

The stability and collapse of lava domes:



insight from UAS-derived 4D structure and slope stability models

Brett B. CARR^{1,4}, Einat LEV¹, Lojç VANDERKLUYSEN², Danielle MOYER², Gayatri Indah MARLIYANI³, Amanda B. CLARKE⁴
bcarr@ldeo.columbia.edu

Motivation

- Dome collapse-generated pyroclastic flows are a primary hazard of lava dome eruptions
- Dome-forming eruptions can last for years to decades, creating a persistent hazard
- Improved understanding of collapse mechanisms and how to estimate the risk of collapse can improve hazard assessment for these eruptions



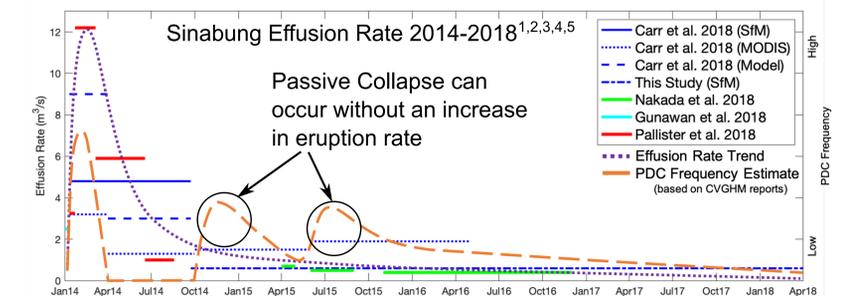
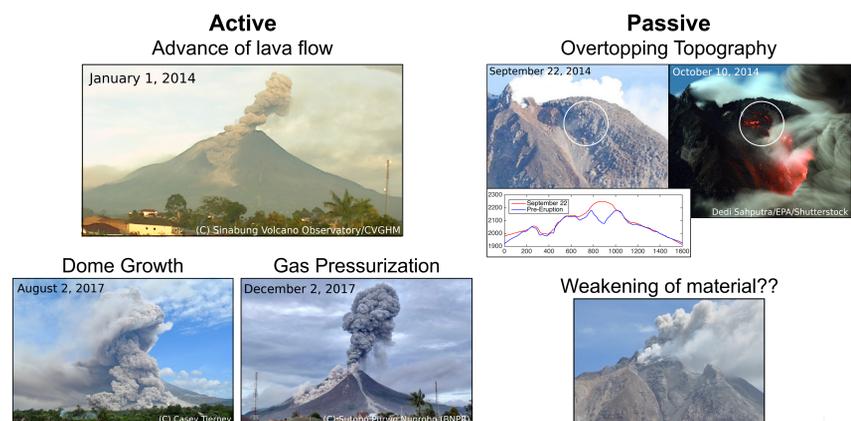
Dome Collapse at Sinabung

The eruption of Sinabung Volcano (2013 - present)^{1,2} has included explosions, emplacement of a 3 km long lava flow^{3,4}, and frequent dome collapse⁵



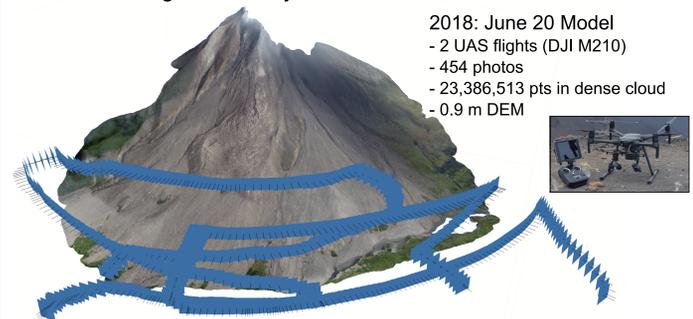
Dome Collapse during the eruption is caused by multiple processes:

- Active Collapse⁶:**
- Caused by effusion of lava and growth of domes and/or flows ("pushed")
 - Size and/or frequency generally correlates with eruption rate, can be anticipated by monitoring eruption signals
- Passive Collapse⁶:**
- Caused by weakening of the internal structure of erupted lava ("pulled" by gravity)
 - Not correlated to other activity, can occur unexpectedly

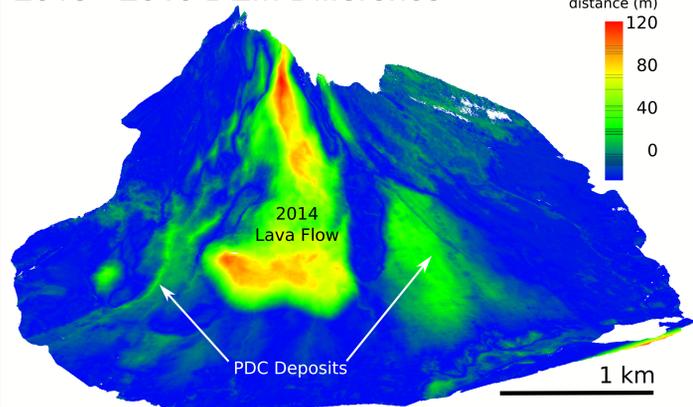


Topographic Change

We create digital elevation models (DEMs) of Sinabung by applying Structure-from-Motion photogrammetry⁷ to image sets collected during field surveys in 2014³ and 2018

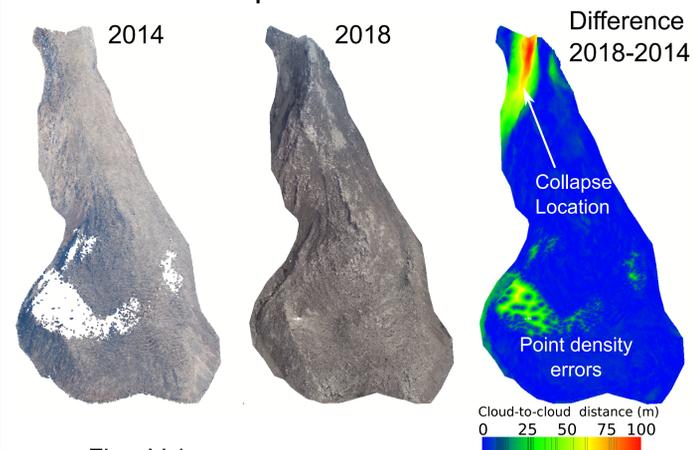


2018 - 2010 DEM Difference



Total Erupted Volume (2013-2018): $173 \times 10^6 \text{ m}^3$
Volume of PDC (Collapse) Deposits: $76 \times 10^6 \text{ m}^3$

Lava Flow Collapse 2014 - 2018



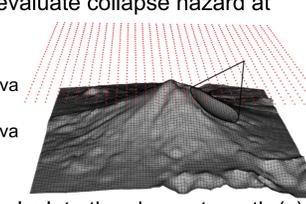
Flow Volume:
Sept 2014³ (Before Collapse): $103 \times 10^6 \text{ m}^3$
June 2018 (After Collapses): $97 \times 10^6 \text{ m}^3$
Initial Collapse Volume: $0.2 \times 10^6 \text{ m}^3$
Total Collapse Volume: $9.4 \times 10^6 \text{ m}^3$

Slope Stability

Scoops3D

We apply the Scoops3D slope stability model⁸ to evaluate collapse hazard at Sinabung. For given input parameters:

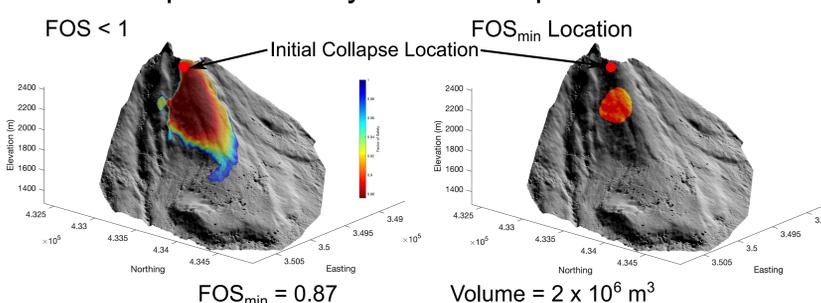
DEM:	SfM Photogrammetry
Cohesion (c) (kPa):	100 - 500 [8,9,10,11] erupted lava 1000 [8,9] edifice
Angle of Internal Friction (φ):	25 - 40 [8,9,10,11] erupted lava 40 [8,9] edifice
Unit Weight (kN m-3):	24.5 [5]



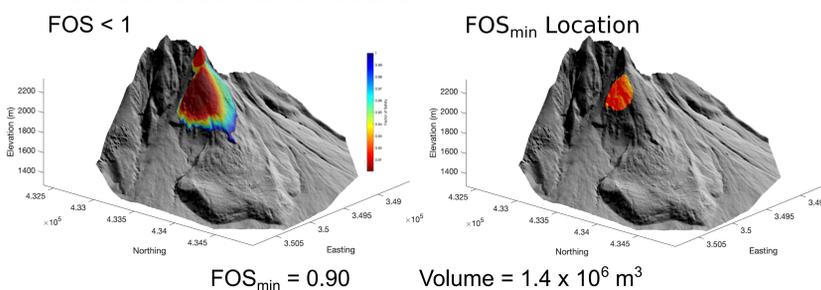
Scoops3D applies the Coulomb failure criteria to calculate the shear strength (s)
 $s = c + \sigma_n \tan(\phi)$ and then the Factor of Safety (FOS) $FOS = \frac{s}{\tau}$ shear strength / shear stress

for thousands of potential rotational, spherical slip surfaces. A FOS < 1 indicates instability. FOS_{min} is the lowest FOS found for the DEM

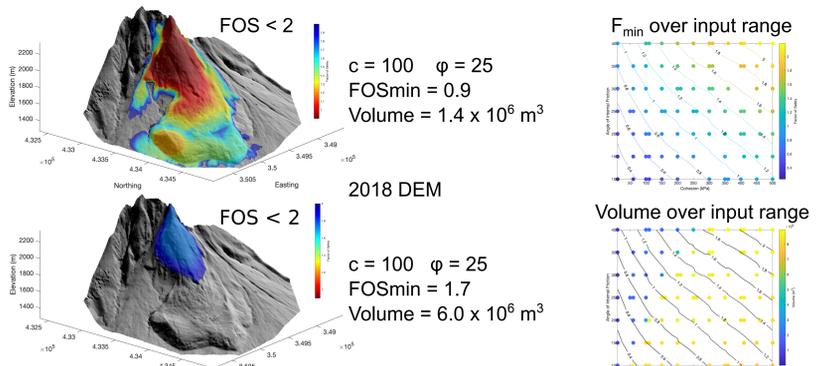
Can Scoops3D Identify 2014 Collapse?



What is the current hazard?



What is the effect of Material Parameters?



Accurate constraint of material properties is essential for assessing the FOS of a failure surface. However, the location and volume of the surface where FOS_{min} occurs is not strongly affected by the material properties.

Conclusions

- For the 2014 DEM:
 - The initial 2014 collapse location has FOS < 1
 - The FOS_{min} is located in material that collapsed
 - The FOS_{min} volume is similar to observed collapse volumes
- For the 2018 DEM:
 - A large region of potential instability still exists
 - The potential collapse size is similar to that from earlier periods of active lava effusion
 - The FOS_{min} is located in the same region as in 2014
- Accurate constraint of material properties is needed to determine if FOS < 1 for potential failure surfaces
- The location and volume of the FOS_{min} can still be reasonably assessed without well-constrained material properties
- Application of Scoops3D with SfM-generated DEMs presents a means to assess passive collapse hazards in near-real-time during an eruption

Acknowledgements

Support for author BC came from NSF EAR Postdoctoral Fellowship Award #1725768. Support for author EL came from NSF EAR Award #1654588.

Research and field work in Indonesia is conducted in cooperation with the Geological Engineering Department at Universitas Gadjah Mada in Yogyakarta, Java, Indonesia through a Memorandum of Understanding with the School of Earth and Space Exploration at Arizona State University, Tempe, AZ.

The Center for Volcanology and Geological Hazard Management (CVGHM), Sinabung Volcano Observatory, and Badan Informasi Geospasial generously shared data that contributed to this study.

References

- Gunawan et al., 2019, J. Volcanol. Geotherm. Res. (382), p. 103-119
- Pallister et al., 2019, J. Volcanol. Geotherm. Res. (382), p. 149-163
- Carr et al., 2019a, J. Volcanol. Geotherm. Res. (382), p. 164-172
- Carr et al., 2019b, J. Volcanol. Geotherm. Res. (382), p. 137-148
- Nakada et al., 2017, J. Volcanol. Geotherm. Res. (382), p. 120-136
- Calder et al., 2002, Geol. Soc. London Memoirs (21), p. 173-190
- James & Robson, 2012, J. Geophys. Res. (117), F03017
- Reid et al., 2000, J. Geophys. Res. (105), p. 6043-6056
- Ball et al., 2018, J. Geophys. Res.: Solid Earth (123), p. 2787-2805
- Schaefer et al., 2019, Earth-Sci. Rev. (192), p. 236-257
- Voight & Ellsworth, 2000, Geophys. Res. Lett. (27), p. 1-4