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The ongoing need for rates: can physiology and omics come together to co-design the measurements needed to understand complex ocean biogeochemistry?

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Abstract

29

30 The necessity to understand the influence of global ocean change on biota has exposed wide-
31 ranging gaps in our knowledge of the fundamental principles that underpin marine life.
32 Concurrently, physiological research has stagnated, in part driven by the advent and rapid evolution
33 of molecular biological techniques, such that they now influence all lines of enquiry in biological
34 and microbial oceanography. This dominance has led to an implicit assumption that physiology is
35 outmoded, and advocacy that ecological and biogeochemical models can be directly informed by
36 omics. However, the main modelling currencies continue to be biological rates and biogeochemical
37 fluxes. Here we ask: how do we translate the wealth of information on physiological potential from
38 omics-based studies to quantifiable physiological rates and, ultimately, to biogeochemical fluxes?
39 Based on the trajectory of the state-of-the-art in biomedical sciences, along with case-studies from
40 ocean sciences, we conclude that it is unlikely that omics can provide such rates in the coming
41 decade. Thus, while physiological rates will continue to be central to providing projections of global
42 change biology, we must revisit the metrics we rely upon. We advocate for the co-design of a new
43 generation of rate measurements that better link the benefits of omics and physiology.

44

Introduction

46

47 A major challenge for ocean scientists is to address key questions on future ecosystem services. For
48 example, how will global climate change alter low latitude primary productivity and hence food
49 security? A powerful tool to address these global-scale questions is Earth system models, such as
50 those within the Coupled Model Intercomparison Project (CMIP6) (Kwiatkowski et al., 2020). The
51 CMIP currencies are mainly the rates at which metabolism occurs in living organisms (i.e.,
52 physiological rates) and the biogeochemical fluxes of bioactive elements. It is unlikely that these
53 currencies will change in the coming decade, for example when CMIP7 is developed. At present,
54 the accuracy of the model projections is hindered by two issues: 1) computational limitations to
55 developing more complex parameterisations for processes such as nitrogen (N) fixation
56 (Kwiatkowski et al., 2020) and 2) our inability to untangle how marine life responds to complex
57 ocean change. For the latter, we need to decipher the fundamental physiological rules that govern
58 biological responses to ocean change. These include the metabolic co-dependencies in response to
59 multiple stressors, and strategies to buffer responses to rapid change, such as phenotypic plasticity
60 and microevolution.

The physiological metrics used to quantify biological rates that are the cornerstones of Earth system models, such as primary productivity, have not fundamentally changed in decades. In contrast, omics techniques have evolved rapidly this century and have superseded physiological metrics as the main approach to study the fundamental principles driving marine life. With this dominance has come an implicit assumption by many that measuring physiological rates directly is obsolete, as they can be inferred from omics (Hellweger, 2020; McCain et al., 2021). However, omics provides a surfeit of data, at a level of detail that is often difficult to relate to the information provided by physiological rate measurements and the current needs of Earth system models. This growing mismatch between the currencies of global-scale models (rates and fluxes) and the aspirations of omics (coupling cellular potential via omics to Earth system model projections) must be addressed urgently.

Here we ask: How do we translate the wealth of information on physiological *potential* from omics-based studies to quantifiable physiological rates and, ultimately, to biogeochemical processes and their representation in Earth system models? We employ three approaches to address this question. First, we examine the evolution of research into ocean N₂ fixation from the perspective of advances in physiology and omics (Fig. 1). Second, we examine the recent trajectory of biomedical research to forecast how ocean sciences might evolve in the next decade. Third, we broaden our view by examining insights that can be gained for understanding the ocean phosphorus (P) and iron (Fe) cycles by better linking omics and physiology. We conclude with advocacy for the co-design of better physiological tools.

Lessons from marine diazotrophy

Here, we use the history of N₂ fixation (diazotrophy) research to reveal the benefits and limitations of physiological rate measurements, and how these measurements are complemented by more recent omics approaches (Fig. 1).

The contribution of diazotrophy to the supply of new N is central to understanding ocean N cycling (Fogg, 1942; Dugdale et al., 1961). Physiological studies played an important early role by quantifying rates of diazotrophy (e.g., Dilworth, 1966). These measurements provided the integrated rates necessary to estimate global biogeochemical fluxes of N (Karl et al., 2002), and to identify the environmental drivers of N₂ fixation (see Carpenter and Capone, 2008), including how climate changes may affect future diazotrophy (Garcia et al., 2011; Hutchins et al., 2013). However, imbalances in these N fluxes have uncovered unidentified N sources, and the subsequent application

of genetic tools has identified additional diazotrophic taxa that contribute to ocean N₂ fixation (Zehr and Capone, 2020).

Nitrogen fixation provides clear examples of both the limitations and benefits of non-targeted omics-based discoveries (Fig. 1). Nitrogenase (*nif*) genes can be used to detect N₂ fixation potential, and their expression is used as an index of N₂ fixation activity (Zehr et al., 1996; Zehr and Montoya, 2007). Omics has revealed diverse N₂ fixers including the unicellular cyanobacteria *Crocospaera* and UCYN-A, and endosymbiotic and heterotrophic diazotrophs (Mehta et al., 2003; Church et al., 2005; Martinez-Perez et al., 2016). However, *nif* gene abundance does not directly equate to N₂ fixation rates (Turk-Kubo et al., 2013). Transcriptomics and proteomics targeting *nif* genes provide more relevant information about nitrogenase *activity* than genomics. However, taxon-specific dynamics can complicate estimates of community N₂ fixation rates (Church et al., 2005), and measurements of *nif* expression are not well correlated with ¹⁵N-based rates of N₂ fixation (Turk et al., 2011).

Thus, despite the insights gained from omics, critical gaps remain in our understanding of the phylogenies, distribution, and physiology of marine N₂ fixers, and accurate global estimates of N₂ fixation remain elusive (Zehr and Capone, 2020). Measuring N₂ fixation remains critical to estimate the biogeochemical processing and ecological fates of new N. However, N₂ fixation is not included in the CMIP6 models, which presently project declining productivity in low latitude oceans in coming decades (Kwiatkowski et al., 2020). So, both rates and omics will be needed increasingly to reveal and quantify currently unknown (but biogeochemically important) pathways for the turnover of N (Fig. 1) to improve global models.

Resolving these unknowns will require combined measurements of *nif* gene expression with rate measurements based on nitrogenase enzyme activity (e.g., Turk et al., 2011). Broader application of flow-through high-throughput rate measurements can improve the spatial and temporal coverage of N₂ fixation (Cassar et al., 2018). Rates, when coupled with omics approaches to N₂ fixation research (Tang et al., 2020), will continue to expand our understanding of diazotroph diversity and could help focus N₂ fixation rate measurements on these emerging diazotrophic groups (Zehr and Capone, 2020). Mechanistic controls on diazotrophy can be revealed through variations in *nif* gene expression (Church et al., 2005), supporting prior conclusions that local environmental conditions influence N₂ fixation rates (Carpenter and Capone, 2008; Capone, 1993). Such environmental controls could be further explored using targeted proteomics analyses (e.g., Saito et al., 2011).

129 The status of omics

130

131 Both marine and biomedical sciences focus on the genome, transcriptome, proteome, and
132 metabolome, with most research on the first three. In the field of meta-omics, marine metagenomics
133 has set the pace, and is directly influencing research into the human microbiome (Pocevičiute and
134 Ismagilo, 2019). Here, we focus on genomics through to proteomics at the cellular level where, in
135 contrast to meta-omics, biomedical research has led the way (Okada and Kuroda 2019). Genomics
136 demonstrates the breadth of possible gene functions, but only catalogues the functional potential of
137 an organism (Sunagawa et al., 2015). Transcriptomics is a popular approach to explore how
138 organisms respond to environmental change by characterizing shifts in mRNA abundance (Evans,
139 2015). Feder and Walser (2005) offered a pointed description of the major issues facing the use of
140 transcriptomics in finding the genes that matter for environmental adaptation. Their critique focused
141 on three major issues: (1) genes with large impacts on fitness are rare and therefore unlikely to be
142 identified with transcriptomics, (2) the relationship between gene expression and fitness is
143 unreliable, and (3) fitness is primarily determined by proteins, and mRNA abundance is a poor
144 proxy for protein abundance. Proteomics, on the other hand, provides taxonomically specific
145 information on structural and metabolic enzymes with tighter correlation to functional activity.
146 Proteomics has advanced methodologically, with more accurate standardized quantitative analyses
147 (Collins et al., 2017; Pino et al., 2020) and protein identifications that allow metabolic profiling
148 (Nunn et al., 2013; Mikan et al., 2020).

149 Numerous efforts have been made to identify correlations between omics layers. However, evidence
150 from both marine and biomedical science reveals that making these linkages is not straightforward.
151 For example, in marine sciences it is well recognized that the amplitude and timing of the mRNA
152 pool does not align with protein expression. This misalignment was illustrated in Waldbauer et al.
153 (2012) while tracking diel changes in the transcriptome and proteome within a single cyanobacteria
154 species (Fig. 2). Subsequent research on the model diatom *Phaeodactylum tricornutum* used
155 multiple omics layers to explore the regulation of N limitation and again reported mismatches
156 between transcript, protein, and metabolite abundance (Remmers et al., 2018). In the further
157 advanced biomedical field, it remains difficult to obtain mechanistic and functional insights by
158 simply integrating multiomics data (Okada and Kuroda, 2019). As far back as the late eighties,
159 Kurland and Ehrenberg (1987) discussed the challenges of linking cellular design and molecular
160 design (such as via enzyme expression) in the context of physiology. More recently, Lalanne et al.
161 (2018) uncovered post-transcriptional controls that ensure the maintenance of the protein
162 stoichiometries required for specific biological pathways. This compensatory mechanism rectifies

163 divergences in regulation driven by changes of internal promoters and terminators. Hence, even in
164 advanced biomedical research there are confounding issues, driven by post-transcriptional and post-
165 translational modifications to enzymes, in deriving metabolic rates from omics.

166 In the marine context, omics has clearly demonstrated large scale patterns in microbial diversity
167 across oceanic provinces and provided insights into which metabolic pathways are active (Fig. 1).
168 However, omics-based approaches provide static ‘snap-shots’ of physiological potential, and we
169 need to improve our quantitative, process-level understanding of the roles of marine microbes in
170 biogeochemical cycles. Indeed, it is physiological activity or realized potential – the chemical
171 fluxes generated by cellular metabolism as modified by biological species differences, external
172 environmental drivers, and the interactions between the two – that drives biogeochemical cycles.

173

174 Linking physiology and omics: the need for co-design

175

176 We propose that physiological rates can bridge biogeochemistry and omics. Physiological rates
177 quantify the integrated activity of proteins that drive marine biogeochemical cycles in units that
178 modellers can use (Fig. 3). Research into the ocean’s N cycle reveals the potential of using the joint
179 expertise of the physiology and omics communities (i.e., co-design) to guide future research (Fig.
180 1). We can extend this complementary approach to use omics datasets to develop new targeted
181 physiological metrics that improve the parameterisation of biogeochemical processes. Here, we
182 explore the feasibility of co-design using case studies of the ocean P and Fe cycles that illustrate
183 how physiological metrics may act as a ‘currency converter’ to link omics datasets and
184 biogeochemical models.

185 In the case of P, a lab study used proteomics and physiological metrics to explore the cumulative
186 effect of five climate-change stressors on a subpolar diatom (Boyd et al., 2015). A central finding
187 was that the effect of decreased nutrient supply in a future ocean was offset by warming.

188 Proteomics revealed that a decreased need for P was driven by the under-expression of P-containing
189 proteins associated with translation (Fig. 3). Physiological metrics corroborated this finding, with
190 lower cellular P quotas under warming. Hence, P quotas acted as a currency converter between
191 protein synthesis and the biogeochemical cycle of P. They showed serendipitously a link between
192 protein synthesis and P quotas. In the future, we must actively seek conceptual linkages, rather than
193 uncovering them by chance. Better links from omics via physiology to biogeochemistry would
194 benefit from input from the modelling and biogeochemical research communities.

195 Physiology was established earlier than omics or biogeochemistry and so many of the conventional
196 metrics used preceded developments in these disciplines. This begs the question: are we currently
197 measuring the best physiological metrics to mesh omics with biogeochemistry? Two examples that
198 begin to straddle the gaps between omics and physiology come from Saito et al. (2011) and Wu et
199 al. (2019). The former revealed diel changes in the proteome, including Fe-metalloproteins involved
200 in N₂ fixation and photosynthesis of *Crocospheara watsonii* resulting in more efficient use of Fe,
201 which is essential for N₂ fixation. In the latter case, protein expression and physiological metrics
202 were coupled to examine the influence of Fe and manganese on *Phaeocystis antarctica*.

203 Although our current choice of physiological metrics needs urgent attention, there is compelling
204 evidence of the utility of long-established assays, such as those used to determine the
205 macromolecular P content of cells from Liefer et al. (2019), for more innovative phytoplankton
206 cellular P models (Inomura et al., 2020). But can we be inventive, and use omics to interpret P
207 physiology in a more holistic manner (Fig. 4)? Physiology can provide valuable insights, even when
208 considering only a few components of the cellular P cycle. Imagine the progress if we developed
209 better metrics jointly with omics (Feng et al., 2014; Lin et al. 2016). So, the way ahead may be to
210 use molecular biology to ‘reverse engineer’ the most pertinent physiological metrics (Fig. 4). For
211 example, a useful point of departure would be to select processes in which protein abundance
212 correlates with quantifiable metabolic activity. Such co-design, in our opinion, will further facilitate
213 the transition from lab- to field-based omics and will lead in the coming decades to incorporation of
214 omics into biogeochemical models.

215 The transition to field studies will face additional challenges that centre on how marine biota
216 integrate environmental history (i.e., cellular status imposed by conditions encountered prior to
217 sampling; Fig. 2) (Prairie et al., 2012; Deutschmann et al., 2021). This requires a multi-stranded
218 approach. First, placing the sampling locale in a wider environmental context (Figure 5A). For
219 example, profiling robotic floats with multiple sensors are providing synoptic snapshots of spatial
220 variability in ocean properties along with the prior seasonal dynamics of key resources such as
221 nutrients (Claustre et al., 2021). Second, how do such prior oceanic conditions set cellular status, for
222 example the degree of Fe stress (Fig. 5B)? An open question is whether the relationship between
223 environmental forcing and cellular status is instantaneous or lagged (Fig. 2). Will such co-designed
224 metrics reconcile a biological product with a chemical residual since different physiological metrics
225 display a range of response times (Boyd et al., 2005; Baker et al., 2018), as do different omics
226 layers (Waldbauer et al., 2012)? One promising approach to probe environmental history and
227 cellular status is physiological titration. For example, by manipulating Fe availability to
228 contextualize cellular Fe status (Fig. 5B).

229

230 Towards the future

231

232 We conclude with recent field-leading examples from ocean sciences, that seek to derive metabolic
233 rates from omics, explored through the lens of biomedical sciences. Saito et al. (2020) conducted
234 metaproteomic analysis on subsurface biota in the Tropical North Pacific to pinpoint commonly
235 occurring enzymes. They reported that nitrite oxidoreductase associated with the bacterium
236 *Nitrospina* was abundant in this stratum and explored whether they could estimate rates of nitrite
237 oxidation using wide-ranging methods, including biochemistry (specific activity), physiology
238 (Michaelis-Menten kinetics), and omics. Despite employing this innovative suite of approaches
239 derived rates ranged >200-fold, pointing to the need to develop targeted physiological assays (c.f.
240 Fig. 4). There are also promising initial developments from the emergence of phenomenological
241 models based on simple geochemical/taxonomic principles that yield phytoplankton growth rates
242 assuming steady-state growth (McCain et al., 2021).

243 The latest developments in biomedical and model-system omics suggest obtaining rates from omics
244 is still under development. First, holistic investigations of well-characterized model organisms have
245 tracked every metabolite and protein to generate enzyme-directed functional rates in the bacterium
246 *Escherichia coli* (Taniguchi et al., 2010) and the yeast *Saccharomyces cerevisiae* (Ho et al., 2018),
247 but this approach is restricted to the organisms for which the function of every gene and protein is
248 known. Second, expression-fitness landscapes (linking enzyme expression with growth rate) have
249 revealed that enzyme expression can have a ‘ripple’ effect across layers of biological organisation
250 ranging from mechanistic, regulatory to systemic (Lalanne et al., 2021), which adds further
251 complexity to deriving growth rates from enzymatic fluxes. Third, sophisticated microbiome studies
252 (from cheese to the human gut) (Pocevičiute and Ismagilo, 2019), which are more akin to oceanic
253 microbial systems, reveal that there are still a high number of metabolic functions that remain
254 uncharacterized (Price et al., 2018). Fourth, progress in tackling cell regulatory mechanisms using
255 multiomic modelling has been made but requires complex computing using deep neural networks
256 such as GEMS (Genome-scale metabolic models) (Okada and Kuroda 2019).

257 These four categories of advanced well-resourced research point to challenges yet to be surmounted
258 in obtaining physiological rates from omics for biomedical sciences. But, they also provide
259 cautionary lessons for ocean sciences. In our opinion, it may be more straight-forward to co-design
260 targeted physiological metrics that better link omics with marine biogeochemistry.

261

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440

441

442 Figure legends

443 Figure 1. The contributions of physiology and omics to understanding the role of diazotrophy in the
444 ocean N cycle (based on Zehr and Capone, 2020). Key events in the physiology timeline (top green
445 line) include estimation of N fluxes through nitrogenase (Dugdale and Dugdale, 1962), initial
446 estimates of global marine N₂ fixation rates (Capone et al., 1982), and the combining of lab and
447 field measurements to understand individual diazotrophs and community contributions and
448 constraints. Pivotal events in the omics timeline (lower green line) include problem solving (Zehr
449 and Montoya, 2007) and discovery of diazotroph diversity including in unicellular cyanobacteria
450 group A (UNCYN-A) and diverse uncultured heterotrophic bacteria (UHB) (Martinez-Perez et al.,
451 2016). Recent examples of more integrated physiological and omics co-designed studies (Walworth
452 et al., 2016; Held et al. 2020) offer an important way forward.

453

454 Figure 2. Examples of the potential for mismatches in transcriptomics versus proteomics in a pico-
455 prokaryote over the diurnal cycle. A) The diel cycling and amplitudes of transcripts and proteins in
456 *Prochlorococcus* for Ribonucleotide reductase (nrdJ), the large sub-unit of Rubisco (rbcL), and
457 Geranylgeranyl diphosphate reductase (chlP). B) Histogram of lag-times for proteins and their
458 transcripts for a 312 gene dataset. Antiphase refers to genes that are offset by ~12 h (i.e., 50%) of
459 the diel cell cycle. Redrawn from Waldbauer et al. (2012).

460

461 Figure 3. An example illustrating the utility of physiological metrics as a ‘currency converter’ to
462 link omics and biogeochemical modelling. A) The under- (downward arrows) and over-expression
463 (upwards arrows) of proteins in 4 treatments within a climate change manipulation experiment
464 measured with proteomics (Boyd et al., 2015). Warming results in an under-expression of P-
465 containing proteins associated with translation. B) Corresponding changes to the cellular P quotas
466 of the study subject, a lab culture of a subantarctic diatom, across the treatments A-D. This
467 physiological metric reveals the causal link between under-expression of translation proteins and
468 decreased P quotas (as previously described by Toseland et al., 2013). C) A subset of global model
469 projections of upper ocean phosphate (PO₄⁻) stocks across biogeochemical models of different
470 complexity (Kriest et al., 2010). The approaches employed in panels A and C can be linked using
471 the cellular P quotas obtained from panel B.

472

473 Figure 4. The potential of reverse-engineering physiological metrics to provide better linkages with
474 molecular tools using the example of P. A) Findings of a physiological study (Leifer et al., 2019)
475 using a cluster of long-established metrics (residual P pools/intracellular storage of inorganic P) to
476 compare the P allocation strategies of a diatom (*Thalassiosira pseudonana*) and a prasinophyte
477 (*Micromonas sp.*). B) Cartoon summarizing the known PO_4^- acquisition and metabolic pathways
478 that may be present in most phytoplankton species (for details see Fig. 4 in Lin et al., 2016). C) A
479 KEGG map from I-PATH (Letunic et al., 2008; Darzi et al., 2018) overlaid with the functional
480 categories of differentially expressed proteins (mapped to KEGG pathways) involved in various
481 biological processes for P limitation by a *Phaeocystis* species (Feng et al., 2014). I-PATH is a web-
482 application for the visualization and analysis of cellular pathways from omics (e.g., see Nunn et al.,
483 2013 for Fe replete versus Fe deplete proteomes).

484 Figure 5. Utility of environmental context to define the present physiological status of cells in
485 relation to prior oceanic conditions. A) Dissolved Fe time series for the upper ocean in the
486 subtropical Atlantic (BATS site) that reveals conspicuous aerosol Fe inputs ($>0.5 \text{ nmol L}^{-1}$) along
487 with the influence of eddy activity ($<0.3 \text{ nmol L}^{-1}$) on dissolved Fe concentrations (Sedwick et al.,
488 2020). B) Photosynthetic efficiency of PSII (F_v/F_m) measured in deckboard incubation experiments
489 ‘titrated’ with dissolved Fe concentrations by either reducing bioavailable Fe using the fungal
490 siderophore desferrioxamine B (DFB) or increasing it with chelated inorganic Fe addition (from
491 Wilhelm et al., 2013). The circles denote putative linkages between chemical stocks and biological
492 responses (red = high Fe; green = low Fe).

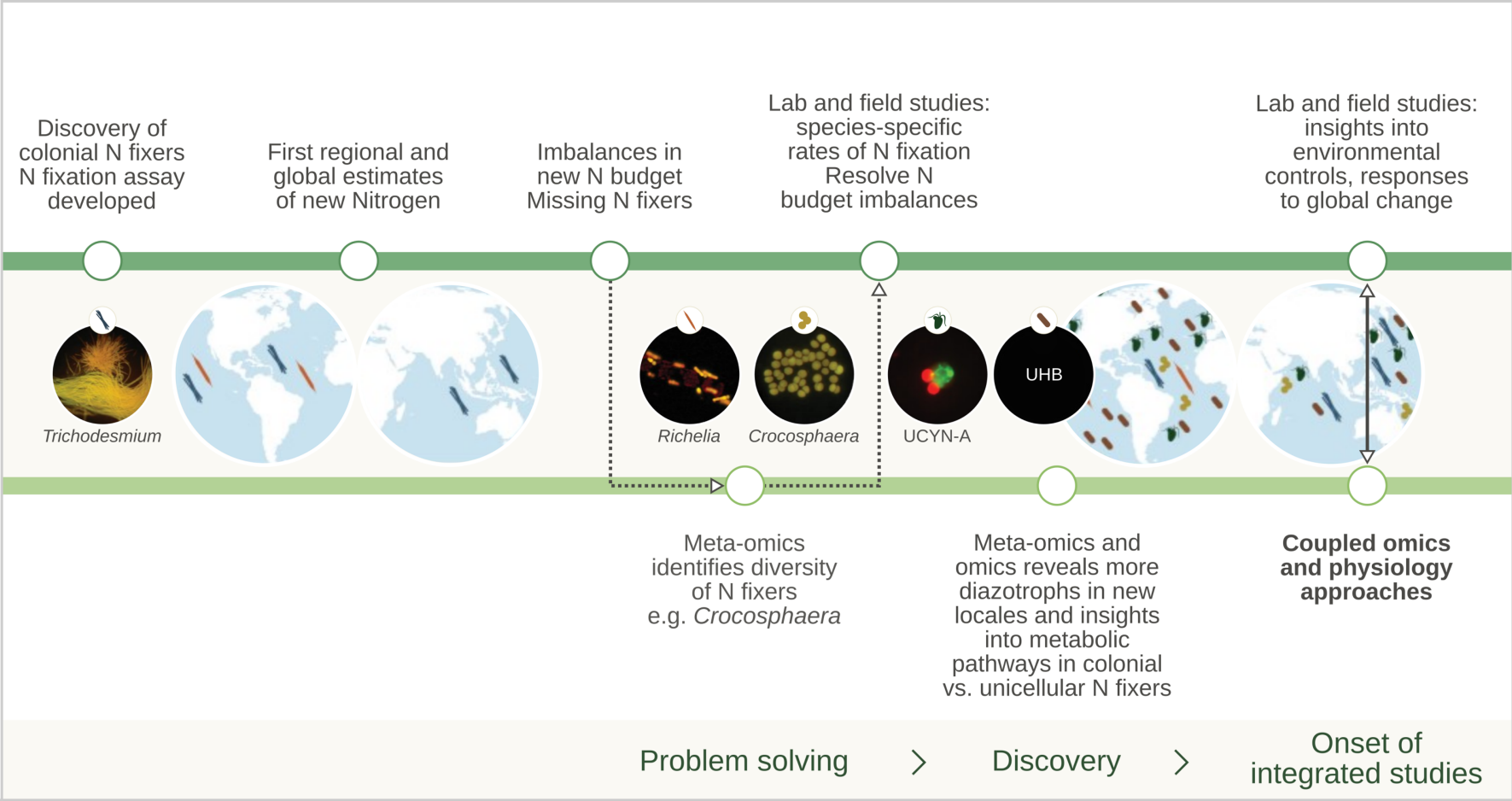


Figure 1

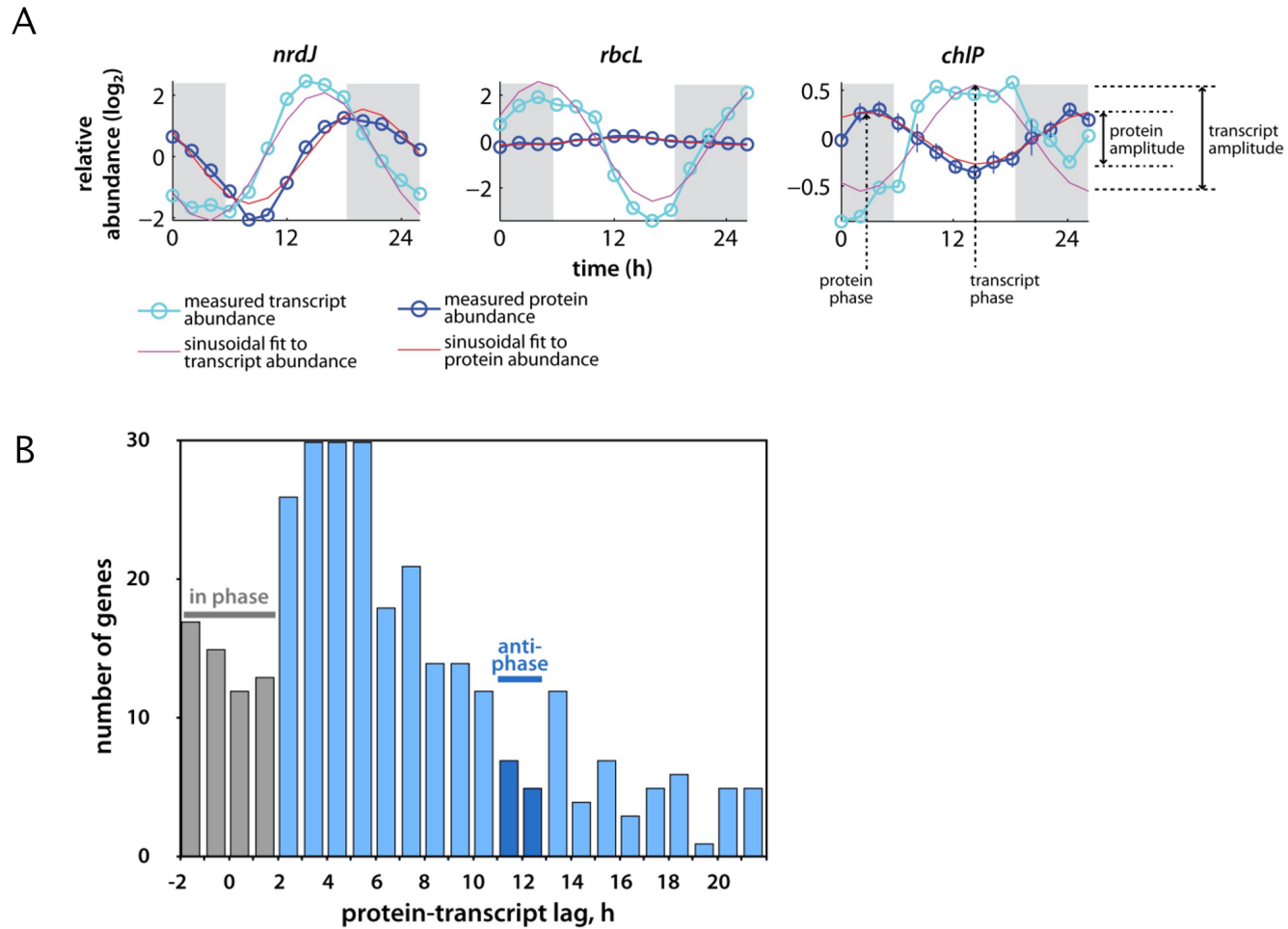


Figure 2



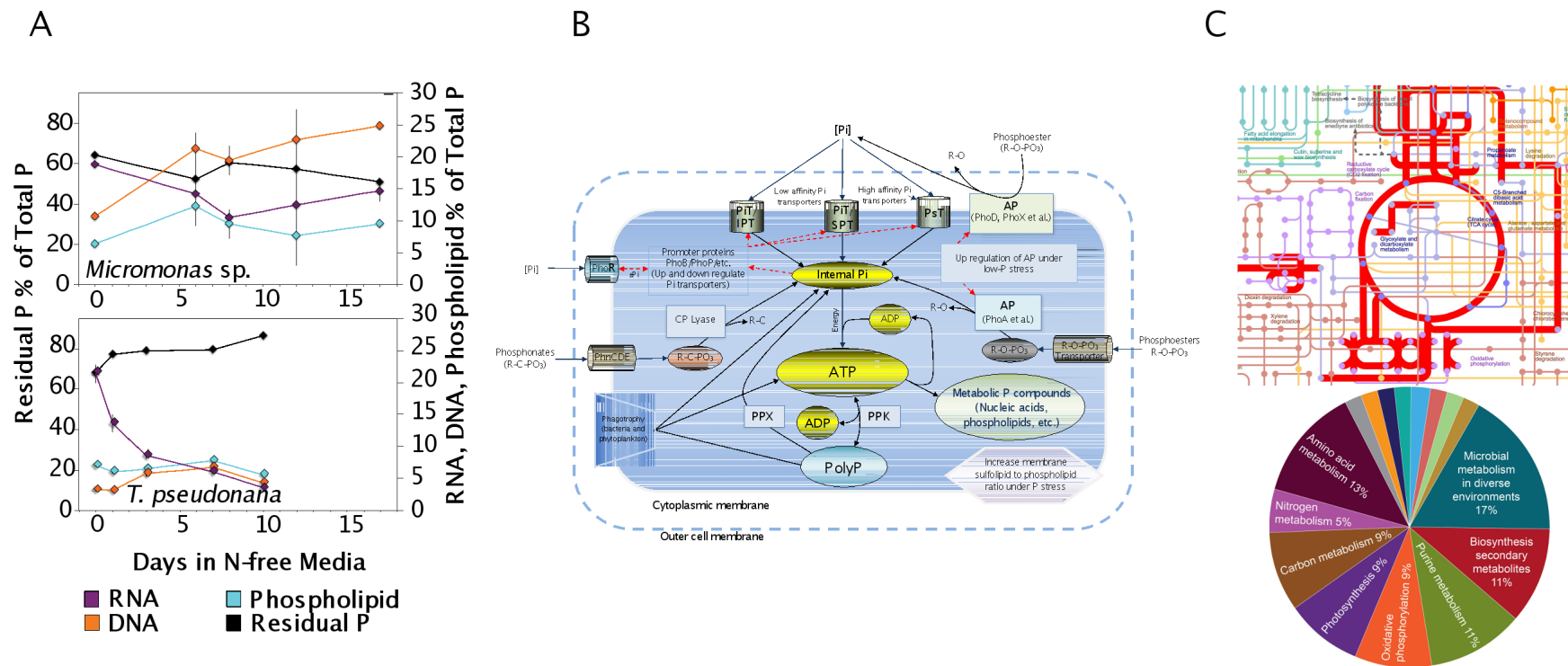


Figure 4

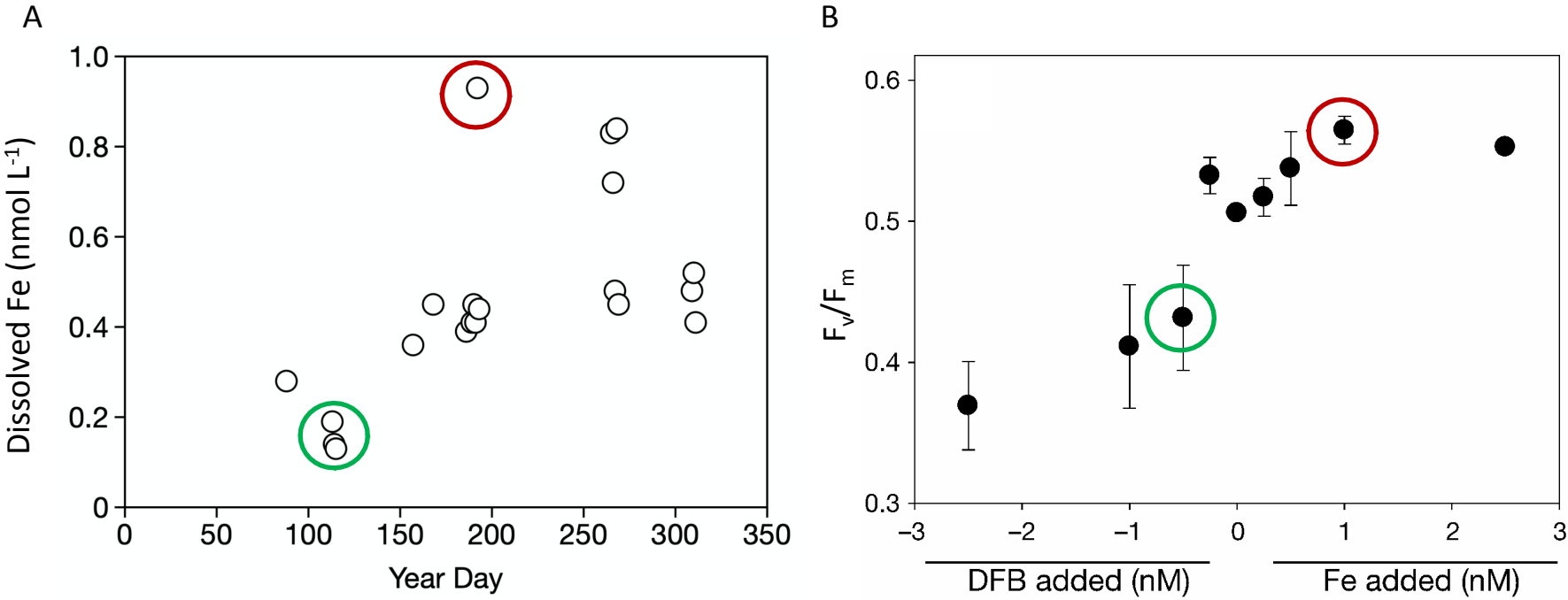


Figure 5