

Ana S. Casas<sup>1</sup>, Fabian B. Wadsworth<sup>2</sup>, Adrian Hornby<sup>3</sup>, Ulrich Kueppers<sup>1</sup>, Pierre Delmelle<sup>4</sup> and Donald B. Dingwell<sup>1</sup>

<sup>1</sup>Ludwig-Maximilians-Universität, Germany, <sup>2</sup>Durham University, U.K., <sup>3</sup>Cornell University, USA, <sup>4</sup>Catholic University of Louvain, Belgium.

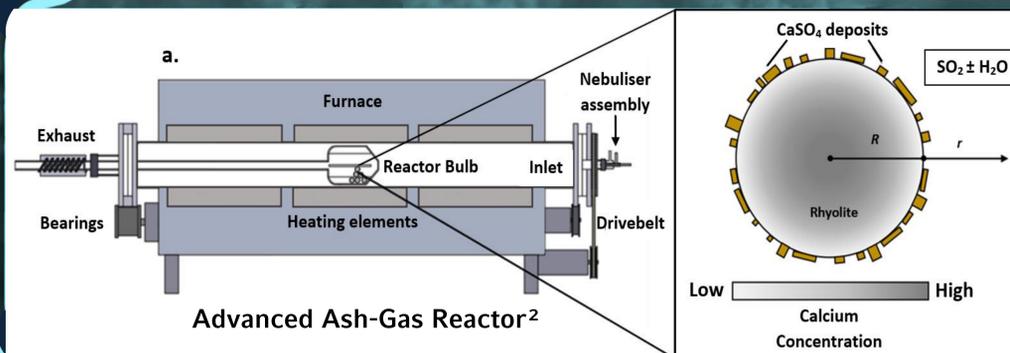
## 1 Introduction

The presence and nature of crystalline phases in pyroclasts (i.e., ash, bombs, lapilli), bear first-hand information about pre- and syn-eruptive magma conditions. For instance, crystal grain size pertains to cooling rate estimations during both magma storage in the chamber and magma ascent through the conduit. Changes in the eruptive style, magma degassing and magma fluid dynamics have also been linked to nucleation and growth of crystalline phases, yet, the influence of the crystallinity of pyroclasts on post-eruptive processes has been given little attention.

Here, we tested the reaction potential of various natural volcanic samples (crushed to fine-ash GSD) by exposing them to a SO<sub>2</sub>-Ar gas mixture (1 mol.% in 99 mol.%) at high temperatures, and characterized the samples prior and after gas treatments.

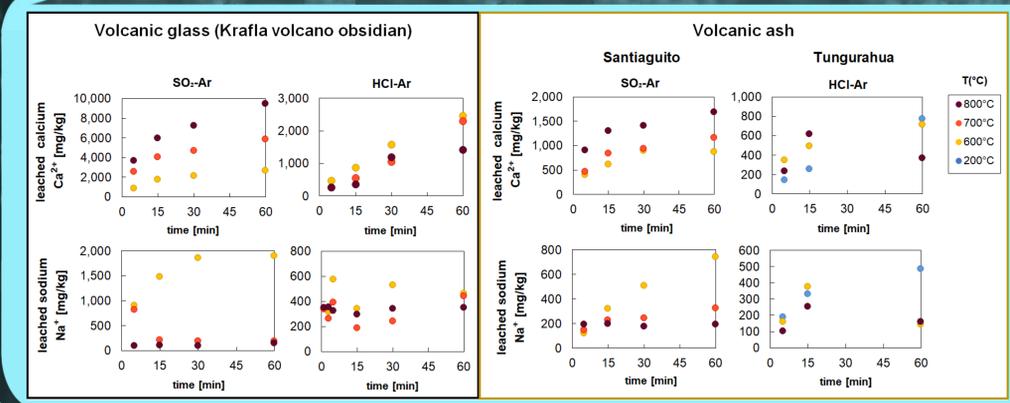
Using a natural volcanic bomb, we could test, for the first time, the role that a more complex ash-componentry performs in SO<sub>2</sub>-ash reactions, in particular that of the mineral phases in high temperature calcium diffusion that results in surficial salt formation, thus in sinking SO<sub>2</sub> emissions during explosive volcanic eruptions.

## 2 Methods: high temperature SO<sub>2</sub>-glass interactions

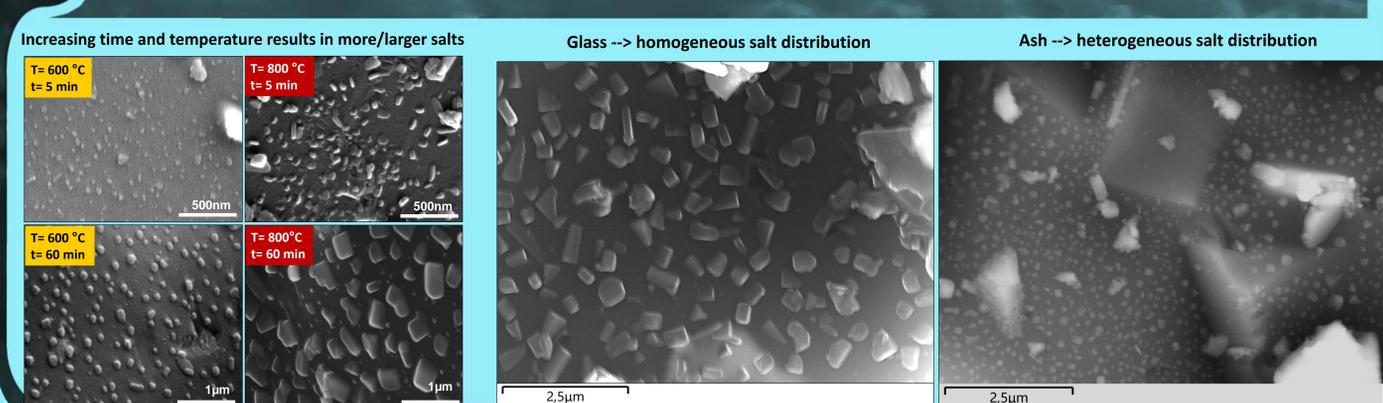


Samples	Variables	Characterization
<ul style="list-style-type: none"> <li>1) <b>Santiaguito bomb</b> (2013 eruption)</li> <li>2) <b>Tungurahua ash</b> (2014 eruption)</li> </ul>	<ul style="list-style-type: none"> <li>Temperature (°C) — 200, 600, 700, 800</li> <li>Gas mixture — SO<sub>2</sub>-Ar and HCl-Ar (1 mol.% in 99 mol.%)</li> <li>Exposure time (min.) — 5, 15, 30, 60</li> </ul>	<ul style="list-style-type: none"> <li>Leaching (1:250 ratio)</li> <li>IC analysis</li> <li>Particle size distribution</li> <li>SEM, BET analysis</li> <li>Microprobe analysis</li> <li>Fe<sup>3+</sup>/Fe<sup>2+</sup> ratio determination</li> <li>Potentiometric K<sub>2</sub>Cr<sub>2</sub>O<sub>7</sub> titration</li> </ul>

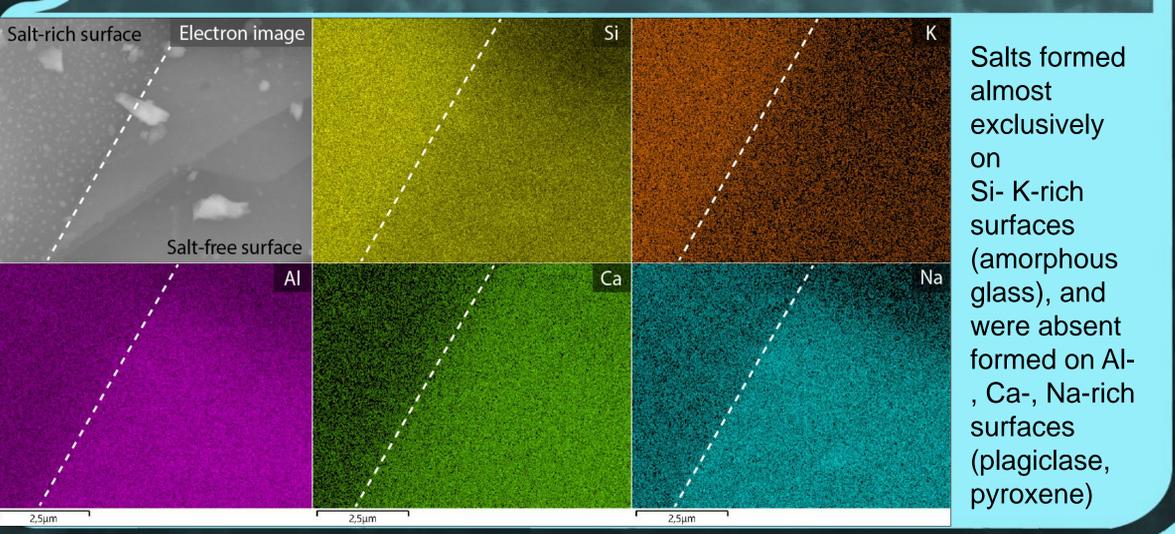
## 3 Results: Leachate analysis



## 4 SEM-BSE images



## 5 SEM-EDS element mapping



## 6 Discussion

We here observed that, while, exposure temperature, reaction time and bulk composition strongly influence salt formation (nucleation-, growth-rates), the surface chemistry and componentry is a critical parameter that governs solid-gas reaction in the first place; chemisorption of gas molecules (HCl, SO<sub>2</sub>) is perhaps unlikely to occur on pre-existent surficial mineral phases as it is on glass network surfaces. This could be due to the strong crystalline bonding in mineral structures, which prevents atoms, to act as sorbents for gas molecules, thus depriving the surface from reactive adsorption sites. Absence of salts on crystal-rich surfaces, was observed for both HCl- and SO<sub>2</sub> experiments, which implies that regardless of the solid-gas chemisorption mechanism (different for SO<sub>2</sub> and HCl), gas uptake is less likely to occur on surficial mineral phases, thus nucleation, and growth of surficial salts, and high temperature gas scavenging will be less for crystall-rich ash, than for glassy ash.

## 7 Future work

Our final goal is enable the modelling of high temperature volcanic gas uptake for any given eruption. Knowing the mechanisms and variables controlling Calcium diffusion, we can modify the model of Suzuki et al. (2005) and predict the mass of gas uptaken.

This image is an example of a volcanic plume simulation, where the size particle distribution of ash is been tracked in 3 spatial coordinates, for an eruption of the magnitude of Pinatubo, 1991. Courtesy of Yujiro J. Suzuki.