

1 **Dynamical coupling between the low-latitude lower**
2 **thermosphere and ionosphere via the non-migrating**
3 **diurnal tide as revealed by concurrent satellite**
4 **observations and numerical modeling**

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14 **Key Points:**

- 15 • First results from the SORTIE CubeSat show a large low-latitude ionospheric wave-
16 4 (WN4) structure observed concurrently by ICON and SABER
- 17 • Spectral analyses of SORTIE IVM, ICON IVM and model output demonstrate
18 that the DE3 tide is responsible for this strong IT WN4 coupling
- 19 • SORTIE and ICON IVM provide insights into tropical tropospheric influences on
20 the IT by measuring poorly sampled altitudes simultaneously

Abstract

The diurnal, eastward propagating tide with zonal wavenumber 3 (DE3) is an important tidal component due to its ability to effectively couple the ionosphere-thermosphere (IT) with the tropical troposphere. In this work, we present the first results of a prominent zonal wavenumber 4 (WN4) structure in the low-latitude ionosphere observed by the Scintillation Observations and Response of The Ionosphere to Electrodynamics (SORTIE) CubeSat mission during May 27 - June 5, 2020. Least-squares analyses of concurrent in-situ ion number density measurements from the SORTIE and the Ionospheric Connection Explorer (ICON) satellites near 420 and 590 km show this pronounced WN4 to be driven by DE3. Thermosphere Ionosphere Mesosphere Energetics Dynamics Sounding of the Atmosphere using Broad band Emission Radiometry (TIMED/SABER) temperatures and Specified-Dynamics Whole Atmosphere Community Climate Model with thermosphere and ionosphere eXtension (SD/WACCM-X) output demonstrate that the ionospheric WN4 structure is driven by DE3 propagating from the lower thermosphere.

Plain Language Summary

The extent to which terrestrial weather (below ~ 30 km) can influence the ionosphere and thermosphere (IT) is a fascinating discovery of the last two decades or so. The IT is known to vary significantly from day to day, and this day-to-day weather is largely driven by processes originating in the lower atmosphere, especially during periods of quiet solar activity. Accurate forecasting of the IT variability thus depends on the ability to forecast the component that originates in the lower atmosphere. Ionospheric variability translates to uncertainty in navigation and communications systems, while thermospheric variability translates to uncertainty in orbital and reentry predictions. In this work, we present the first results from the SORTIE CubeSat showing a large amplitude structure with four longitudinal peaks in the ionosphere associated with the well-known diurnal eastward propagating tide with zonal wavenumber 3 (DE3) originating in the tropical troposphere. Our analyses presented in this paper use concurrent SORTIE, ICON, and TIMED satellite observations during May 27 - June 5, 2020 and a whole atmosphere model to interpret SORTIE measurements. Our results suggest SORTIE and ICON to be excellent and complementary observational platforms for studying the influence of terrestrial weather on IT variability.

1 Introduction

The lower atmosphere drives variability in the ionosphere-thermosphere (IT) system through the vertical propagation of waves, including tides, planetary waves, and Kelvin waves. These waves are periodic in time and longitude due to the rotation of the Earth, and interact with the lower IT region to modulate electric fields that map to higher altitudes and redistribute plasma in the 200-1000 km region. Due to the geometry of magnetic field lines near the equator, much of this variability occurs at low latitudes and is driven by waves that are excited by deep convective processes in the tropical troposphere and that propagate upwards into the IT system. Tropical troposphere variability is essentially mapped to the IT system through a variety of neutral-plasma coupling processes, and over a range of spatial and temporal scales. One important class of global-scale atmospheric waves characterized by periods that are harmonics of a solar day are thermal tides. These atmospheric tides can be generated in different altitudinal regions due to tropospheric latent heating, absorption of tropospheric infrared radiation by water vapor, absorption of solar ultraviolet radiation by stratospheric ozone, thermosphere molecular oxygen absorption of extreme ultraviolet radiation, and wave-wave interactions (e.g., Chapman and Lindzen, 1970; Hagan and Forbes, 2002; Hagan et al., 2007; Liu, 2016). The main pathways responsible for the modulation of the ionosphere by tides are direct propagation of atmospheric tides into the ionosphere and thermosphere (e.g., Hagan et al., 2007; Oberheide et al., 2009) and indirect coupling via the ionosphere E-region dynamo (e.g., Jin et al., 2008; Ren et al., 2010; Wan et al., 2008, 2010, 2012).

Several studies investigated the importance of IT variability driven by lower-atmosphere wave sources. Initial work (e.g., Hagan et al., 2007; Jin et al., 2008; Fang et al., 2009; Wan et al., 2010) focused on verifying Sagawa et al.'s (2005) and Immel et al.'s (2006) suggestion that the wavenumber-4 (hereafter, WN4) structure seen in satellite-borne Sun-synchronous F-region ionospheric data was due to the modulation of dynamo electric fields by non-migrating tides propagating from below, and in particular due to the diurnal eastward-propagating tide with zonal wavenumber 3 ($s = -3$, i.e., DE3). DE3 originates in the tropical troposphere by latent heat release in deep convective clouds (e.g., Hagan, 1996; Hagan and Forbes, 2002; Lieberman et al., 2007), and its first equatorially-symmetric Hough mode is the largest component in the lower thermosphere (Oberheide et al., 2009; Truskowski et al., 2014), capable of propagating well into the middle thermosphere (Oberheide et al., 2011; Gasperini et al., 2015, 2017a, 2018) due to its longer vertical wavelength. When

85 viewed at quasi-fixed local time (LT) from slowly precessing satellites, this feature man-
86 ifests as a 4-peak longitude structure. Subsequent studies examined the LT and seasonal
87 variations of related WN4 and WN3 structures (e.g., Lin et al., 2007; Liu and Watan-
88 abe, 2008; Ren et al., 2009), and investigated the underlying mechanisms in further depth
89 (e.g., Oberheide et al., 2011; He et al., 2011; Mukhtarov and Pancheva, 2011; Maute et
90 al., 2012; Chang et al., 2013; Lei et al., 2014; Cho et al., 2015; Onohara et al., 2018).

91 In this work, we present observational evidence of prominent WN4 coupling between
92 the lower thermosphere near 105 km and the ionospheric F-region at heights near 420
93 km and 590 km during May 27 - June 5, 2020. Our results take advantage of ion num-
94 ber density (hereafter, ion density) measurements from the Ion Velocity Meter (IVM)
95 instruments onboard the Scintillation Observations and Response of The Ionosphere to
96 Electrodynamic (SORTIE) and Ionospheric Connection Explorer (ICON) spacecraft and
97 kinetic temperature observations from the Sounding of the Atmosphere using Broadband
98 Emission Radiometry (SABER) instrument onboard the Thermosphere Ionosphere Meso-
99 sphere Energetics Dynamics (TIMED) satellite. We further connect this WN4 longitude
100 variability to DE3 by spectrally analyzing ICON and SORTIE IVM data and investi-
101 gate its latitude-height amplitude and phase structure using a Whole Atmosphere Com-
102 munity Climate Model with thermosphere and ionosphere eXtension (SD/WACCM-X)
103 simulation with realistic wave forcing imposed by nudging Modern-Era Retrospective anal-
104 ysis for Research and Applications version-2 (MERRA-2) reanalysis data in the tropo-
105 sphere and stratosphere. After a brief description of the data, models, and methods (Sec-
106 tion 2), we present in detail the observational and modeling results regarding the WN4
107 variability and its connection to the DE3 tide (Section 3), and conclude with a brief sum-
108 mary (Section 4).

109 **2 Data, Models, and Methods**

110 **2.1 SORTIE/IVM**

111 SORTIE is a NASA Heliophysics System Observatory (HSO) 6U CubeSat mission
112 to investigate the underlying causes behind the appearance of plasma structures in the
113 F-region ionosphere, leading to equatorial plasma bubbles, and the evolution of these struc-
114 tures after their formation (Crowley et al., 2016). SORTIE was launched onboard Dragon
115 CRS-19 to the International Space Station (ISS), from where it was deployed on Febru-

116 ary 19, 2020 in a nearly circular orbit near 420 km with $\sim 51.6^\circ$ inclination. It carries
117 two science instruments, a miniature IVM (MIVM), and a micro-Planar Langmuir Probe
118 (Crowley et al., 2016). This study employs SORTIE’s IVM Level 2 ion density data prod-
119 uct with a 4-second temporal cadence.

120 **2.2 ICON/IVM**

121 ICON is a NASA HSO mission designed to study the fundamental connections be-
122 tween the dynamics of the neutral atmosphere at altitudes between 100 km and 300 km
123 and the charged particle motions at low and middle latitudes from a nearly circular or-
124 bit at an altitude near 590 km with $\sim 27^\circ$ inclination (Immel et al., 2018). The ICON
125 payload includes an IVM instrument that provides in situ measurements of the ion drift
126 motions, density, temperature and major ion composition (Heelis et al., 2017). The IVM
127 is comprised of two instruments, the Retarding Potential Analyzer (RPA) and the Drift
128 Meter (DM). This study employs ICON’s IVM-A Data Product 2.7 v4 determined from
129 RPA measurements, with a 1-second temporal cadence. Preliminary validation work by
130 the ICON IVM team reports accuracy around $\pm 10^3 \text{ cm}^{-3}$ for this data product (see the
131 acknowledgments for further information).

132 **2.3 TIMED/SABER**

133 The SABER instrument was launched onboard the TIMED satellite (also a NASA
134 HSO mission) on December 7, 2001. SABER provides measurements of kinetic temper-
135 ature from ~ 20 km to ~ 120 km altitude (Mertens et al., 2001). SABER views the at-
136 mospheric limb from an orbit of ~ 625 km altitude and $\sim 73^\circ$ inclination, so that the lat-
137 itude coverage on a given day extends from about 53° in one hemisphere to about 83°
138 in the other. This viewing geometry alternates once every ~ 60 days. Errors in the re-
139 trieved temperatures in the 80-105 km region are estimated to be ± 1.5 -5 K (Garcia-Comas
140 et al., 2008). This study takes advantage of the SABER Level 2a Data Product with a
141 ~ 55 -second temporal cadence.

142 **2.4 WACCM-X**

143 WACCM-X (Liu et al., 2018) is a configuration of the National Center for Atmo-
144 spheric Research (NCAR) Community Earth System Model (CESM) (Hurrell et al., 2013)

145 that extends the atmospheric component into the thermosphere, with a model top bound-
 146 ary between 500 and 700 km. The standard spatial resolution of WACCM-X is 1.9° by
 147 2.5° in latitude and longitude, respectively, and 1/4 scale height vertically above the up-
 148 per stratosphere. To simulate the lower and middle atmosphere, WACCM-X can be run
 149 with specified dynamics (SD) that constrains the troposphere and stratosphere dynam-
 150 ics using MERRA-2 reanalysis data (Gelaro et al., 2017). This study employs output from
 151 an extended SD/WACCM-X version v2.1 simulation that covers the May 27 - June 5,
 152 2020 period. The empirical ion convection patterns are specified using the Heelis et al.
 153 (1982) empirical model. The MERRA-2 forcing provides a realistic representation of the
 154 wave forcing in the lower and middle atmosphere making this simulation an ideal physics-
 155 based framework for interpreting atmosphere-ionosphere coupling by tides.

156 2.5 Analysis of Satellite Data

157 When combined, the ascending and descending nodes of SORTIE and ICON cover
 158 24-hour of LT in ~ 30 days (i.e., each node samples ~ 4 hours of LT in about 10 days).
 159 Figure 1 shows the LT coverage of ICON (blue plus signs) and SORTIE (red squares)
 160 IVM measurements at 25°N (panel *b*), at the equator (panel *c*), and at 25°S (panel *d*)
 161 during May 27 - June 5, 2020. At the equator, SORTIE's (ICON's) orbit precesses from
 162 ~ 2 LT/ ~ 14 LT (~ 7 LT/ ~ 19 LT) on May 27 to ~ 22 LT/ ~ 10 LT (~ 2 LT/ ~ 14 LT) on
 163 June 5 (Figure 1*c*). As shown in Figure 1*a*, this 10-day period is characterized by low
 164 solar activity (average F10.7 ~ 70) with two minor geomagnetic disturbances (ap ~ 25 on
 165 May 30, 2020 and ap ~ 15 on June 1-2, 2020). For this study, ICON and SORTIE IVM
 166 ion density are analyzed by both adding and subtracting ascending and descending node
 167 measurements. Wave analysis is performed by applying least-squares methods to SOR-
 168 TIE IVM and ICON IVM ion density, and to SD/WACCM-X ion density and neutral
 169 temperature.

170 3 Results

171 The global distribution of ion density observed by SORTIE (ICON) IVM near 420
 172 km (590 km) altitude during May 27 - June 5, 2020 is presented in Figure 2. SORTIE
 173 daytime (~ 10 -14 LT) and nighttime (~ 22 -2 LT) 10-day averages are shown in panels *a*
 174 and *b*, respectively. Figures 2*a'* and 2*b'* show the corresponding ICON daytime (~ 14 -18
 175 LT) and nighttime (~ 2 -6 LT) 10-day averages, respectively. Excellent low-latitude cov-

176 erage is achieved by both SORTIE (about $\pm 51.6^\circ$ latitude) and ICON (about $\pm 27.1^\circ$ lat-
 177 itude) during this 10-day period, as evidenced by the distribution of the black dots in
 178 panels *a-a'* and *b-b'* showing the measurement locations. Ion density values up to $\sim 4.4 \times 10^5$
 179 cm^{-3} ($\sim 2.1 \times 10^5 \text{ cm}^{-3}$) are observed during daytime and up to $\sim 1.7 \times 10^5 \text{ cm}^{-3}$ ($\sim 0.5 \times 10^5$
 180 cm^{-3}) during nighttime by SORTIE (ICON) IVM. Note that a scaling factor is applied
 181 multiplying the ICON ion density by a factor of 2 given ICON's higher mean altitude
 182 (i.e., ~ 590 km versus ~ 420 km for SORTIE). Figures *2a''-2b''* display ascending and de-
 183 scending node differences for SORTIE and ICON IVM ion density, respectively, each scaled
 184 (multiplied) by 0.5. Both SORTIE and ICON show enhanced ion density at low latitudes
 185 around 30°S - 30°N with a prominent WN4 structure and peaks near 180°E , 100°E , 45°E ,
 186 and 100°W . This type of feature in the low-latitude ionosphere was reported in previ-
 187 ous satellite-based measurements (e.g., Immel et al., 2006; Lin et al., 2007; Liu and Watan-
 188 abe, 2008, Liu et al., 2009), as discussed in Section 1. This notable WN4 structure ob-
 189 served by SORTIE and ICON points to a possible modulation of the background ion den-
 190 sity by the non-migrating DE3 tide as observed from their slowly precessing orbits. Note
 191 that while SORTIE IVM observes an O^+ -dominated ionosphere, ICON IVM samples non-
 192 negligible H^+ (especially during nighttime). Thus, to best interpret the concurrent iono-
 193 spheric WN4 signatures observed by ICON and SORTIE, Figure *2b-2b''* shows ICON's
 194 O^+ density. The Pearson correlation coefficient between the WN4 structure seen in the
 195 SORTIE total ion density and ICON O^+ density (see Figure *2a''-2b''*) is calculated to
 196 be ~ 0.87 . Also note that the equatorial ionization anomaly (EIA) is not evident in Fig-
 197 ure 2 likely due SORTIE's and ICON's mean altitude (~ 420 km and ~ 590 km, respec-
 198 tively) being significantly higher than the F-layer peak height (~ 200 - 350 km), in agree-
 199 ment with Figure 2 of Mukhtarov and Pancheva (2011).

200 It is important to realize that while the variability due to the semidiurnal eastward-
 201 propagating tide with $s = -2$ (SE2) and stationary planetary wave 4 (SPW4) would largely
 202 be eliminated by the ascending/descending node differences shown in Figures *2a''-2b''*
 203 for latitudes less than about $\pm 10^\circ$ ($\pm 30^\circ$) for ICON IVM (SORTIE IVM), aliasing from
 204 the diurnal westward-propagating tide with $s = 5$ (DW5) and from the terdiurnal eastward-
 205 propagating tide with $s = -1$ (TE1) may also contribute to a WN4 longitude structure
 206 seen at a constant LT (see Lieberman, 1991; Oberheide et al., 2000; Gasperini et al., 2015,
 207 2017b, 2018, 2020). Additionally, the LT difference between ICON's (SORTIE's) ascend-
 208 ing and descending node IVM measurements is < 9 hours for latitudes greater than around

209 $\pm 10^\circ$ ($\pm 30^\circ$) (see Figure 1). Hence one can anticipate significant aliasing due to vari-
 210 ability associated with SE2 and SPW4 at latitudes greater than around $\pm 10^\circ$ for ICON
 211 and $\pm 30^\circ$ for SORTIE due to this LT difference being < 9 hours. Thus, to avoid these
 212 aliasing issues in the analysis of the WN4 structure we (1) combine 30 days of SORTIE
 213 and ICON data collected during May 27 - June 25, 2020 to acquire full 24-hour LT cov-
 214 erage, and (2) use output from a 3-hourly SD/WACCM-X simulation. Figure 3 shows
 215 the period versus zonal wavenumber amplitude spectra of SORTIE IVM ion density and
 216 ICON IVM O^+ ion density obtained by combining ascending and descending node mea-
 217 surements during May 27 - June 25, 2020, and of SD/WACCM-X ion density during May
 218 27 - June 5, 2020 derived using the full model output. Around 10°S - 10°N magnetic lat-
 219 itude (MLAT) SORTIE (panel *a'*), ICON (panel *b'*) and WACCM-X (panel *c'*) ion den-
 220 sity spectra reveal the existence of a pronounced DE3 component with amplitudes up
 221 to $\sim 2 \times 10^4 \text{ cm}^{-3}$. DE3 is found to be nearly absent in both WACCM-X and the SOR-
 222 TIE (ICON) observations at 40°N MLAT (25°N MLAT) and 40°S MLAT (25°N MLAT),
 223 in agreement with the well-known Kelvin wave behavior of the first-symmetric equatorially-
 224 trapped Hough mode of DE3 that efficiently propagates to the IT. The ion density DE3
 225 observed by ICON is about 30% smaller than that observed by SORTIE. This ampli-
 226 tude difference between $\sim 420 \text{ km}$ and $\sim 590 \text{ km}$ is likely due to the combined influence
 227 of dissipation, zonal mean winds, wave-wave interactions, and inherent transience on the
 228 upward propagating thermospheric DE3. Also present in the 10°S - 10°N MLAT spectra
 229 is the migrating DW1 tide, as shown in previous ionospheric observations (e.g., Chang
 230 et al., 2013). SW2, DW5 and TE1 amplitudes are found to be negligible (TE1 results
 231 not shown here), providing confidence to our assertion that the WN4 structure observed
 232 by SORTIE and ICON is in fact due to the non-migrating DE3 tide.

233 To examine the latitudinal structure associated with this prominent F-region ion
 234 density DE3 (and other diurnal tides), Figure 4 shows the MLAT-wavenumber diurnal
 235 amplitude spectra of SORTIE IVM (panels *a* and *a'*), ICON IVM O^+ (panels *b* and *b'*)
 236 and SD/WACCM-X (panels *c* and *c'*) ion densities with and without the migrating DW1
 237 tide set to zero. Similar to the results presented in Figure 3, SORTIE, ICON, and WACCM-
 238 X show general agreement with large DW1 ($\sim 7 \times 10^4 \text{ cm}^{-3}$) and DE3 ($\sim 2.1\text{-}3 \times 10^4 \text{ cm}^{-3}$)
 239 amplitude variations primarily confined to low latitudes. Along with pronounced DW1
 240 and DE3 signals, the SORTIE and ICON spectra also exhibit D0 and DW2 variations
 241 (and some DE2 for SORTIE), while the WACCM-X spectra show DE1 and DW2 vari-

242 ations. Note that a $\sim 2 \times 10^4 \text{ cm}^{-3}$ DE2 amplitude variation is found in the model around
 243 5°N - 20°N MLAT, but not in the SORTIE or ICON observations. Larger DE3 amplitudes
 244 (up to $3.06 \times 10^4 \text{ cm}^{-3}$) are found in the SORTIE and ICON spectra around 25°N - 10°S
 245 MLAT (panels $4a'$ and $4b'$, respectively) consistent with the results shown in Figures $3a'$
 246 and $3b'$, while WACCM-X maxima are found near 10°N - 20°S MLAT. Some differences
 247 between the observed and modeled diurnal spectra are not unexpected and are likely due
 248 to (1) the different averaging window used (30 days for SORTIE and ICON, and 10 day
 249 for WACCM-X; note that a 10-day window for WACCM-X is adopted for consistency
 250 with the SORTIE and ICON results shown in Figure 2), (2) differences in the longitude-
 251 local time sampling between model and observations, and (3) inherent tidal variability
 252 observed by SORTIE and ICON that is not reproduced by the model. Even with the am-
 253 plitude suppression associated with the monthly averaging, panel $4a'$ shows SORTIE IVM
 254 ion density DE3 amplitudes exceeding $\sim 3 \times 10^4 \text{ cm}^{-3}$, i.e, over 15% of the observed zonal
 255 mean. As noted above in the context of Figure 3, the ICON observed ion density DE3
 256 amplitudes are about 30% smaller than SORTIE. Although not addressed in this study,
 257 the noted amplitude difference may be ascribable to reduced thermospheric DE3 ampli-
 258 tudes near 590 km due to the combined influence of dissipation, zonal mean winds, and
 259 wave-wave interactions that may result in reduced in-situ generated ion density DE3 at
 260 ICON heights. Follow on work will investigate these possible contributions to the ob-
 261 served amplitude differences (and day-to-day variability) using concurrent thermospheric
 262 observations by ICON. On the other hand, it is important to point out the remarkable
 263 degree of consistency between the SORTIE and ICON diurnal ion density spectra be-
 264 tween ~ 420 km and ~ 590 km during May 27-June 25, 2020. The latitude structure of
 265 DW1, DE3, D0, and DE2 are similarly captured by both space-borne observational plat-
 266 forms. It will be interesting to investigate in detail the diurnal, semidiurnal, and terdi-
 267 urnal tidal coupling between these two ionospheric regions and possible sources of lat-
 268 itude and temporal variability in the various tidal components. A few likely candidates
 269 responsible for the observed variability with height in the tidal amplitudes and latitude
 270 structures are noted above. A more systematic study that employs the entire ~ 5 -month
 271 period of simultaneous SORTIE and ICON observations and leverages other ICON remotely-
 272 sensed thermospheric observations is currently underway.

273 Additional analysis of the DE3 tide in the ionosphere is performed using a 30-day
 274 sliding window of low-MLAT ($\pm 20^\circ$) SORTIE and ICON IVM ion density during the

275 4-month interval from May 27, 2020 through September 15, 2020 to investigate its tem-
276 poral variability. Figure 5 shows daily variations of the estimated DE3 amplitudes (panel
277 *a*) and phases (panel *a'*) derived using a least squares method and the combined ascend-
278 ing/descending node data (i.e., covering ~ 24 -hr LT). Both SORTIE and ICON DE3 am-
279 plitudes are found to be larger ($\sim 1.9 \times 10^4 \text{ cm}^{-3}$ for SORTIE and $\sim 1 \times 10^4 \text{ cm}^{-3}$ for ICON)
280 around May 27 - June 25, quickly decreasing to minima ($10^3 - 10^4 \text{ cm}^{-3}$ for SORTIE and
281 $\sim 800 - 2 \times 10^3 \text{ cm}^{-3}$ for ICON) near June - July, with largest amplitudes ($\sim 2.1 \times 10^4 \text{ cm}^{-3}$
282 for SORTIE and $\sim 1.3 \times 10^4 \text{ cm}^{-3}$) observed during mid-August through early Septem-
283 ber 2020. General agreement is found in the seasonal variation of the ion density DE3
284 amplitudes observed by SORTIE near 420 km and by ICON near 590 km, with some dif-
285 ferences likely associated to additional complexities introduced by the influences of dis-
286 sipation, zonal mean winds, and wave-wave interactions. The latitude structure of DE3
287 presented in Figure 4 and its temporal variation shown in Figure 5*a*, with largest am-
288 plitudes occurring around August-September and a second period of enhanced activity
289 around April-May, is consistent with previous modeling work and observations (e.g., Mukhtarov
290 and Pancheva, 2011; Truskowski et al., 2014; Gasperini et al., 2015). To investigate the
291 altitude-latitude structure of the DE3 tide in the IT region, Figure 5*b* shows the SD/WACCM-
292 X height (~ 100 -450 km) - MLAT(60°S - 60°N) structure of ion density (panels *b* and *b'*)
293 and temperature (panels *c* and *c'*) amplitudes and phases during May 27 - June 5, 2020.
294 Ionospheric DE3 amplitudes are found to be largest around the F-layer peak near 200-
295 350 km, with a defined 2-peak structure and maxima of $\sim 8 \times 10^4 \text{ cm}^{-3}$ around 15°N and
296 15°S associated with the EIA (Appleton, 1946; Balan and Bailey, 1995). Previous stud-
297 ies (e.g., Sagawa et al., 2005; Immel et al., 2006; Wan et al., 2008) showed the EIA to
298 exhibit a WN4 longitudinal variation and it is now well accepted that this structure in
299 the ionospheric F-region can form as a result of the combined effect of the E-region dy-
300 namo modulation by the lower thermospheric DE3 and from the direct propagation of
301 DE3 into the F-region. It is fairly well established that neutral density variations, changes
302 in thermospheric atomic oxygen to nitrogen ratio, and meridional winds at F-region al-
303 titudes (e.g., Liu et al., 2009; England et al., 2010; Maute et al., 2012) can contribute
304 to the coupling between the tides and the ionospheric plasma. Previous modeling and
305 observational results (e.g., Oberheide et al., 2011) indicate small amplitudes ($< 1 \text{ m/s}$)
306 meridional DE3 winds in the low latitude (± 15 - 20°) middle and upper thermosphere,
307 which suggests the thermospheric DE3 winds to be a minor contributor the observed iono-

308 spheric DE3 signature. Similarly, neutral density and thermospheric atomic oxygen to
309 nitrogen ratio DE3 amplitudes are generally small at SORTIE and ICON altitudes. It
310 would be beyond the scope of this study to investigate the relative contribution of these
311 effects to the ionospheric DE3 observed by SORTIE and ICON and this effort is left for
312 follow on work. The lower thermospheric temperature DE3 exhibits largest amplitudes
313 of ~ 12 K around 110-120 km, in accord with previous modeling (e.g., Gasperini et al.,
314 2015, 2017a) and observational (e.g., Truskowski et al., 2014) results. The vertical wave-
315 length of the modeled temperature DE3 inferred from its vertical phase progression (panel
316 c') is ~ 49.8 km, in line with a predominant first-symmetric Hough mode. Latitudinal
317 asymmetries and broadening of the latitude structure with height in the thermospheric
318 DE3 are likely due to the combined effect of mean winds and dissipation (e.g., Forbes,
319 2000; Gasperini et al., 2015, 2017a).

320 Figure 6 displays maps of TIMED/SABER temperatures near 105 km observed con-
321 currently (i.e., May 27 - June 5, 2020) with the SORTIE and ICON IVM ion density maps
322 presented in Figure 2. Panels *a* and *b* show results for the descending (~ 3 -5 LT) and as-
323 cending (~ 15 -17 LT) nodes, respectively; while panel *c* contains results for half ascend-
324 ing/descending node differences. A prominent WN4 structure is found to dominate the
325 low-latitude ($\pm 30^\circ$) lower thermosphere with observed amplitude maxima of ~ 30 K peak-
326 ing near 135°E , 75°E , 45°W , 135°W . Pearson correlation coefficients between the equa-
327 torial longitudinal WN4 structure in TIMED/SABER temperatures and the WN4 struc-
328 ture in SORTIE (ICON) IVM ion density are calculated to be $r=0.71$ ($r=0.67$). The cor-
329 relation is computed on half ascending/descending node differences combining data within
330 10°S - 10°N MLAT. This level of correlation, along with the modeling work discussed in
331 Figure 5, provides strong evidence that the ionospheric WN4 structure observed by SOR-
332 TIE and ICON is in fact associated with the lower thermospheric DE3. This level of cor-
333 relation also indicates that there is significant phase coherence between the lower atmo-
334 spheric DE3 and the F-region ion density DE3 near 420 km and 590 km. This result agrees
335 with our hypothesis that the F-region ion density DE3 is primarily driven by the E-region
336 dynamo. As previously discussed, modeling and observational results generally demon-
337 strate small meridional wind, neutral density, and atomic oxygen to nitrogen ratio DE3
338 amplitudes in the low-latitude (± 15 - 20°) middle and upper thermosphere. These con-
339 siderations suggest that the in-situ-driven component may be a minor component. An
340 investigation on the relative contribution of the E-region dynamo modulation by the lower

341 thermospheric DE3 versus the direct propagation of DE3 into the F-region is considered
 342 beyond the purview of the current investigation and will be the subject of a follow on
 343 work.

344 4 Summary and Conclusions

345 Results presented above provide a clear picture of a marked longitudinal WN4 vari-
 346 ation observed during May 27 - June 5, 2020 by SORTIE and ICON IVM in the low-
 347 latitude ionosphere near 420 km and 590 km, respectively. Taking advantage of output
 348 from an SD/WACCM-X simulation nudged with MERRA-2 reanalysis data in the tro-
 349 posphere and stratosphere and monthly-averaged SORTIE and ICON IVM data, this
 350 prominent WN4 structure in the F-region ion density is demonstrated to be due to the
 351 well-known DE3 tide. This non-migrating tide has gained significant attention due to
 352 its ability to preferentially propagate to the IT from the tropical troposphere, thus ef-
 353 fectively coupling these regions. SORTIE IVM (ICON IVM) observations and SD/WACCM-
 354 X output are shown to exhibit ion density DE3 amplitudes upward of $3 \times 10^4 \text{ cm}^{-3}$ (2.1×10^4
 355 cm^{-3}), i.e., over 15% of the zonal mean, around $\pm 20^\circ$ MLAT, with general agreement
 356 between model and observations. The monthly-averaged SORTIE and ICON diurnal ion
 357 density spectra are found to be in strong agreement, with similar latitudinal structures
 358 observed and predominant DW1, DE3, D0, and DE2 tidal signals. DE3 ion density am-
 359 plitudes are found to be about 30% smaller in ICON observations near 590 km compared
 360 to SORTIE observations near 420 km, possibly due to the combined effect of dissipation,
 361 mean winds, and wave-wave interactions on the thermospheric DE3 propagating between
 362 420 km and 590 km (e.g., Gasperini et al., 2017a).

363 Least squares fitting of monthly-averaged SORTIE and ICON IVM ion density dur-
 364 ing May 27 - September 15, 2020 showed larger ($\sim 1.9 \times 10^4 \text{ cm}^{-3}$ for SORTIE and $\sim 1 \times 10^4$
 365 cm^{-3} for ICON) DE3 amplitudes around May 27 - June 25, rapidly decreasing to smaller
 366 values ($10^3 - 10^4 \text{ cm}^{-3}$ for SORTIE and $\sim 800 - 2 \times 10^3 \text{ cm}^{-3}$ for ICON) around June - July,
 367 and becoming largest ($\sim 2.1 \times 10^4 \text{ cm}^{-3}$ for SORTIE and $\sim 1.3 \times 10^4 \text{ cm}^{-3}$) around mid-
 368 August through early September 2020, in agreement with the well-known seasonal vari-
 369 ation of DE3 amplitudes highlighted in previous modeling and observational studies (e.g.,
 370 Mukhtarov and Pancheva, 2011; Truskowski et al., 2014; Gasperini et al., 2015, 2017a).
 371 Concurrent TIMED/SABER temperature observations near 105 km are shown to ex-
 372 hibit a pronounced WN4 structure ($\pm 30 \text{ K}$) in the low-latitude lower thermosphere as-

373 sociated with DE3 that is found to be correlated with the ionospheric WN4 ($r=0.71$ for
374 SABER/SORTIE and $r=0.67$ for SABER/ICON around 10°N - 10°S MLAT). Strong cor-
375 relation ($r=0.87$) is found between the WN4 structure observed simultaneously by SOR-
376 TIE and ICON IVM. This level of correlation and the similarities in the SORTIE and
377 ICON ion density DE3 latitude structures demonstrate that this tidal component is ef-
378 fective at coupling these two different ionospheric regions near 420 km and 590 km si-
379 multaneously. Note that this period is generally characterized by small thermospheric
380 DE3 amplitudes (e.g., Oberheide et al., 2011, Forbes et al., 2014, Gasperini et al., 2015).
381 The less than perfect agreement (i.e., $r=1.0$) between the two heights is ascribable to ad-
382 ditional complexities on the thermospheric DE3 introduced by wave dissipation, the pres-
383 ence of zonal mean winds, possible wave-wave interactions, and inherent transience.

384 The latitude-height structure of this prominent DE3 signal is further investigated
385 using SD/WACCM-X. DE3 amplitudes were found to be largest around the F-layer peak
386 near 200-350 km, with two well-defined enhancements near 15°N and 15°S MLAT as-
387 sociated with the EIA and maxima of $\sim 8 \times 10^4 \text{ cm}^{-3}$. The modeled DE3 temperature am-
388 plitudes are found to exhibit largest amplitudes around 110-120 km, in accord with pre-
389 vious modeling and observational results (e.g., Truskowski et al., 2014; Gasperini et al.,
390 2015, 2017a). Departure from a purely equatorially symmetric latitude structure and broad-
391 ening of latitude structures with height of the thermospheric DE3 is explained in terms
392 of the combined effect of mean winds and dissipation. The vertical wavelength of the mod-
393 eled temperature DE3 is found to be ~ 49.8 km indicative of a predominant first sym-
394 metric Hough mode.

395 This study presented first results of a prominent F-region ion density WN4 struc-
396 ture driven by the nonmigrating DE3 tide observed concurrently by the SORTIE Cube-
397 Sat and ICON. The results herein contained provide evidence for the extent to which
398 DE3 is capable of affecting simultaneously the global structure of the F-region ion den-
399 sity near 420 km and 590 km. This study further demonstrates the degree to which global-
400 scale tidal activity related to the tropical troposphere can influence the IT system. This
401 study also demonstrates SORTIE and ICON to be excellent observational platforms for
402 studying the influence of terrestrial weather on IT variability at low latitudes. Comple-
403 mentary and concurrent measurements from SORTIE's and ICON's near identical IVM
404 instruments at different altitudes are shown to be particularly valuable. Future work will

405 take advantage of comprehensive sets of SORTIE and ICON observations to investigate
 406 in further detail the coupling of different IT regions by global-scale waves.

407 **Acknowledgments**

408 The SORTIE mission is supported from NASA HQ by Grant 80NSSC18K0094 to At-
 409 mospheric & Space Technology Research Associates (ASTRA) LLC. ICON is supported
 410 by NASA's Explorers Program through contracts NNG12FA45C and NNG12FA42I. AS-
 411 TRA is grateful for support from U.C. Berkeley and the ICON mission via Subcontract
 412 No: 00008210. ICON IVM-A data products (Level 2.7, Version 4) are publicly available
 413 at <https://spdf.gsfc.nasa.gov/pub/data/icon/l2/>. Post-processed SORTIE IVM Level
 414 2 ion density data can be accessed on Zenodo at doi:10.5281/zenodo.4589362. TIMED/SABER
 415 temperatures (version v2.0) can be freely accessed at <http://saber.gats-inc.com/data.php/>.
 416 WACCM-X history files can be accessed at the NCAR/CDG ([https://doi.org/10.26024/5b58-
 417 nc53](https://doi.org/10.26024/5b58-nc53)) and are archived on the NCAR/HAO Campaign Space. For further information
 418 on the ICON IVM data used see: <https://spdf.gsfc.nasa.gov/pub/data/icon/documentation/>.
 419 The post-processed data used for Figures 1-6 are available at doi:10.5281/zenodo.4746615.

421 **References**

- 422 Appleton, E. V. (1946), Two anomalies in the ionosphere, *Nature*, 157, 691.
- 423 Balan, N., and G. J. Bailey (1995), Equatorial plasma fountain and its effects:
 424 Possibility of an additional layer, *J. Geophys. Res.*, 100, 21421-21432,
 425 doi:10.1029/95JA01555.
- 426 Chang, L. C., Lin, C.-H., Liu, J.-Y., Balan, N., Yue, J., and Lin, J.-T. (2013), Sea-
 427 sonal and local time variation of ionospheric migrating tides in 2007-2011
 428 FORMOSAT-3/COSMIC and TIE-GCM total electron content, *J. Geophys.*
 429 *Res. Space Physics*, 118, 2545-2564, doi:10.1002/jgra.50268.
- 430 Chapman, S., and R. S. Lindzen (1970), *Atmospheric Tides*, Springer, New York.
- 431 Cho, Y.-M., and G. Shepherd (2015), Resolving daily wave 4 nonmigrating tidal
 432 winds at equatorial and midlatitudes with WINDII: DE3 and SE2, *J. Geophys.*
 433 *Res. Space Physics*, 120, 10,053–10,068, doi:10.1002/2015JA021903.
- 434 Crowley, G., C. Fish, M. Pilinski, E. Stromberg et al. (2016), Scintillation Observa-
 435 tions and Response of The Ionosphere to Electrodynamics (SORTIE), Proceed-

- 436 ings of the 30th Annual AIAA/USU SmallSat Conference, paper: SSC16-VI-3.
437 England, S. L., T. J. Immel, J. D. Huba, M. E. Hagan, A. Maute, and R. DeMa-
438 jistre (2010), Modeling of multiple effects of atmospheric tides on the iono-
439 sphere: An examination of possible coupling mechanisms responsible for the
440 longitudinal structure of the equatorial ionosphere, *J. Geophys. Res.*, 115,
441 A05308, doi:10.1029/2009JA014894.
- 442 Fang, T.-W., H. Kil, G. Millward, A. D. Richmond, J.-Y. Liu, and S.-J. Oh
443 (2009), Causal link of the wave-4 structures in plasma density and vertical
444 plasma drift in the low-latitude ionosphere, *J. Geophys. Res.*, 114, A10315,
445 doi:10.1029/2009JA014460.
- 446 Forbes, J. M. (2000), Wave coupling between the lower and upper atmosphere: Case
447 study of an ultra-fast Kelvin wave, *J. Atmos. Terr. Phys.*, 62, 1603-1621.
- 448 Garcia-Comas, M., et al. (2008), Errors in SABER kinetic temperature
449 caused by non-LTE model parameters, *J. Geophys. Res.*, 113, D24, doi:
450 10.1029/2008JD010105.
- 451 Gasperini, F., J. M. Forbes, E. N. Doornbos, and S. L. Bruinsma (2015), Wave cou-
452 pling between the lower and middle thermosphere as viewed from TIMED and
453 GOCE, *J. Geophys. Res.*, 120, 5788-5804, doi:10.1002/2015JA021300.
- 454 Gasperini, F., Forbes, J. M., and Hagan, M. E. (2017a), Wave coupling from the
455 lower to the middle thermosphere: Effects of mean winds and dissipation, *J.*
456 *Geophys. Res.*, 122, 7781-7797, doi:10.1002/2017JA024317.
- 457 Gasperini, F., M. E. Hagan, and Y. Zhao (2017b), Evidence of tropo-
458 spheric 90-day oscillations in the thermosphere, *Geophys. Res. Lett.*,
459 doi:10.1002/2017GL075445.
- 460 Gasperini, F., Forbes, J. M., Doornbos, E. N., and Bruinsma, S. L. (2018),
461 Kelvin wave coupling from TIMED and GOCE: Inter/intra-annual vari-
462 ability and solar activity effects, *J. Atmos. Sol.-Terr. Phys.*, 171, 176-187,
463 doi:10.1016/j.jastp.2017.08.034.
- 464 Gasperini, F., Liu, H., and McInerney, J. (2020), Preliminary evidence of Madden-
465 Julian Oscillation effects on ultrafast tropical waves in the thermosphere, *J.*
466 *Geophys. Res.*, 125, e2019JA027649, doi:10.1029/2019JA027649.
- 467 Gelaro, R., McCarty, W., Suarez, M. J., Todling, R., Molod, A., Takacs, L., et al.
468 (2017), The modern-era retrospective analysis for research and applications,

- 469 version 2 (MERRA-2), *J. Climate*, 30, 5419-5454, doi:10.1175/JCLI-D-16-
470 0758.1.
- 471 Hagan, M. E. and J. M. Forbes (2002), Migrating and nonmigrating diurnal tides in
472 the middle and upper atmosphere excited by tropospheric latent heat release,
473 *J. Geophys. Res.*, 107(D24), 4754, doi: 10.1029/2001JD001236.
- 474 Hagan, M.E., A. Maute, R.G. Roble, A.D. Richmond, T.J. Immel, and S.L. England
475 (2007), Connections between deep tropical clouds and the Earth's ionosphere,
476 *Geophys. Res. Lett.*, 34, L20109, doi:10.1029/2007GL030142.
- 477 Hagan, M. E. (1996), Comparative effects of migrating solar sources on tidal signa-
478 tures in the middle and upper atmosphere, *J. Geophys. Res.*, 101, 21213-21222.
- 479 He, M., L. Liu, W. Wan, and Y. Wei (2011), Strong evidence for couplings between
480 the ionospheric wave-4 structure and atmospheric tides, *Geophys. Res. Lett.*,
481 38, L14101, doi:10.1029/2011GL047855.
- 482 Heelis, R. A., Lowell, J. K., and Spiro, R. W. (1982), A model of the high-latitude
483 ionospheric convection pattern, *J. Res. Lett.*, 87, 63390-6345.
- 484 Heelis, R.A., Stoneback, R.A., Perdue, M.D., Depew, M.D., Morgan, W.A., Mankey,
485 M.W., Lippincott, C.R., Harmon, L.L., and Holt, B.J., (2017), Ion Velocity
486 Measurements for the Ionospheric Connections Explorer, *Sp. Sci. Rev.*, 212(1-
487 2), 615-629, doi:10.1007/s11214-017-0383-3.
- 488 Hurrell, J. W., Holland, M. M., Gent, P. R., Ghan, S., Kay, J. E., Kushner, P. J., et
489 al. (2013), The community earth system model: A framework for collaborative
490 research, *Bull. of the Am. Met. Soc.*, 94(9), 1339-1360, doi:10.1175/BAMS-D-
491 12-00121.1.
- 492 Immel, T. J., E. Sagawa, S. L. England, S. B. Henderson, M. E. Hagan, S. B.
493 Mende, H. U. Frey, C. M. Swenson, and L. J. Paxton (2006), The control of
494 equatorial ionospheric morphology by atmospheric tides, *Geophys. Res. Lett.*,
495 33 (15), doi:10.1029/2006GL026161.
- 496 Immel, T.J., England, S.L., Mende, S.B. et al, (2018), The Ionospheric Connec-
497 tion Explorer Mission: Mission Goals and Design, *Sp. Sci. Rev.* 214, 13,
498 doi:10.1007/s11214-017-0449-2.
- 499 Jin, H., Y. Miyoshi, H. Fujiwara, and H. Shinagawa (2008), Electrodynamics of the
500 formation of ionospheric wave number 4 longitudinal structure, *J. Geophys.*
501 *Res.*, 113, A09307, doi:10.1029/2008JA013301.

- 502 Lei, J., J.P. Thayer, W. Wang, J. Yue, and X. Dou (2014), Nonmigrating tidal mod-
 503 ulation of the equatorial thermosphere and ionosphere anomaly, *J. Geophys.*
 504 *Res.*, 119, 3036–3043, doi:10.1002/2013JA019749.
- 505 Lieberman, R. S., Riggin, D. M., Ortland, D. A., Nesbitt, S. W., and Vin-
 506 cent, R. A. (2007), Variability of mesospheric diurnal tides and tropo-
 507 spheric diurnal heating during 1997-1998, *J. Geophys. Res.*, 112, D20110,
 508 doi:10.1029/2007JD008578.
- 509 Lieberman, R. S. (1991), Nonmigrating diurnal tides in the equatorial middle atmo-
 510 sphere, *J. Atmos. Sci.*, 48, 1112 - 1123.
- 511 Lin, C.H., C.C. Hsiao, J.Y. Liu, and C.H. Liu (2007), Longitudinal structure of the
 512 equatorial ionosphere: Time evolution of the four-peaked EIA structure, *J.*
 513 *Geophys. Res.*, 112, A12305, doi:10.1029/2007JA012455.
- 514 Liu, H., and S. Watanabe (2008), Seasonal variation of the longitudinal structure
 515 of the equatorial ionosphere: Does it reflect tidal influences from below?, *J.*
 516 *Geophys. Res.*, 113, A08315, doi:10.1029/2008JA013027.
- 517 Liu, H., M. Yamamoto, and H. Luhr (2009), Wave-4 pattern of the equatorial mass
 518 density anomaly: A thermosphere signature of tropical deep convection, *Geo-*
 519 *phys. Res. Lett.*, 36, L18104, doi:10.1029/2009GL039865.
- 520 Liu, H. -L. (2016), Variability and predictability of the space environment
 521 as related to lower atmosphere forcing, *Space Weather*, 14, 634-658,
 522 doi:10.1002/2016SW001450.
- 523 Liu, J., Liu, H., Wang, W., Burns, A. G., Wu, Q., Gan, Q., et al. (2018),
 524 First results from the ionospheric extension of WACCM-X during the
 525 deep solar minimum year of 2008, *J. Geophys. Res.*, 123, 1534-1553,
 526 doi:10.1002/2017JA025010.
- 527 Maute, A., A. D. Richmond, and R. G. Roble, Sources of low-latitude ionospheric
 528 E x B drifts and their variability (2012), *J. Geophys. Res.*, 117, A06312,
 529 doi:10.1029/2011JA017502.
- 530 Mertens Christopher J., M Martin G. Mlynczak, Manuel López-Puertas, Peter P.
 531 Wintersteiner, R. H. Picard, Jeremy R. Winick, Larry L. Gordley, and James
 532 M. Russell III (2001), Retrieval of mesospheric and lower thermospheric kinetic
 533 temperature from measurements of CO₂ 15 μ m Earth limb emission under
 534 non-LTE conditions, *Geophys. Res. Lett.*, 28, 1391-1394.

- 535 Mukhtarov, P., and D. Pancheva (2011), Global ionospheric response to nonmigrat-
 536 ing DE3 and DE2 tides forced from below, *J. Geophys. Res.*, 116, A05323,
 537 doi:10.1029/2010JA016099.
- 538 Oberheide, J., J. M. Forbes, X. Zhang, and S. L. Bruinsma (2011), Wave-driven vari-
 539 ability in the ionosphere-thermosphere-mesosphere system from TIMED ob-
 540 servations: What contributes to the “wave 4”?, *J. Geophys. Res.*, 116, A01306,
 541 doi:10.1029/2010JA015911.
- 542 Oberheide, J., J. Forbes, K. Hausler, Q. Wu, and S. L. Bruinsma (2009), Tropo-
 543 spheric tides from 80–400 km: Propagation, inter-annual variabil-ity and solar
 544 cycle effects, *J. Geophys. Res.*, doi:10.1029/2009JD012388.
- 545 Oberheide, J., Hagan, M. E., Ward, W. E., Riese, M., and Offermann, D. (2000),
 546 Modeling the diurnal tide for the Cryogenic Infrared Spectrometers and Tele-
 547 scopes for the Atmosphere (CRISTA) 1 time period, *J. Geophys. Res.*, 105(
 548 A11), 24917-24929, doi:10.1029/2000JA000047.
- 549 Onohara, A. N., Batista, I. S., and Batista, P. P. (2018), Wavenumber-4 structures
 550 observed in the low-latitude ionosphere during low and high solar activity peri-
 551 ods using FORMOSAT/COSMIC observations, *Ann. Geophys.*, 36(2), 459-471,
 552 doi:10.5194/angeo-36-459-2018.
- 553 Ren, Z., W. Wan, L. Liu, and J. Xiong (2009), Intra-annual variation of wave num-
 554 ber 4 structure of vertical E B drifts in the equatorial ionosphere seen from
 555 ROCSAT-1, *J. Geophys. Res.*, 114, A05308, doi:10.1029/2009JA014060.
- 556 Sagawa, E., T. J. Immel, H. U. Frey, and S. B. Mende (2005), Longitudinal struc-
 557 ture of the equatorial anomaly in the nighttime ionosphere observed by IM-
 558 AGE/FUV, *J. Geophys. Res.*, 110, A11302, doi:10.1029/2004JA010848.
- 559 Truskowski, A.O., Forbes, J.M., Zhang, X., and S.E. Palo (2014), New perspec-
 560 tives on thermosphere tides - 1. Lower thermosphere spectra and seasonal-
 561 latitudinal structures, *Earth, Planets and Space*, 66-136, doi:10.1186/s40623-
 562 014-0136-4.
- 563 Wan, W., L. Liu, X. Pi, M.-L. Zhang, B. Ning, J. Xiong, and F. Ding (2008),
 564 Wavenumber-4 patterns of the total electron content over the low latitude
 565 ionosphere, *Geophys. Res. Lett.*, 35, L12104, doi:10.1029/2008GL033755.
- 566 Wan, W., J. Xiong, Z. Ren, L. Liu, M.-L. Zhang, F. Ding, B. Ning, B. Zhao,
 567 and X. Yue (2010), Correlation between the ionospheric WN4 signature

568 and the upper atmospheric DE3 tide, *J. Geophys. Res.*, 115, A11303,
569 doi:10.1029/2010JA015527.

570 Wan, W., Z. Ren, F. Ding, J. Xiong, L. Liu, B. Ning, B. Zhao, G. Li, M.-L. Zhang
571 (2012), A simulation study for the couplings between DE3 tide and longi-
572 tudinal WN4 structure in the thermosphere and ionosphere, *J. of Atm. and*
573 *Sol.-Terr. Phy.*,, 90-91, 52-60, doi:10.1016/j.jastp.2012.04.011.

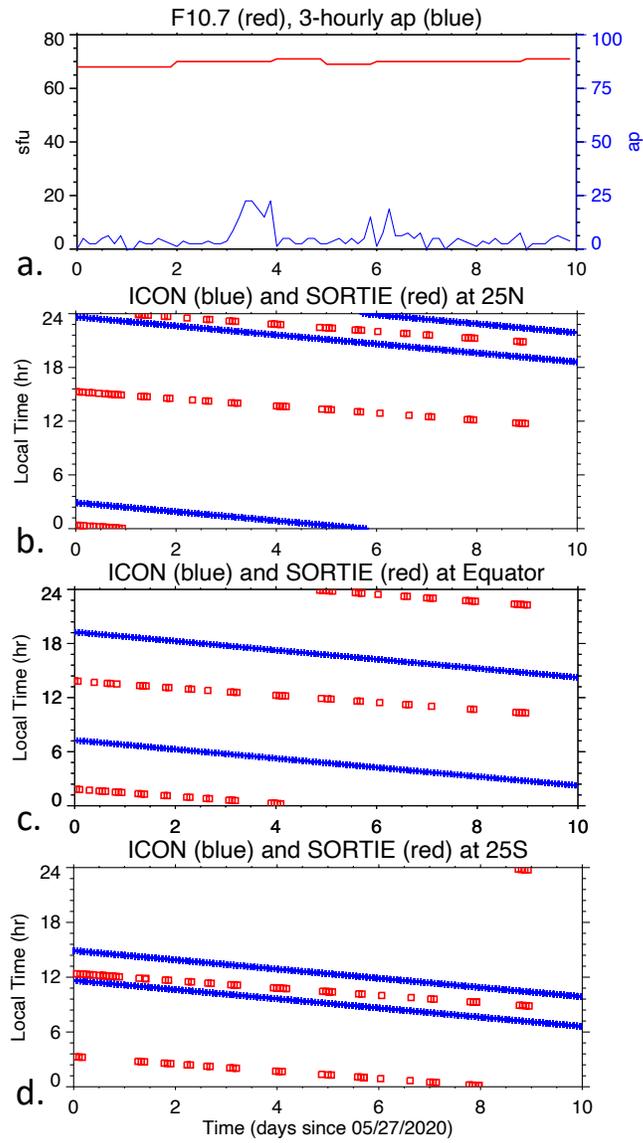


Figure 1. (a) Time series of the daily F10.7 solar flux (red curve) and 3-hourly ap (blue curve) during May 27 - June 5, 2020. (b)-(d) Time series of ICON (blue plus signs) and SORTIE (red squares) local solar time at 25°N (panel *b*), equator (panel *c*), and 25°S (panel *d*).

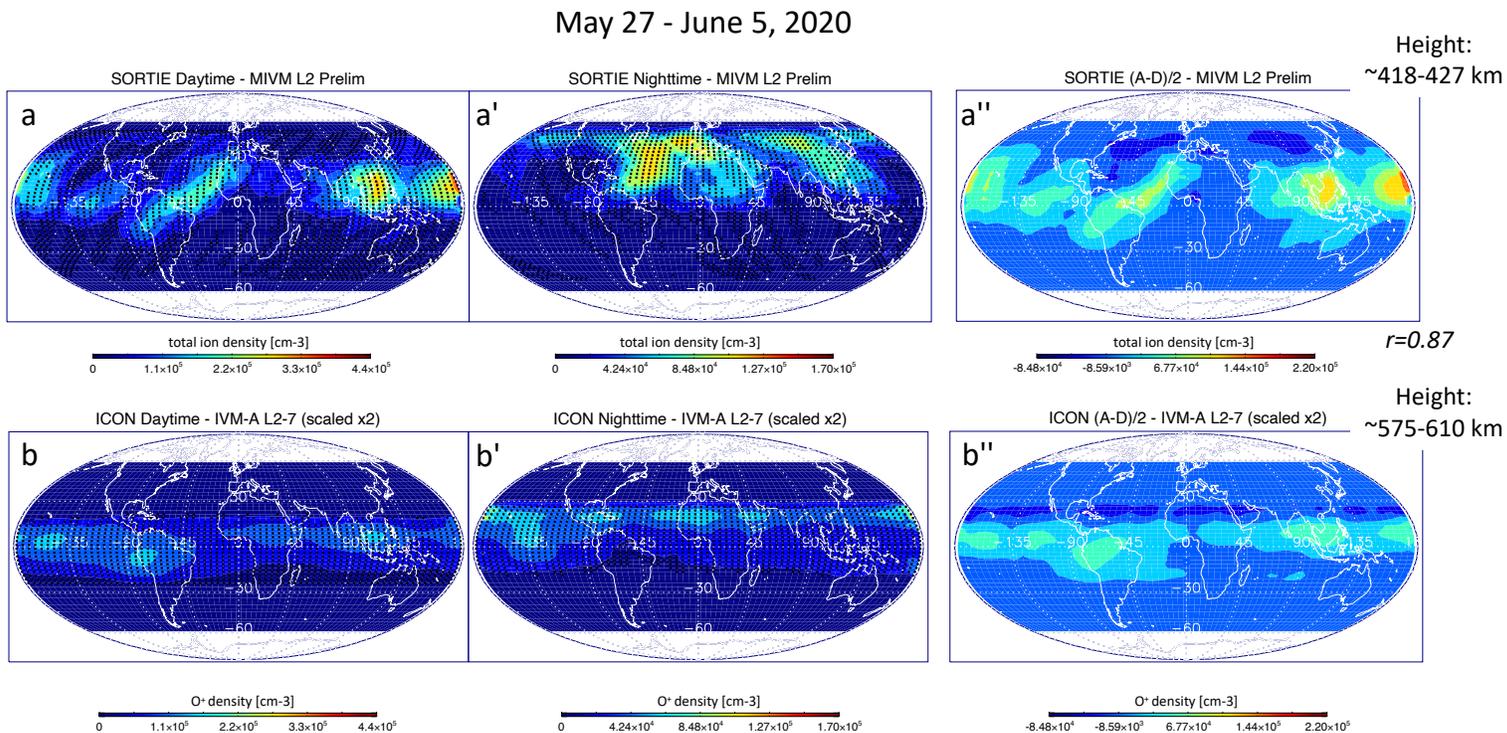
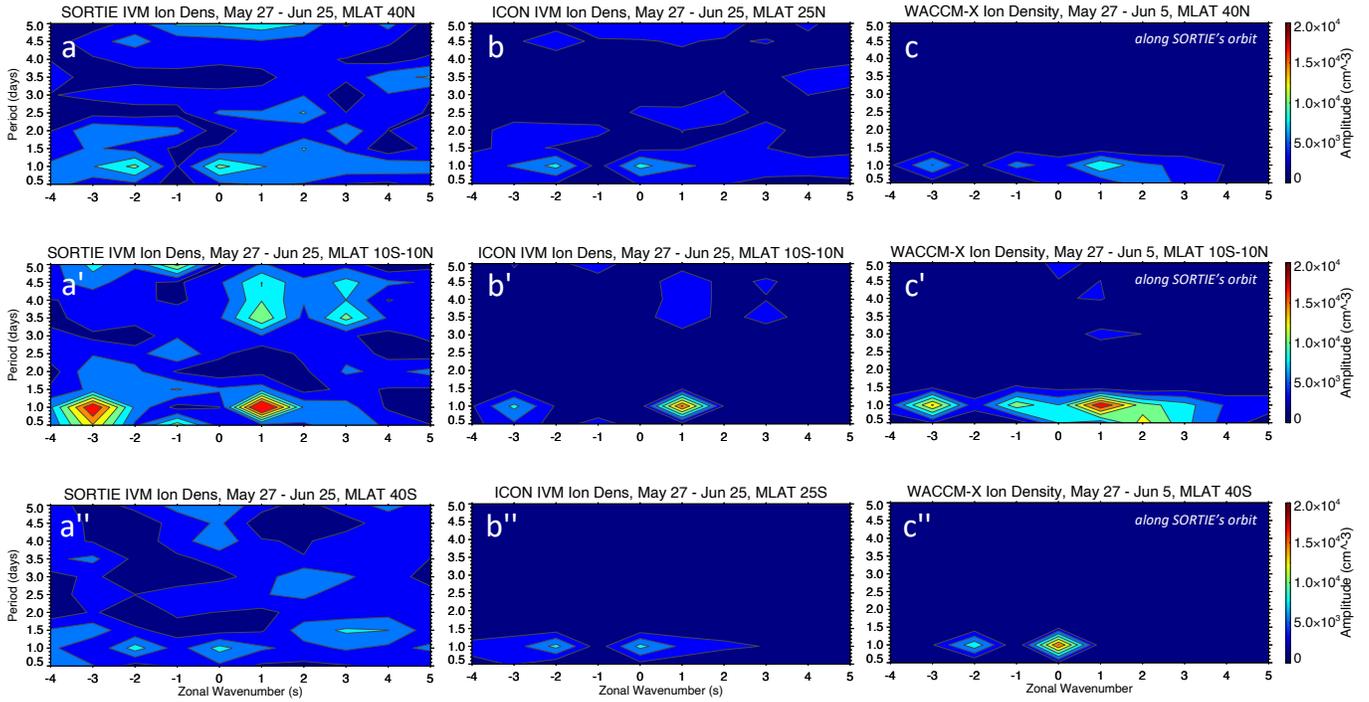


Figure 2. Latitude (60°S - 60°N) versus longitude (180°W - 180°E) maps of SORTIE total ion density (upper panels) and ICON O^+ density (lower panels) Ion Velocity Meter (IVM) ion density measured during May 27 - June 5, 2020. SORTIE (ICON) daytime, i.e., ~ 10 - 14 LT (~ 14 - 18 LT), averages are shown in panel *a* (panel *b*); while SORTIE (ICON) nighttime, i.e., ~ 22 - 2 LT (~ 2 - 6 LT) averages are shown in panel *a'* (*b'*). Panels *a''* and *b''* show maps of half ascending and descending node differences (i.e., daytime-nighttime differences). The black dots in panels *a-a'* and *b-b'* show the measurement locations. ICON O^+ ion density is scaled (multiplied) by a factor of 2 given ICON's higher mean altitude (~ 575 - 610 km) compared to SORTIE (~ 418 - 427 km). This scaling is performed to use the same color bar for both ICON and SORTIE. The correlation between the WN4 in SORTIE and ICON (panels *a''* and *b''*) is $r=0.87$ near the equator (10°S - 10°N), as noted at the bottom of panel *a''*.



Negative for eastward propagation

Figure 3. Period versus zonal wavenumber amplitude spectra of SORTIE IVM ion density at MLAT 40°N (panel *a*), 10°S-10°N (panel *a'*), and 40°S (panel *a''*) during May 27 - June 25, 2020. (*b*)-(*b''*) Same as (*a*)-(*a''*), but for ICON IVM O⁺ ion density at MLAT 25°N (panel *b*), 10°S-10°N (panel *b'*), and 25°S (panel *b''*). (*c*)-(*c''*) Same as (*a*)-(*a''*), but for SD/WACCM-X ion density during May 27 - June 5, 2020 and sampled along SORTIE's orbit. This 10-day window for WACCM-X is adopted for consistency with the SORTIE and ICON results shown in Figure 2. Eastward propagating waves are indicated with negative wavenumbers. A prominent diurnal s=−3 DE3 signal is clearly seen in the SORTIE, ICON, and SD/WACCM-X low-latitude ionosphere spectra, along with large DW1.

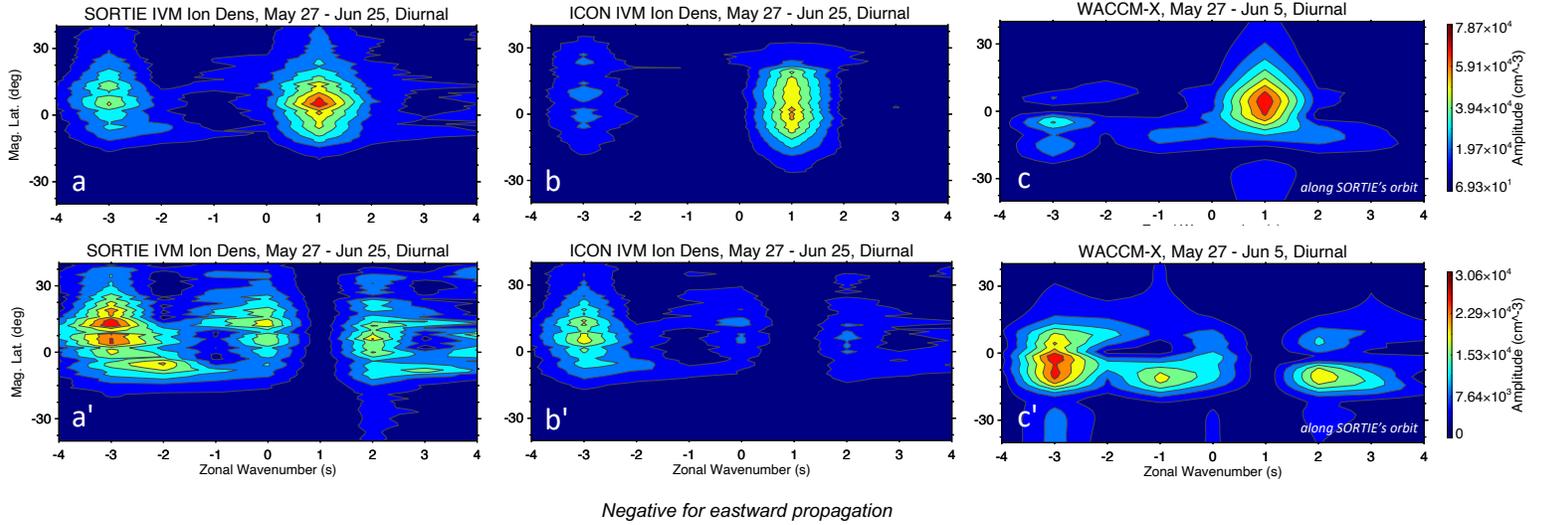


Figure 4. (a) MLAT versus zonal wavenumber diurnal amplitude spectra of SORTIE IVM ion density during May 27 - June 25, 2020. (b) Same as (a), but for ICON IVM O^+ ion density. (c) Same as (a), but from SD/WACCM-X output sampled along SORTIE's orbit during May 27 - June 5, 2020. (a')-(c') Same as (a)-(c), but with the migrating tide set to zero. The prominent ion density $s=-3$ DE3 variation exhibits amplitudes upward of $3 \times 10^4 \text{ cm}^{-3}$ in both the SORTIE observations (panel a') and the model (panel c'). ICON's DE3 ion density amplitude maxima are shown to be around $2.1 \times 10^4 \text{ cm}^{-3}$, i.e., 30% lower than those found in SORTIE. The monthly-averaged SORTIE and ICON diurnal ion density spectra are in general agreement, with similar latitudinal structures observed and predominant DW1, DE3, D0, and DE2 tidal signals.

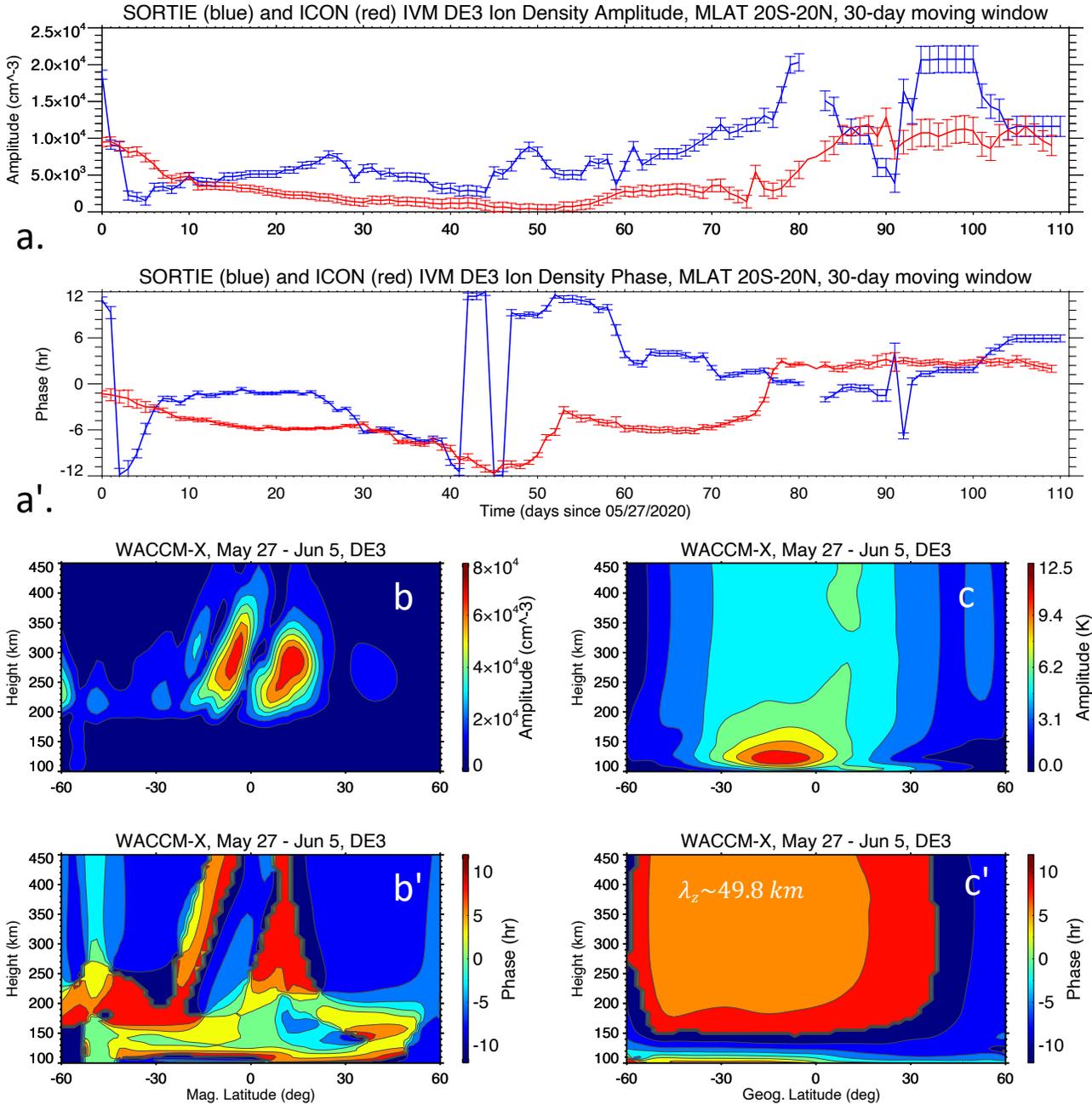


Figure 5. Time series of SORTIE (blue line) and ICON (red line) IVM ion density DE3 amplitudes (panel *a*) and phases (panel *a'*) derived using least squares fitting within 30-day sliding windows during May 27 - September 15, 2020. The vertical boxes identify 1- σ uncertainty estimates in the amplitudes and phases output by the fitting procedure. Panels *b-b'* show the SD/WACCM-X height (~ 100 -450 km) versus MLAT (60° S- 60° N) ion density DE3 amplitudes and phases, respectively. Panels *c-c'* show the same results as in panels *b-b'*, but for neutral temperature DE3. The vertical wavelength of the modeled temperature DE3 inferred from its vertical phase progression (panel *c'*) is ~ 49.8 km, consistent with a predominant first symmetric Hough mode.

TIMED/SABER Temperature V2.0, May 27 – June 5, 2020

Height: 105 km

Descending Node (D)

Ascending Node (A)

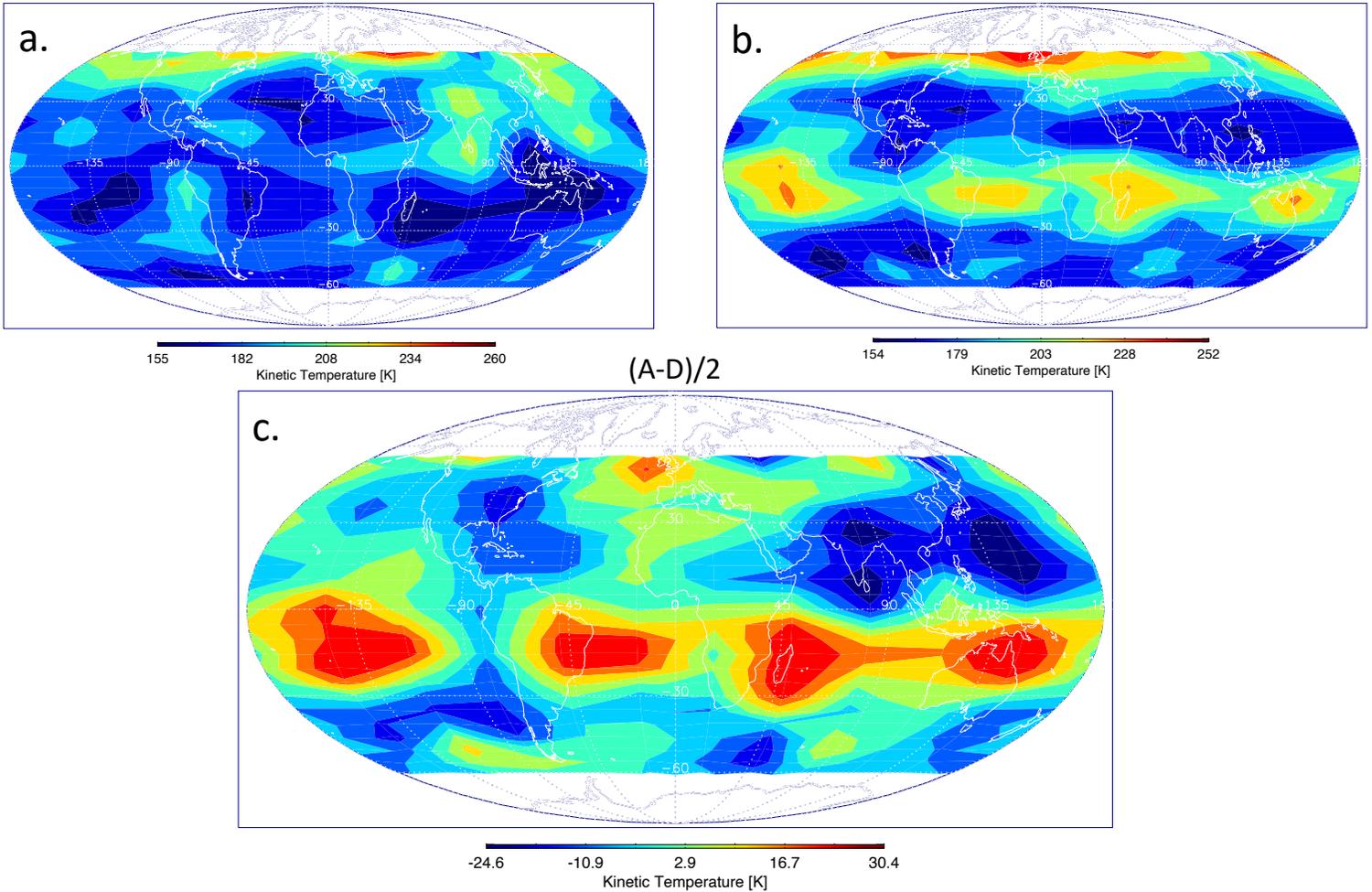


Figure 6. Latitude (60°S-60°N) versus longitude (180°W-180°E) maps of TIMED/SABER kinetic temperatures near 105 km during May 27 - June 5, 2020 at the descending node (panel *a*) near 3-5 LT, ascending node (panel *b*) near 15-17 LT, and their half difference (panel *c*). A prominent longitudinal WN4 structure is evident in the differences (panel *c*) consistent with the concurrent ionospheric DE3 signature observed by SORTIE (ICON) IVM near 420 km (590 km).