|) dt + 2\sqrt{2Ddt^*} \quad (4)$$

147 is an estimate of the maximum displacement. No flux boundary conditions  
 148 at the fluid-solid boundary are imposed as elastic reflections.



## 149 2.2. Spatial Markov Model

150 We upscale transport using the framework of the trajectory-based spatial  
 151 Markov model proposed in [30]. The methodology is based upon pore scale  
 152 transport trajectories and their associated travel times. In the following, we  
 153 first describe the pore scale trajectories simulations and then how these are  
 154 used to parameterize the tSMM.

### 155 2.2.1. Pore scale trajectories

156 Parameterization of the model is grounded on the pore scale simulation of  
 157 a set  $S = \{s_1, \dots, s_{N_s}\}$  of  $N_s$  advective-diffusive trajectories, for a specific  $Pe$ .  
 158 These particle trajectories are simulated by solving Eq. (3) across a single  
 159 cell in the longitudinal direction, i.e., between the inlet location  $x = 0$  and  
 160 the outlet location  $x = L_x$ . Figure 2 represents a sample of 100 trajectories  
 161 across the considered unit cell selected from a flux weighted initial condition  
 162 and setting initial location of particles distributed along the entire unit cell  
 163 cross section.

164 For each trajectory  $s_i$  we record the travel time  $\tau$  needed to travel across a  
 165 distance  $L_x$  in the longitudinal direction and the  $y$  positions ( $y_{in}$ ,  $y_{out}$ ) of the  
 166 particle at the inlet and outlet as the particle enters and exits the domain.  
 167 Particles are injected at locations  $x = 0, y_{in} \in [-L_y/2, L_y/2]$ . Particles may  
 168 cross into adjacent cells along the  $y$  direction, but due to the periodicity  
 169 of the cell geometry each location  $y_{out}$  can be mapped to a corresponding  
 170  $\hat{y}_{out}$  using Eq. (1). Therefore, the coordinate  $y_{out}(s_i)$  can be determined as  
 171  $y_{out}(s_i) = \hat{y}_{out}(s_i) + \Delta_C(s_i)L_y$  where  $\Delta_C(s_i)$  is an integer that indicates the  
 172 net number of cell transitions in the transverse direction observed for a given  
 173  $s_i$  trajectory path. We can then compute  $\Delta y(s_i) = y_{out}(s_i) - y_{in}(s_i)$ . The

174 trajectories are subdivided into  $N_B$  equiprobable bins that are assigned by  
 175 considering the starting locations  $y_{in}(s_i)$  in ascending order. This implicitly  
 176 defines a discretization of the  $\hat{y}$  axis in terms of the binning of the trajectories.  
 177 To exemplify this binning, the trajectories in Figure 2 are subdivided into 10  
 178 bins, indicated by different colours. The trajectories  $s_i$  consider all simulated  
 179 pathways between the locations  $x = 0, y_{in} \in [-L/2, L_y/2]$  and  $x = L_x, y_{out} \in$   
 180  $(-\infty, +\infty)$ . We observe that some trajectories may even travel backwards  
 181 along  $x$  close to the inlet section before traveling downstream, as indicated  
 182 in the highlighted parts in Figure 2. These effects are due to the combined  
 183 action of advection and diffusion and are present for both the considered  $Pe$ .  
 184 The comparison between the two considered cases allows for identification  
 185 of the effects of diffusion on the pore scale trajectory paths. In particular,  
 186 for  $Pe = 100$  particles explore a wider portion of the pore space than for  
 187  $Pe = 1000$ .

### 188 2.2.2. The tSMM parameterization

189 The information collected in the parameterization step is then used to  
 190 build the following trajectory-based Spatial Markov model (tSMM)

$$\begin{aligned}
 x_i^{k+1} &= (k+1)L_x \\
 y_i^{k+1} &= y_i^k + \Delta y [s_i^k | \hat{y}_{out}(s_i^{k-1})] \\
 t_i^{k+1} &= t_i^k + \tau [s_i^k | \hat{y}_{out}(s_i^{k-1})]
 \end{aligned} \tag{5}$$

191 where both  $y_i^{k+1}$  and  $\tau_i^{k+1}$  are determined through a Markov chain, which is  
 192 related to the transverse location assigned to the particle in the periodic cell

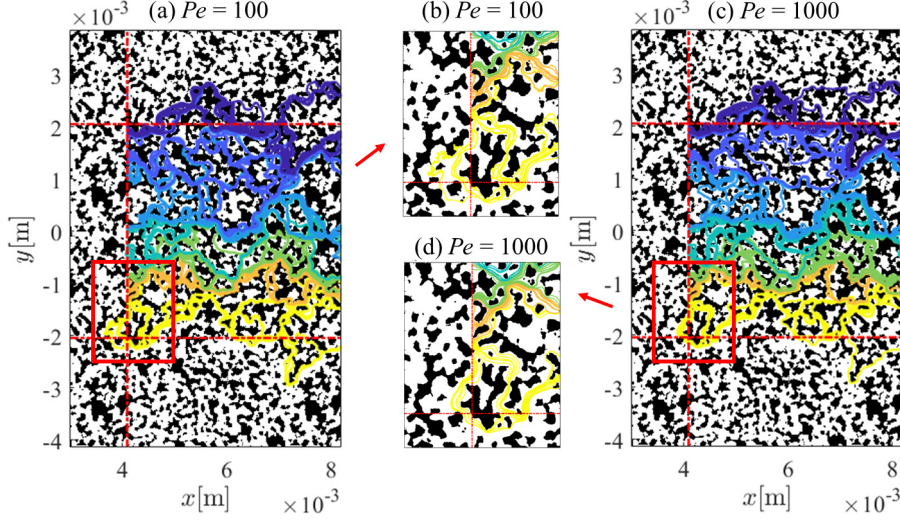


Figure 2: Sample of 100 trajectories employed for model parameterization for (a), (b)  $Pe = 100$  and (c), (d)  $Pe = 1000$ , dashed red lines indicate the unit cell boundaries, the trajectories are binned in different colours as a function of the  $y_{in}$  location. The two middle panels represent a zoom on the region highlighted in red for the two cases.

193 during successive steps  $k$ , leveraging the information given by the trajectories  
 194 in  $s_i$ . The innovative feature of the model in Eq. (5) with respect to previous  
 195 implementations [30, 21] is that it allows for predictions of transverse spread-  
 196 ing over successive Markov steps. This is achieved by considering  $y$  as a con-  
 197 tinuous variable, i.e., the Markov chain has a longitudinal fixed spatial step  
 198  $L_x$  while transport along  $y$  is considered through the  $\Delta y(s_i)$  obtained from  
 199 the trajectories  $s_i$  recorded during the parameterization stage. In essence, at  
 200 the beginning of the simulation (step  $k = 0$ ) each particle  $i$  is assigned to an  
 201 initial location  $y_i^0$  corresponding to a selected initial or boundary condition  
 202 (e.g., flux weighted or uniform distribution on the inlet boundary). From  
 203 this information we select a trajectory  $s_i^1$ , randomly sampling from those  
 204 whose  $y_{in}(s_i)$  lies in the same bin interval as  $y_i^0$ . By selecting the trajectory

we also obtain a given travel time  $\tau(s_i^1)$  and transverse displacement  $\Delta y(s_i^1)$ , from which we evaluate  $\hat{y}_{out}(s_i^1)$ . The latter can be then used to select a new trajectory  $s_i^2$  for the next transition and the procedure can then be repeated for any arbitrary step number  $k > 0$ .

### 2.3. Model outputs

To analyze the outputs of our tSMM model, our analysis relies on the following dimensionless space-time coordinates

$$\tilde{x} = \frac{x^k}{L}, \quad \tilde{y} = \frac{y_c}{L}, \quad \tilde{t} = \frac{tU}{L} \quad (6)$$

where  $L = L_x = L_y$ ,  $x^k$  corresponds to the longitudinal spatial coordinate of  $k^{th}$  Markov step (see Eq. (5)) and  $y_c(x) = y(x) - \bar{y}(x)$ , i.e., is the transverse location centered with respect to the average transverse position  $\bar{y}(x)$  observed at a given  $x$ . The value of  $\bar{y}(x)$  is not constant with  $x$  because the average transverse velocity component is not exactly equal to zero. While this component is only approximately 1% of the longitudinal mean velocity  $U$ , it still induces plume migration along  $y$  after a number of cells.

The key output of the tSMM is the joint probability distribution  $P(\tilde{x}, \tilde{y}, \tilde{t})$ . In this distribution, the variable  $\tilde{x}$  can only assume discrete values, while the  $\tilde{y}, \tilde{t}$  are continuous.

Physically meaningful information related to the plume can then be extracted from this joint probability by considering conditional and marginal distributions. In our analysis we will consider the conditional distributions  $P(\tilde{t}, \tilde{y}|\tilde{x})$  for a given dimensionless downstream distance  $\tilde{x}$ , or  $P(\tilde{x}, \tilde{y}|\tilde{t})$  for a given dimensionless time,  $\tilde{t}$ . Examples of these conditional distributions

227  $P(\tilde{t}, \tilde{y}|\tilde{x})$  for  $\tilde{x} = 5, 10, 25$  and  $50$  and  $P(\tilde{x}, \tilde{y}|\tilde{t})$  for  $\tilde{t} = 20$ , obtained from the  
 228 tSMM are shown in Figure 3a and b for  $Pe = 100$  and  $1000$ , respectively.

229 In addition, we consider the marginal probability distributions  $P(\tilde{x}, \tilde{y})$ ,  
 230 and  $P(\tilde{t}|\tilde{x})$ ,  $P(\tilde{y}|\tilde{x})$  conditional to a given dimensionless downstream distance.  
 231 These distributions have a clear physical meaning: the marginal distribution  
 232  $P(\tilde{x}, \tilde{y})$  represents the steady state distribution of the particle plume, while  
 233  $P(\tilde{t}|\tilde{x})$  corresponds to the breakthrough curve, i.e., the first passage time  
 234 probability distribution at distance  $\tilde{x}$ . Finally, the probability distribution  
 235  $P(\tilde{y}|\tilde{x})$  provides the probability distribution associated with transverse po-  
 236 sition at a control plane and is related solely to transport in the transverse  
 237 direction.

238 To produce benchmark data against which to test the tSMM we run  
 239 a 50 cell high resolution random walk direct numerical simulation (DNS)  
 240 using transport Eq. (3). The accuracy of the tSMM defined in section 2.2  
 241 will be tested by comparing the above mentioned probability distributions  
 242 with their analogs obtained from the DNS. For each of the above defined  
 243 distributions we provide a quantitative evaluation of the mismatch between  
 244 DNS and tSMM using the Hellinger distance [35]

$$HD[F_1, F_2] = \frac{1}{\sqrt{2}} \sqrt{\sum_{i=1}^N \left( \sqrt{f_{1,i}} - \sqrt{f_{2,i}} \right)^2} \quad (7)$$

245 where  $F_1$  is any of the above-mentioned marginal or conditional distribu-  
 246 tions predicted by DNS and  $F_2$  corresponds to its counterpart obtained with  
 247 the tSMM. These distributions are approximated through  $N$  discrete bins  
 248 and  $f_{1,i}$ ,  $f_{2,i}$  are the values of the distributions in the  $i^{th}$  bin.

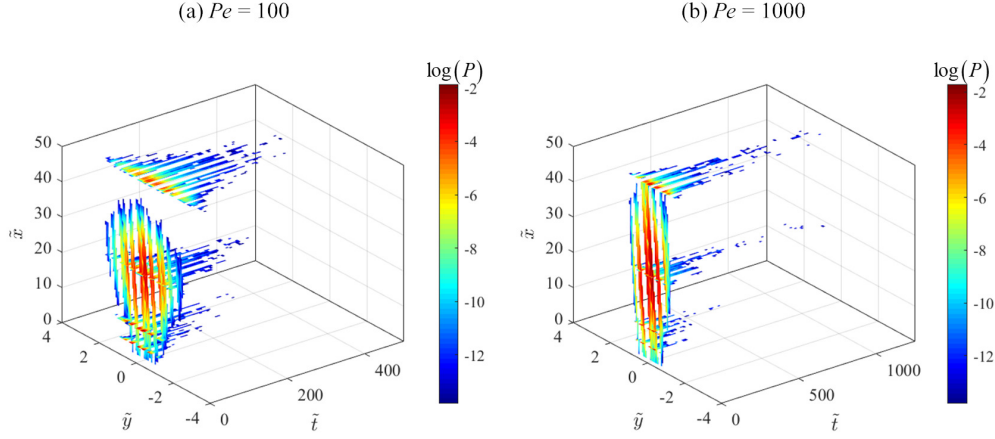


Figure 3: Conditional joint distributions  $P(\tilde{t}, \tilde{y}|\tilde{x})$  for  $\tilde{x} = 5, 10, 25, 50$  and  $P(\tilde{x}, \tilde{y}|\tilde{t})$  for  $\tilde{t} = 20$  predicted by the tSMM considering (a)  $Pe = 100$  and (b)  $Pe = 1000$

249 The  $HD$  metric quantifies the distance between two probability measures  
 250 and it is a proper distance metric in the mathematical sense, by satisfying  
 251 the properties of non-negativity, symmetry, and triangle inequality.  $HD$  is  
 252 also bounded between 0 and 1, where 0 means that the two distributions are  
 253 indiscernible and 1 that they are maximally distant.

### 254 3. Results

255 First we show results obtained considering parameterization of the tSMM  
 256 with  $N_s = 10^6$ ,  $N_B = 100$ ,  $dt^* = 10^{-5}$  s and  $10^{-6}$  s for  $Pe = 100$  and  
 257  $Pe = 1000$ , respectively. The time step  $dt^*$  indicates the values of  $dt$  eval-  
 258 uated according to Eq. (4). Then, in section 3.3 we analyze the impact of  
 259 parameters  $N_B$  and  $dt$  on the accuracy of the tSMM. For all cases, including  
 260 the DNS reference simulation and tSMM, we impose a flux weighted bound-  
 261 ary conditions. Note that In the reference DNS simulation the  $dt$  parameter  
 262 is kept constant and equal to  $dt^*$ .

### 263 3.1. Model performance as a function of $Pe$

264 The tSMM is able to replicate the shape of the reference conditional dis-  
 265 tribution  $P(\tilde{t}, \tilde{y}|\tilde{x})$  obtained from the DNS for both investigated Péclet num-  
 266 bers. Figure 4a shows the joint distribution  $P(\tilde{t}, \tilde{y}|\tilde{x})$  for  $\tilde{x} = 25$  and  $Pe = 100$   
 267 from the high resolution direct numerical simulations and corresponding re-  
 268 sults obtained with the tSMM (Figure 4b). The agreement between the two  
 269 solutions is significant for all transverse coordinates,  $\tilde{y}$ , and dimensionless  
 270 travel times,  $\tilde{t}$ . Analogous results are obtained for  $Pe = 1000$  (see Figure 4c  
 271 and d) and for all other investigated Markov steps (not shown). For both  $Pe$   
 272 values the maximum value of the probability distributions is found at  $\tilde{y} \approx 0$ .  
 273 Note that  $\tilde{y}$  locations associated with zero probability across the whole time  
 274 window correspond to the occurrence of solid along the considered transverse  
 275 section.

276 To quantify the accuracy of the tSMM outputs with respect to the ref-  
 277 erence DNS, Figure 5 shows the metric  $HD[P_{DNS}(\tilde{y}, \tilde{t}|\tilde{x}); P_{tSMM}(\tilde{y}, \tilde{t}|\tilde{x})]$  de-  
 278 fined in Eq. (7) for both investigated  $Pe$  numbers and all Markov steps. We  
 279 note that the distance between the DNS and the tSMM distributions slightly  
 280 increases with  $\tilde{x}$  and is generally larger for  $Pe = 100$  than for  $Pe = 1000$ .  
 281 This result is likely due to the fact that the effect of noise in low probability  
 282 values increases with the strength of diffusion.

283 Figure 6 depicts conditional joint distribution  $P(\tilde{x}, \tilde{y}|\tilde{t})$  for  $\tilde{t} = 20$ , corre-  
 284 sponding to the time dependent solute plume. DNS and tSMM predictions  
 285 are shown for  $Pe = 100$  (see Figure 6a, b) and for  $Pe = 1000$  (see Figure  
 286 6c, d). Again, the tSMM is able to capture all essential features displayed  
 287 by the fully resolved simulations. Note that tSMM allows for predictions

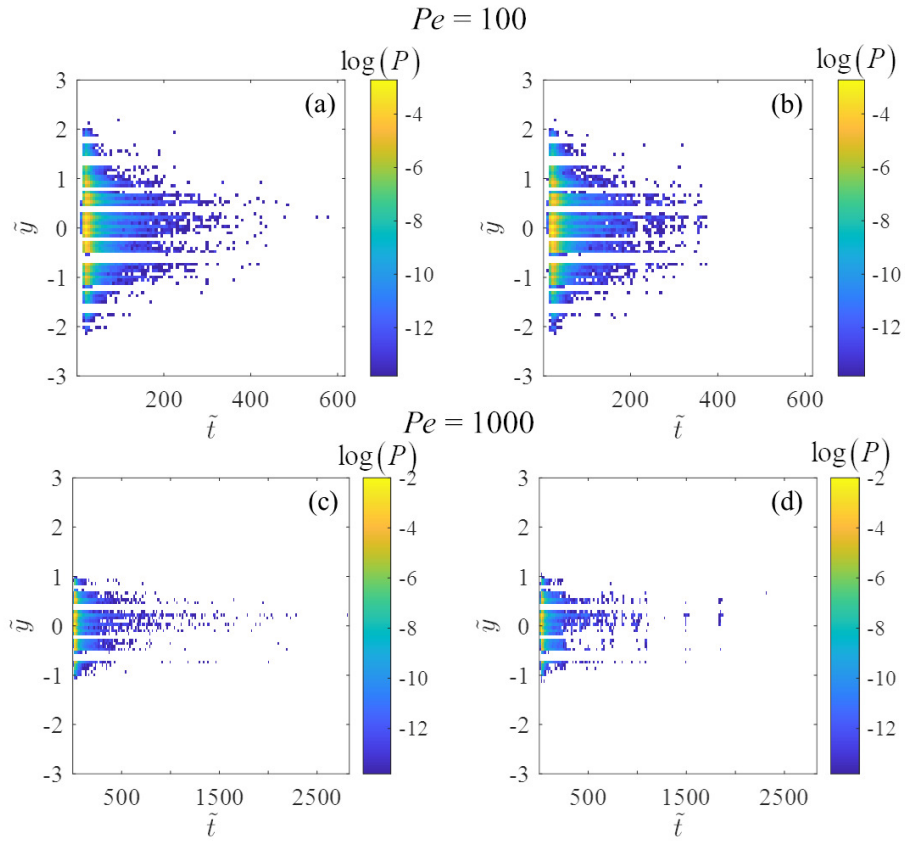


Figure 4: Conditional joint distributions  $P(\tilde{t}, \tilde{y} | \tilde{x})$  for  $\tilde{x} = 25$  and  $Pe = 100$  obtained with (a) DNS, (b) tSMM and for  $Pe = 1000$  predicted by (c) DNS and (d) tSMM.



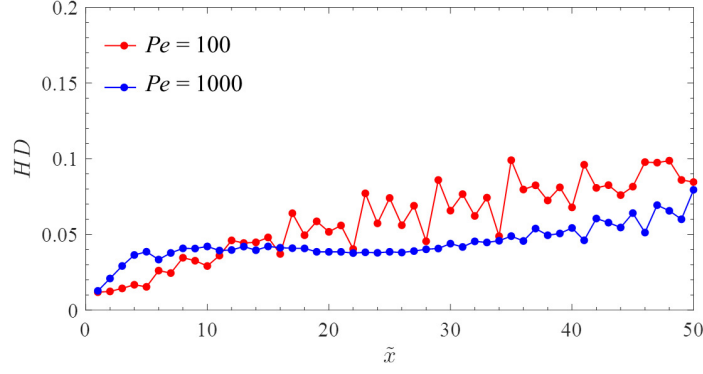


Figure 5:  $HD[P_{DNS}(\tilde{y}, \tilde{t}|\tilde{x}); P_{tSMM}(\tilde{y}, \tilde{t}|\tilde{x})]$  as a function of the downstream location  $\tilde{x}$  for  $Pe = 100$  (red) and 1000 (blue).

288 of the evolution of the plume in the longitudinal and transverse directions  
 289 employing a significantly smaller computational effort than the DNS. As an  
 290 example, the computational time for running the tSMM is approximately 1%  
 291 of that one needed for the DNS results when considering 50 Markov steps.  
 292 Note that this percentage decreases for simulation across higher numbers of  
 293 unit cells (i.e., the computational gain increases with the dimension of the  
 294 system of interest).

295 The marginal distributions  $P(\tilde{x}, \tilde{y})$  are depicted in Figure 7a and b for  
 296  $Pe = 100$  and 1000 respectively. As mentioned above, these distributions  
 297 identify the steady-state plume for a nonreactive solute. These distributions  
 298 are here obtained at no additional computational cost with respect to the  
 299 transient case, which for a DNS would not be the case and significant addi-  
 300 tional cost would be required.

301 Figure 8 displays breakthrough curves  $P(\tilde{t}|\tilde{x})$  considering travel distances  
 302  $\tilde{x} = 5, 10, 25$  and 50 from the injection location and provides a quantitative  
 303 comparison between the reference DNS and the tSMM results. We note that

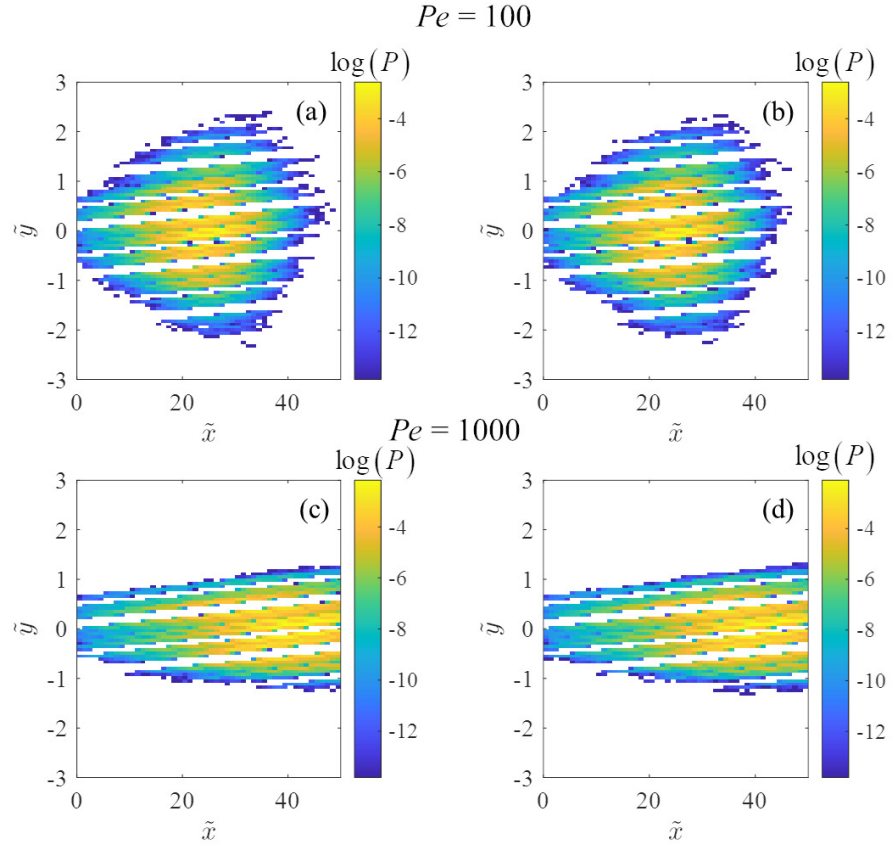


Figure 6: Conditional distributions  $P(\tilde{x}, \tilde{y}|\tilde{t})$  for  $\tilde{t} = 20$  and  $Pe = 100$  obtained with (a) DNS (b) tSMM and for  $Pe = 1000$  predicted by (c) DNS and (d) tSMM

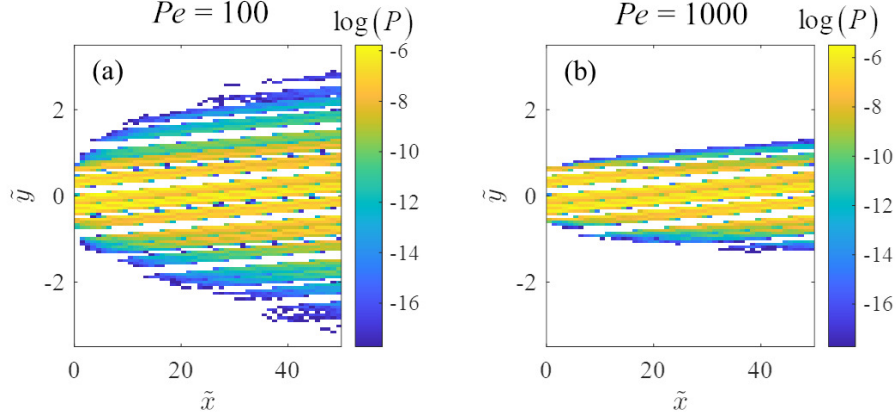


Figure 7: Marginal distribution  $P(\tilde{x}, \tilde{y})$  for (a)  $Pe = 100$  and (b)  $Pe = 1000$ .

the tSMM can reproduce the breakthrough curves across a wide range of distances and both Péclet numbers. This result shows that the trajectory-based upscaled model accurately predicts arrival times in a porous medium made of periodic unit cells displaying a disordered geometry and is in line with those obtained within simpler geometrical settings [30].

The comparison between tSMM and DNS marginal distribution of transverse locations  $P(\tilde{y}|\tilde{x})$  is shown in Figure 9 for two selected distances from the injection ( $\tilde{x} = 5$  and  $\tilde{x} = 25$ ) and for both investigated  $Pe$  numbers. To compare the spreading of the particle plume over all Markov steps we consider the standard deviation of distribution  $P(\tilde{y}|\tilde{x})$  as a function of  $\tilde{x}$  (see Figure 10).

Results obtained through the tSMM are in close agreement with those yielded by the DNS. We observe that the change in  $Pe$  has marked effects on transverse spreading of the solute, as has been previously observed in laboratory and numerical studies [10, 13]. In particular, the standard deviation  $\sigma_y$

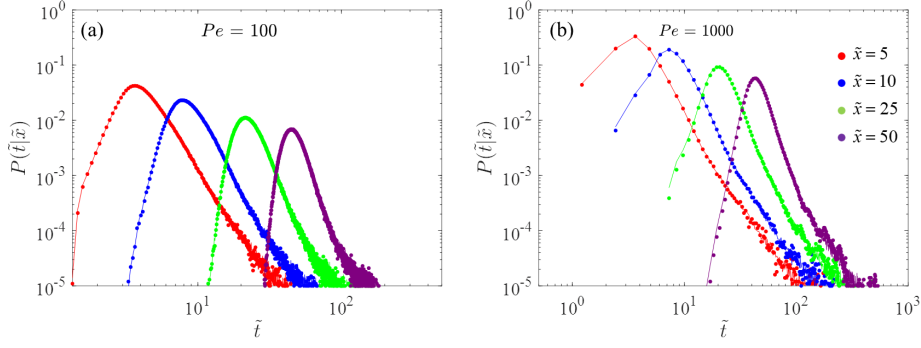


Figure 8: Breakthrough curves  $P(\tilde{t}|\tilde{x})$  obtained at control planes located at distances  $\tilde{x} = 5, 10, 25, 50$  unit cells for (a)  $Pe = 100$  and (b)  $Pe = 1000$ . Symbols and lines represent the DNS and tSMM results, respectively.

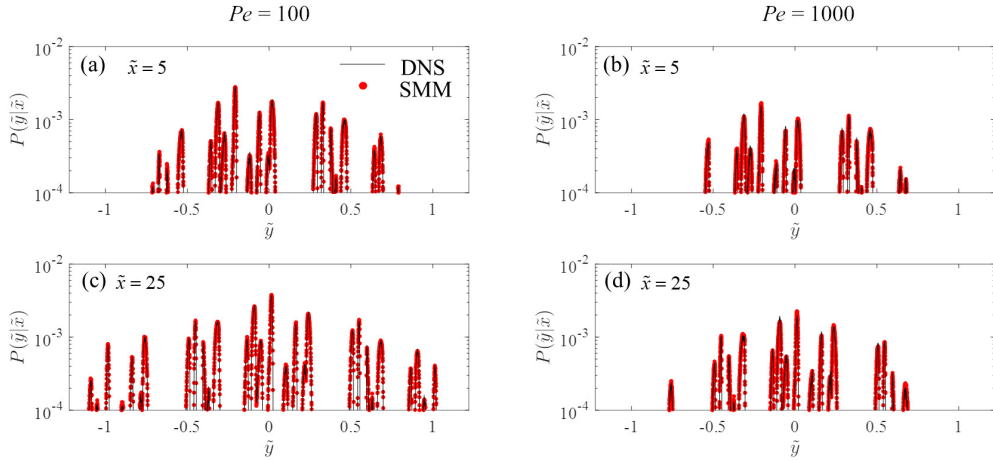


Figure 9: Comparison between the marginal distributions  $P(\tilde{y}|\tilde{x})$  as given by direct numerical simulation (black lines) and tSMM (red dots) for  $\tilde{x} = 5$  and a)  $Pe = 100$ , b)  $Pe = 1000$ , for  $\tilde{x} = 25$  and c)  $Pe = 100$  and d)  $Pe = 1000$ .

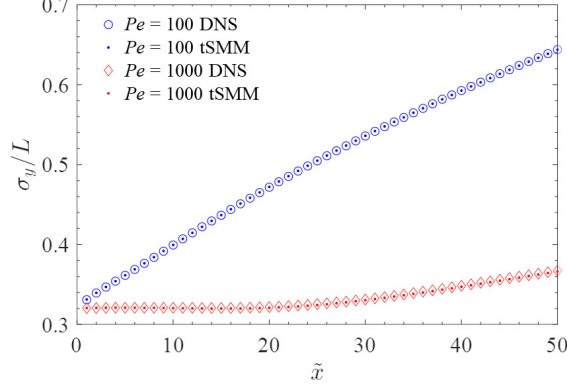


Figure 10: Comparison of the standard deviation,  $\sigma_y$ , of the distribution  $P(\tilde{y}|\tilde{x})$  for the DNS and tSMM as a function of  $\tilde{x}$ .

continuously increases as a function of the longitudinal distance traveled for  $Pe = 100$ . On the contrary,  $\sigma_y$  is approximately constant up to  $\tilde{x} = 20$  for  $Pe = 1000$  and then starts increasing. This result implies that for such an advection-dominated situation we only observe significant transverse spreading after particles have traveled a distance of 20 cells. This result is due to the converging-diverging nature of advective streamlines in two-dimensional fields, and this particular behavior might be different if investigated in three-dimensions. Yet, the tSMM is able to predict these different dynamics based on the simulation of transport across a single unit cell. Note also that the methodology can be extended to three dimensions, upon relying on the same procedure described in Section 2.2.

### 3.2. Steady state plumes with first order degradation

As a showcase application of the capabilities of the tSMM, we also evaluate the influence of a first order reaction on pinching off the steady-state plume. We assume in this application that the solute undergoes degrada-

tion following linear kinetics. This is accounted for in a straightforward manner in the tSMM framework. Starting from the conservative plume results (see Figure 7), for each particle we define the probability of reaction as  $R_i(t, \lambda) = 1 - e^{-\lambda t}$ , where  $\lambda$  is the kinetic degradation rate. Then we compare  $R_i$  with a random number,  $U_i$ , drawn from a standard uniform distribution. If  $U_i \geq P_i$  no reaction occurs while if  $U_i < P_i$  the particle is removed from the system. In our examples  $\lambda$  is chosen based on obtaining specific values of Damkhöler numbers,  $Da = \frac{\ell_s^2 \lambda}{D}$ .

We compute marginal distributions  $P(\tilde{x}, \tilde{y})$  to represent the steady state plume for this reactive scenario. The results obtained for  $Da = 1$  and 5 (corresponding to  $\lambda = 0.1$  and 0.5) and  $Pe = 100$  are depicted in Figures 11a and b while Figures 11c and d show the case associated with  $Da = 5$  and 10 (corresponding to  $\lambda = 0.5$  and 1) and  $Pe = 1000$ . These results portray the ability of our proposed tSMM to predict transport in longitudinal and transverse direction while also accounting for a reactive solute undergoing a first order reaction process. Note that the results are obtained at negligible additional computational cost with respect to the unsteady state, conservative, transport simulations.

### 3.3. Error analysis

To provide a quantitative description of the influence of parameters  $N_B$  and  $dt$  on the accuracy of the proposed spatial Markov model we evaluate  $HD[F_1, F_2]$ , see Eq. 7, choosing as  $F_1$  and  $F_2$  the marginal distributions of travel times,  $P(\tilde{t}|\tilde{x})$ , or transverse positions,  $P(\tilde{y}|\tilde{x})$ , conditional to a given downstream location  $\tilde{x}$  from the injection point, evaluated with DNS and tSMM respectively. Note that the time step is kept constant and equal to

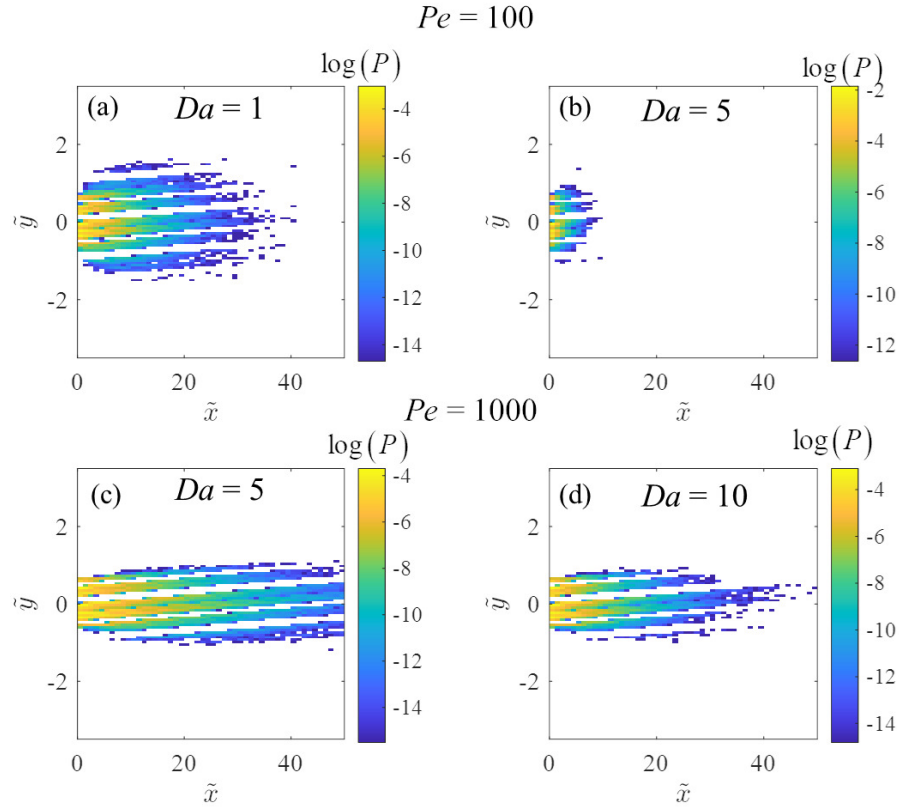


Figure 11: Marginal distribution  $P(\tilde{x}, \tilde{y})$  for a reactive contaminant undergoing a degradation following a linear kinetics for  $Pe = 100$  (a)  $Da = 1$ , (b)  $Da = 5$ , and  $Pe = 1000$  (c)  $Da = 5$ , (d)  $Da = 10$ .

$dt^*$  in the DNS, while we consider in following different values for the simulation of the trajectories employed to parameterize the tSMM, To simplify the notation we introduce here the following indicators

$$HD_t = HD[P_{DNS}(\tilde{t}|\tilde{x}); P_{SMM}(\tilde{t}|\tilde{x})] \quad (8)$$

$$HD_y = HD[P_{DNS}(\tilde{y}|\tilde{x}); P_{SMM}(\tilde{y}|\tilde{x})] \quad (9)$$

353 to assess the model errors. The analysis is performed considering both  
 354  $Pe = 100$  and  $Pe = 1000$  and considering different values of  $N_B$  and  $dt$   
 355 for the parameterization of the tSMM, while, as mentioned above, in the ref-  
 356 erence DNS the  $dt$  is constant and equal to  $dt^*$ . In order to provide an overall  
 357 assessment of the impact of parameters  $dt$  and  $N_B$  we focus on the average  
 358 of  $HD_\alpha$  (with  $\alpha = t, y$ ) across all 50 investigated Markov steps (the averag-  
 359 ing operator is denoted by the symbol  $\langle \cdot \rangle$ ). Figures 12a and b show  $\langle HD_t \rangle$   
 360 and  $\langle HD_y \rangle$  as a function of  $dt/dt^*$  and  $N_B$ , respectively. Continuous lines  
 361 correspond to  $Pe = 100$ , while dashed lines depict results associated with  
 362  $Pe = 1000$ . Red and blue colors are related to arrival time and transverse  
 363 location distributions, respectively. Note that for  $Pe = 100$  both  $\langle HD_t \rangle$  and  
 364  $\langle HD_y \rangle$  are not very sensitive to the choice of parameter  $dt$  (see Figure 12a ).  
 365 On the contrary for  $Pe = 1000$  we observe a sharp increase of  $\langle HD_t \rangle$  and  
 366  $\langle HD_y \rangle$  for  $dt > 2dt^*$ . This is probably due to the fact that advective particle  
 367 displacements depend linearly on  $dt$ , as opposed to the diffusive ones which  
 368 scale with  $dt^{0.5}$ . The variation of  $\langle HD_t \rangle$  and  $\langle HD_y \rangle$  as a function of  $N_B$  are  
 369 displayed in Figure 12b for the two investigated  $Pe$  numbers. These results  
 370 show that the quality of model predictions deteriorates for decreasing num-



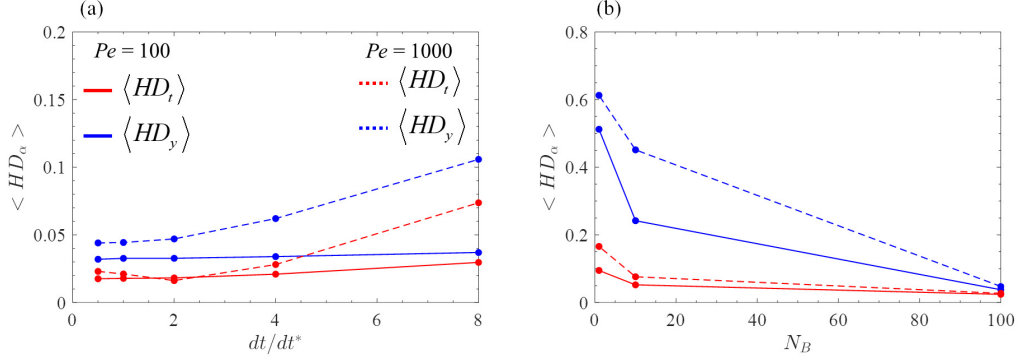


Figure 12:  $\langle HD_\alpha \rangle$  for  $Pe = 100$  and  $Pe = 1000$  as a function of the Markov step number (cell number) for different  $dt$  and numbers of bins  $N_B$  employed in the parameterization step.

ber of bins associated with the tSMM parameterization. Note that  $N_B = 1$  corresponds to considering the particle trajectories as totally uncorrelated across successive Markov steps, while setting  $N_B > 1$  in the tSMM parameterization allows consideration of correlation between particle trajectories belonging to the same bin. We note that considering uncorrelated particle trajectories does not provide good agreement between DNS and tSMM distributions, as indicated by high values of the  $HD_\alpha$  metrics. Also in this case correlation effects at a fixed downstream distance become stronger as Péclet number increases, in line with the results of previous studies [22, 31].

#### 4. Discussion and conclusions

Our study proposes a methodology for upscaling solute plumes in periodic porous media through a trajectory based spatial Markov Model. We extend the work of [30] to the case of a multi-dimensional unsteady solute transport and exemplify our approach considering a two-dimensional porous medium with a disordered geometry. Our framework is based on the simula-

tion of advection-diffusion random walk particle trajectories across a single periodic flow cell with the aim of predicting transport over a much larger scale. In particular, our analysis explicitly includes the evaluation of the joint space-time probability distributions associated with solute plumes providing an efficient and accurate representation of both transient and steady state transport in porous media for different  $Pe$  numbers. Our work leads to the following major conclusions:

1. From a comparison with high resolution direct numerical simulations we show that the proposed tSMM accurately predicts spatial and temporal distributions of a conservative solute plume using information collected from a single cell simulation. The current framework is not restricted to a simplified geometry setting or a particular flow condition, but can be employed to accurately predict multi-dimensional transport in a realistic two-dimensional pore space once the flow field has been evaluated. Note that, in principle our methodology can be used with different type of initial injection condition, e.g. flux-weighted or uniformly distributed, pulse or continuous injection, and can be extended to a three-dimensional setup.
2. Our model is able to predict different transport dynamics, particularly regarding the influence of  $Pe$  on transverse plume spreading. Our results are in line with previous laboratory and numerical studies [10, 13]. In particular, the change in  $Pe$ , due to a change in the diffusion coefficient of the compound, had marked effects on transverse spreading of the solute and this is well captured by our upscaling approach.
3. Particle trajectories and associated travel times, which were simulated

with the proposed tSMM to predict conservative transport, can be naturally extended to simulate reactive transport processes with negligible additional computational cost. As an example, in this work we analyse the influence of a first order kinetic reaction on a steady state plume. To do so our model has been coupled with a probabilistic representation of a linear degradation reaction and applied for several  $Da$  numbers. Note that the methodology is already fully compatible with the analysis of other types of reactions (e.g. sorption/desorption) as was shown by [31] for an idealized benchmark problem.

4. The effect of tSMM parameterization (in particular the parameters  $dt$  and  $N_B$ ) was studied for the two analysed  $Pe$ . As expected, the difference between tSMM and DNS distribution both in space and time increases with increasing  $dt$  and decreasing  $N_B$  respectively. A marked effect of parameterization was observed for  $Pe = 1000$  with respect to  $Pe = 100$  due to fact that (i) the contribution of the advective transport, which increases with  $Pe$ , is more affected by the choice of  $dt$  employed in the tSMM parameterization and (ii) the relevance of particle trajectory correlation increases with  $Pe$ .

For all the above points the simulation time needed for the tSMM is negligible if compared to the one required for high resolution direct numerical simulation. This is one of the great advantages of the proposed tSMM which allows predicts of multi-dimensional transport across large distances (for both conservative and reactive solutes) without the burden of excessive computational resources. Note that, in its current form, the model can be applied to a spatially periodic domain. This restriction is similar to the fact that

436 solving a closure problem on a periodic unit cell is required for many well es-  
437 tablished upscaling procedures. At the same time, broadening the scenarios  
438 of interest, for example extending the methodology to upscaling transport of  
439 a conservative and/or reactive solute in a disordered non periodic porous do-  
440 main would constitute additional elements of interest which are compatible  
441 with the approach we rest upon. A first attempt in this direction has been  
442 provided by [21, 36] obtaining promising results which can be advanced in  
443 the context of future investigations.

## 444 **Acknowledgments**

## 445 **Data Availability Statement**

446 Data sets for this research are available online at  
447 <https://data.mendeley.com/datasets/rzg53tn963/draft>

448 **References**

- 449 [1] A. Scheidegger, Statistical hydrodynamics in porous media, *Journal of*  
450 *Applied Physics* 25 (8) (1954) 994–1001. doi:10.1063/1.1721815.
- 451 [2] A. H.-D. Cheng, J. Bear, *Modeling Groundwater Flow and Contaminant*  
452 *Transport*, Springer Publishing Company, 2016.
- 453 [3] F. Valdés-Parada, D. Lasseux, F. Bellet, A new formulation of the dis-  
454 *persion tensor in homogeneous porous media*, *Advances in Water Re-*  
455 *sources* 90 (2016) 70–82. doi:10.1016/j.advwatres.2016.02.012.
- 456 [4] J. Salles, J.-F. Thovert, R. Delannay, L. Prevors, J.-L. Auriault,  
457 P. Adler, Taylor dispersion in porous media. determination of the dis-  
458 *persion tensor*, *Physics of Fluids A* 5 (10) (1992) 2348–2376.
- 459 [5] B. Berkowitz, A. Cortis, M. Dentz, H. Scher, Modeling non-fickian trans-  
460 *port in geological formations as a continuous time random walk*, *Rev.*  
461 *Geophys.* 44 (2).
- 462 [6] B. Wood, F. Valdés-Parada, Volume averaging: Local and nonlocal clo-  
463 *sures using a green’s function approach*, *Advances in Water Resources*  
464 51 (2013) 139–167. doi:10.1016/j.advwatres.2012.06.008.
- 465 [7] Y. Davit, B. D. Wood, G. Debenest, M. Quintard, Correspondence be-  
466 *tween one-and two-equation models for solute transport in two-region*  
467 *heterogeneous porous media*, *Transport in porous media* 95 (1) (2012)  
468 213–238.

- 469 [8] G. Porta, G. Ceriotti, J.-F. Thovert, Comparative assessment of  
470 continuum-scale models of bimolecular reactive transport in porous me-  
471 dia under pre-asymptotic conditions, *Journal of contaminant hydrology*  
472 185 (2016) 1–13.
- 473 [9] O. Cirpka, A. Valocchi, Two-dimensional concentration distribution for  
474 mixing-controlled bioreactive transport in steady state, *Advances in Wa-  
475 ter Resources* 30 (6-7) (2007) 1668–1679. doi:10.1016/j.advwatres.  
476 2006.05.022.
- 477 [10] G. Chiogna, C. Eberhardt, P. Grathwohl, O. Cirpka, M. Rolle, Evidence  
478 of compound-dependent hydrodynamic and mechanical transverse dis-  
479 persion by multitracer laboratory experiments, *Environmental Science  
480 and Technology* 44 (2) (2010) 688–693. doi:10.1021/es9023964.
- 481 [11] Y. Tang, C. Werth, R. Sanford, R. Singh, K. Michelson, M. Nobu, W.-T.  
482 Liu, A. Valocchi, Immobilization of selenite via two parallel pathways  
483 during in situ bioremediation, *Environmental Science and Technology*  
484 49 (7) (2015) 4543–4550. doi:10.1021/es506107r.
- 485 [12] T. Willingham, C. Werth, A. Valocchi, Evaluation of the effects of  
486 porous media structure on mixing-controlled reactions using pore-scale  
487 modeling and micromodel experiments, *Environmental Science and  
488 Technology* 42 (9) (2008) 3185–3193. doi:10.1021/es7022835.
- 489 [13] M. Rolle, D. Hochstetler, G. Chiogna, P. K. Kitanidis, P. Grathwohl,  
490 Experimental investigation and pore-scale modeling interpretation of  
491 compound-specific transverse dispersion in porous media, *Transport in*

- 492 Porous Media 93 (3) (2012) 347–362. doi:[https://doi.org/10.1016/](https://doi.org/10.1016/j.advwatres.2020.103574)  
493 [j.advwatres.2020.103574](https://doi.org/10.1016/j.advwatres.2020.103574).
- 494 [14] M. Muniruzzaman, M. Rolle, Impact of multicomponent ionic transport  
495 on ph fronts propagation in saturated porous media, Water Resources  
496 Research 51 (8) (2015) 6739–6755. doi:[10.1002/2015WR017134](https://doi.org/10.1002/2015WR017134).
- 497 [15] Y. Ye, G. Chiogna, O. Cirpka, P. Grathwohl, M. Rolle, Experimental in-  
498 vestigation of transverse mixing in porous media under helical flow con-  
499 ditions, Physical Review E 94 (1). doi:[10.1103/PhysRevE.94.013113](https://doi.org/10.1103/PhysRevE.94.013113).
- 500 [16] J.-L. Auriault, C. Moyne, H. Souto, On the asymmetry of the dispersion  
501 tensor in porous media, Transport in Porous Media 85 (3) (2010) 771–  
502 783. doi:[10.1007/s11242-010-9591-y](https://doi.org/10.1007/s11242-010-9591-y).
- 503 [17] S. Pride, D. Vasco, E. Flekkoy, R. Holtzman, Dispersive transport and  
504 symmetry of the dispersion tensor in porous media, Physical Review E  
505 95 (4). doi:[10.1103/PhysRevE.95.043103](https://doi.org/10.1103/PhysRevE.95.043103).
- 506 [18] T. Le Borgne, D. Bolster, M. Dentz, P. Anna, A. Tartakovsky, Effective  
507 pore-scale dispersion upscaling with a correlated continuous time ran-  
508 dom walk approach, Water Resources Research 47 (12) (2011) W12538.
- 509 [19] P. K. Kang, P. Anna, J. P. Nunes, B. Bijeljic, M. J. Blunt, R. Juanes,  
510 Pore-scale intermittent velocity structure underpinning anomalous  
511 transport through 3-d porous media, Geophysical Research Letters  
512 41 (17) (2014) 6184–6190.
- 513 [20] A. Puyguiraud, P. Gouze, M. Dentz, Stochastic dynamics of lagrangian

- 514 pore-scale velocities in three-dimensional porous media, Water Re-  
515 sources Research doi:10.1029/2018WR023702.
- 516 [21] T. Sherman, E. B. Janetti, G. R. Guédon, G. Porta, D. Bolster, Upscal-  
517 ing transport of a sorbing solute in disordered non periodic porous do-  
518 mains, *Advances in Water Resources* (2020) 103574 doi:https://doi.  
519 org/10.1016/j.advwatres.2020.103574.
- 520 [22] D. Bolster, Y. Méheust, T. Le Borgne, J. Bouquain, P. Davy, Model-  
521 ing preasymptotic transport in flows with significant inertial and trap-  
522 ping effects—the importance of velocity correlations and a spatial markov  
523 model, *Advances Water Resources* 70 (2014) 89–103.
- 524 [23] T. Sherman, A. Foster, D. Bolster, K. Singha, Predicting downstream  
525 concentration histories from upstream data in column experiments, *Wa-  
526 ter Resources Research* (2018) 9684–9694.
- 527 [24] A. Russian, M. Dentz, P. Gouze, Time domain random walks for hydro-  
528 dynamic transport in heterogeneous media, *Water Resources Research*  
529 52 (5) (2016) 3309–3323. doi:10.1002/2015WR018511.
- 530 [25] E. Wright, N. Sund, D. Richter, G. Porta, D. Bolster, Upscaling mixing  
531 in highly heterogeneous porous media via a spatial markov model, *Water*  
532 11 (1) (2019) 53.
- 533 [26] M. Schmuck, P. Berg, Effective macroscopic equations for species trans-  
534 port and reactions in porous catalyst layers, *Journal of the Electrochem-  
535 ical Society* 161 (8) (2014) E3323–E3327. doi:10.1149/2.037408jes.



- [27] H. Kim, J. Bae, D. Choi, An analysis for a molten carbonate fuel cell of complex geometry using three-dimensional transport equations with electrochemical reactions, *International Journal of Hydrogen Energy* 38 (11) (2013) 4782–4791. doi:10.1016/j.ijhydene.2013.01.061.
- [28] T. Gebäck, A. Heintz, A pore scale model for osmotic flow: Homogenization and lattice boltzmann simulations, *Transport in Porous Media* 126 (1) (2019) 161–176. doi:10.1007/s11242-017-0975-0.
- [29] F. Municchi, M. Icardi, Macroscopic models for filtration and heterogeneous reactions in porous media, *Advances in Water Resources* 141 (2020) 103605. doi:https://doi.org/10.1016/j.advwatres.2020.103605.
- [30] N. L. Sund, G. M. Porta, D. Bolster, Upscaling of dilution and mixing using a trajectory based spatial markov random walk model in a periodic flow domain, *Advances in Water Resources* 103 (2017) 76–85.
- [31] T. Sherman, A. Paster, G. Porta, D. Bolster, A spatial markov model for upscaling transport of adsorbing-desorbing solutes, *Journal of contaminant hydrology* 222 (2019) 31–40.
- [32] P. K. Smolarkiewicz, C. L. Winter, Pores resolving simulation of Darcy flows, *Journal of Computational Physics* 229 (9) (2010) 3121–3133.
- [33] J. D. Hyman, C. L. Winter, Stochastic generation of explicit pore structures by thresholding Gaussian random fields, *Journal of Computational Physics* 277 (2014) 16–31.

- 558 [34] OpenCFD Limited, OpenFOAM The Open Source CFD Toolbox, user  
559 guide version v1712, <https://www.openfoam.com/>, accessed: 2019-09-  
560 29 (2017).
- 561 [35] E. Hellinger, Neue begründung der theorie quadratischer formen von un-  
562 endlichvielen veränderlichen., Journal Für Die Reine Und Angewandte  
563 Mathematik (1909) 210–271doi:10.1515/crll.1909.136.210.
- 564 [36] T. Sherman, J. Hyman, M. Dentz, D. Bolster, Characterizing the influ-  
565 ence of fracture density on network scale transport, Journal of Geophys-  
566 ical Research: Solid Earth 125 (1) (2020) e2019JB018547.