

**Dissolved Nitrogen Cycling in The Eastern Canadian Arctic Archipelago from Stable Isotopic Data**

H.C. Westbrook<sup>1</sup>, A. Bourbonnais<sup>1</sup>, C.C.M. Manning<sup>2</sup>, M. Ahmed<sup>3</sup>, B. Else<sup>3</sup>, R. Izett<sup>4</sup>,  
and J. Granger<sup>2</sup>

<sup>1</sup> School of the Earth, Ocean and Environment, University of South Carolina, Columbia, SC, United States; <sup>2</sup> Department of Marine Sciences, University of Connecticut, Groton, Connecticut, United States; <sup>3</sup> Department of Geography, University of Calgary, Calgary, Alberta, Canada; <sup>4</sup> Department of Oceanography, Dalhousie University, Halifax, Nova Scotia, Canada

**Contents of this file**

Text S1 to S2  
Figures S1 to S4

**Introduction**

The following is a description of the parameters used in our simple steady-state isotopic model for NO<sub>3</sub><sup>-</sup> cycling in rivers in the study area. This model is used to qualitatively evaluate if the observed Δ(15-18) of NO<sub>3</sub><sup>-</sup> can be reproduced by considering different dissolved NO<sub>3</sub><sup>-</sup> sources (nitrification and atmospheric depositions) in glacial rivers. Subsequently, we provide figures of surface salinity and δ<sup>18</sup>O of water in the study area, and nitrate depth profiles. Description of River Isotopic NO<sub>3</sub><sup>-</sup> model.

**Definition of terms**

*a* = nitrified NO<sub>3</sub><sup>-</sup> (from permafrost or atmospheric NH<sub>4</sub><sup>+</sup>)

*b* = atmospheric NO<sub>3</sub><sup>-</sup> (from glacial snow melt)

*d* = assimilation of NO<sub>3</sub><sup>-</sup>

*e* = recycled production of NO<sub>3</sub><sup>-</sup>

δ<sup>15</sup>N<sub>nit</sub> = δ<sup>15</sup>N of NO<sub>3</sub><sup>-</sup> produced during complete nitrification of permafrost or atmospherically derived ammonium, average of 1.22‰ (range: -6 to 10‰; Arendt et al. 2016; Heikoop et al. 2016)

δ<sup>18</sup>O<sub>nit</sub> = δ<sup>18</sup>O of NO<sub>3</sub><sup>-</sup> produced through nitrification of ammonium, average of -14.21‰ (range: -19.51 to -8.9‰; assuming δ<sup>18</sup>O DO of 23.5-24.2‰ (Kiddon et al., 1993; Wang & Veizer, 2000; Horibe et al., 1973 as in Wynn et al., 2007) and δ<sup>18</sup>O H<sub>2</sub>O of -12 to -22‰ (Arendt et al. 2016, Wynn et al., 2007))

δ<sup>15</sup>N<sub>atm</sub> = δ<sup>15</sup>N of NO<sub>3</sub><sup>-</sup> from atmospheric deposition, -3.54‰ (-8.72-1.40; Ansari et al., 2013; Hastings et al., 2003; Heikoop et al., 2015; Louiseize et al., 2014)

δ<sup>18</sup>O<sub>atm</sub> = δ<sup>18</sup>O of NO<sub>3</sub><sup>-</sup> from atmospheric deposition, 72.07‰ (60.30-80.20; Ansari et al., 2013; Hastings et al., 2003; Heikoop et al., 2015; Louiseize et al., 2014)

ε<sub>as</sub><sup>15</sup> = isotope effect of NO<sub>3</sub><sup>-</sup> assimilation on N (5‰; Altabet, 2001)

$\epsilon_{as}^{18}$  = isotope effect of  $\text{NO}_3^-$  assimilation on O (5.9‰; Rafter & Sigman, 2015)

$\delta^{15}\text{N}_{re}$  =  $\delta^{15}\text{N}$  of  $\text{NO}_3^-$  from recycled production

$\delta^{18}\text{O}_{re}$  =  $\delta^{18}\text{O}$  of  $\text{NO}_3^-$  from recycled production average of -14.21‰ (range: -19.51 to -8.9‰; assuming  $\delta^{18}\text{O}$  DO of 23.5 to 24.2‰ (Kiddon et al., 1993; Wang & Veizer, 2000; Horibe et al., 1973 as in Wynn et al., 2007) and  $\delta^{18}\text{O}$   $\text{H}_2\text{O}$  of -12 to -22‰ (Arendt et al. 2016, Wynn et al., 2007))

$\delta^{15}\text{N}_{as}$  =  $\delta^{15}\text{N}_{box} - \epsilon_{as}^{15}$  =  $\delta^{15}\text{N}$  of  $\text{NO}_3^-$  assimilated

$\delta^{15}\text{N}_{box}$  = model output of  $\delta^{15}\text{N}$  of  $\text{NO}_3^-$  in the river

### Model Conditions

- Nitrified  $\text{NO}_3^-$  is a source
- Atmospheric  $\text{NO}_3^-$  is a source
- $\text{NO}_3^-$  assimilation is a sink
- Recycled production accounts for 50% or 25% of supplied  $\text{NO}_3^-$
- No denitrification

For  $\delta^{15}\text{N}$

$$\delta^{15}\text{N}_{box} = a \times \delta^{15}\text{N}_{nit} + b \times \delta^{15}\text{N}_{atm} - [(d - e) \times (\delta^{15}\text{N}_{box} - \epsilon_{as}^{15})]$$

This can be arranged to

$$\delta^{15}\text{N}_{box} = \frac{a \times \delta^{15}\text{N}_{nit} + b \times \delta^{15}\text{N}_{atm} + [(d - e) \times \epsilon_{as}^{15}]}{(1 + d - e)}$$

For  $\delta^{18}\text{O}$

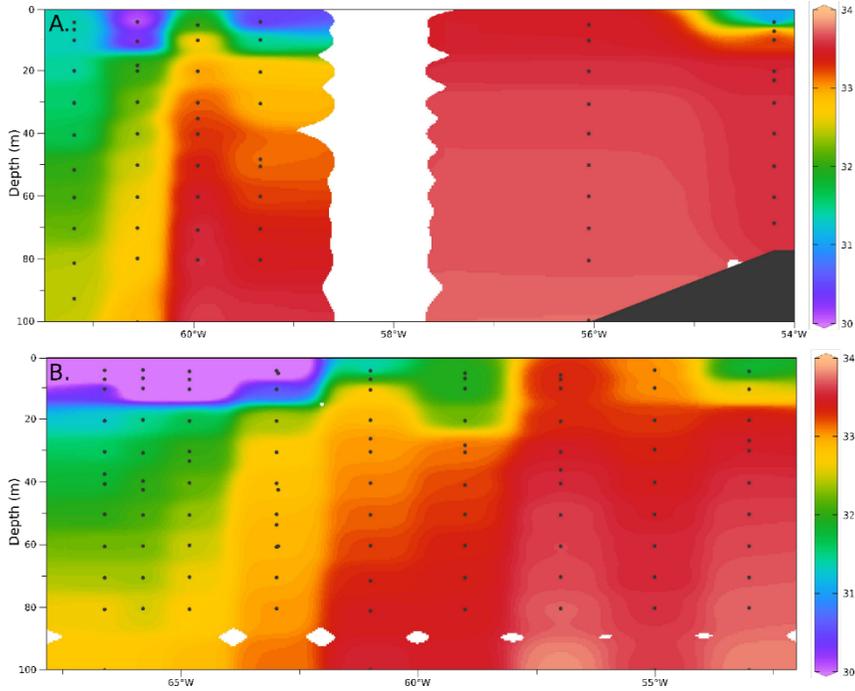
$$\delta^{18}\text{O}_{box} = a \times \delta^{18}\text{O}_{nit} + b \times \delta^{18}\text{O}_{atm} - [d \times (\delta^{18}\text{O}_{box} - \epsilon_{as}^{18})] + e \times \delta^{18}\text{O}_{re}$$

This can be arranged to

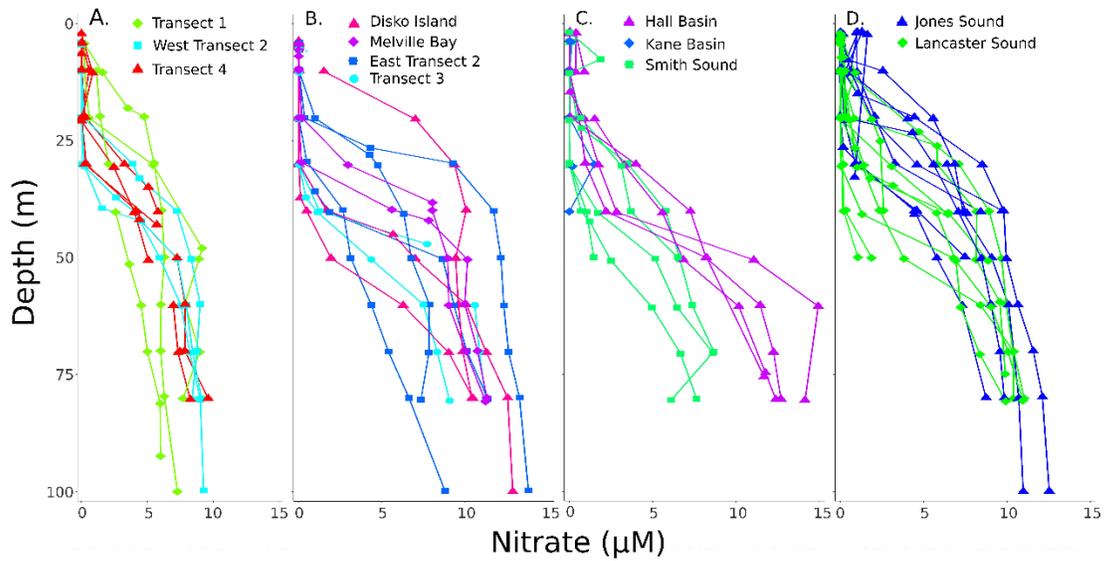
$$\delta^{18}\text{O}_{box} = \frac{a \times \delta^{18}\text{O}_{nit} + b \times \delta^{18}\text{O}_{atm} + d \times \epsilon_{as}^{18} + e \times \delta^{18}\text{O}_{re}}{(1 + d)}$$

The  $\Delta(15,18)$  is calculated according to Rafter and Sigman (2016):

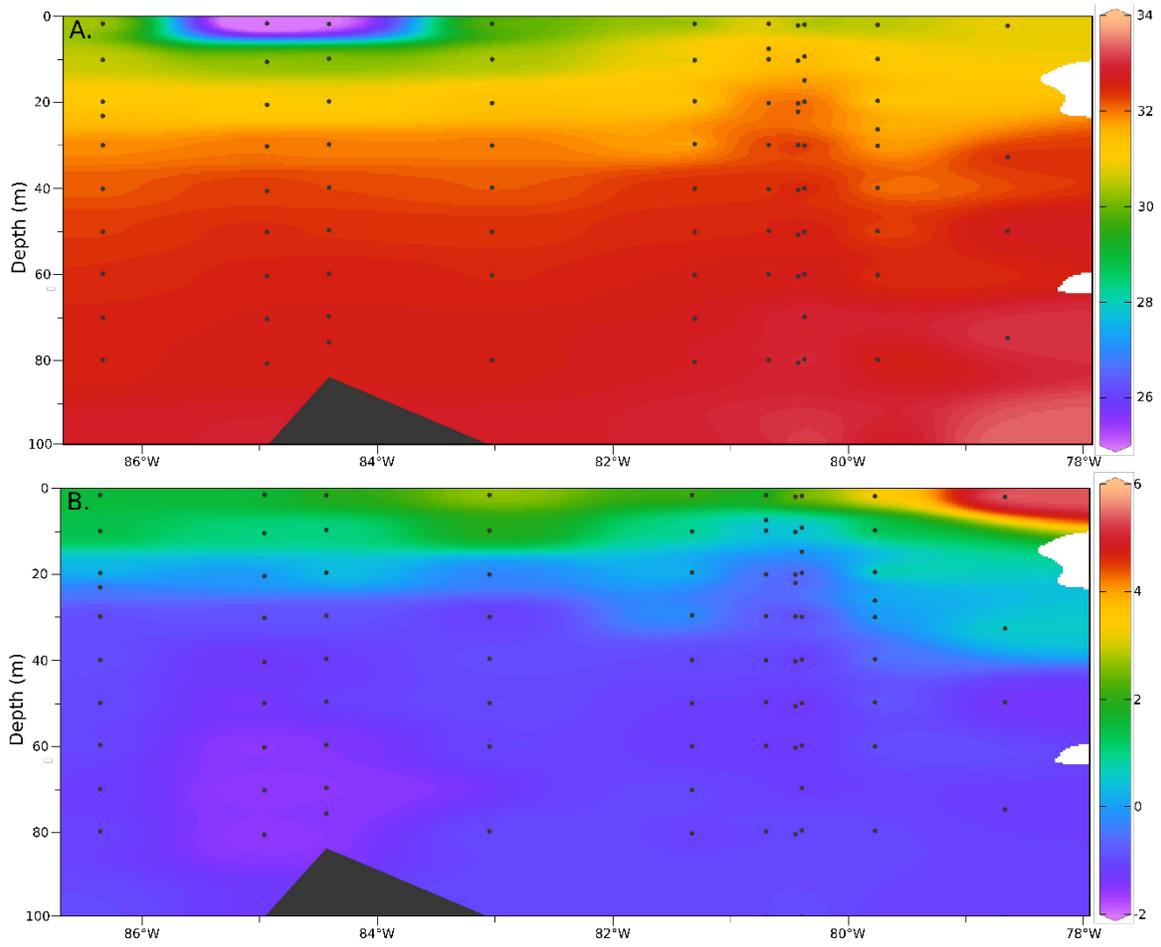
$$\Delta(15,18) = \delta^{15}\text{N} - \delta^{18}\text{O}$$



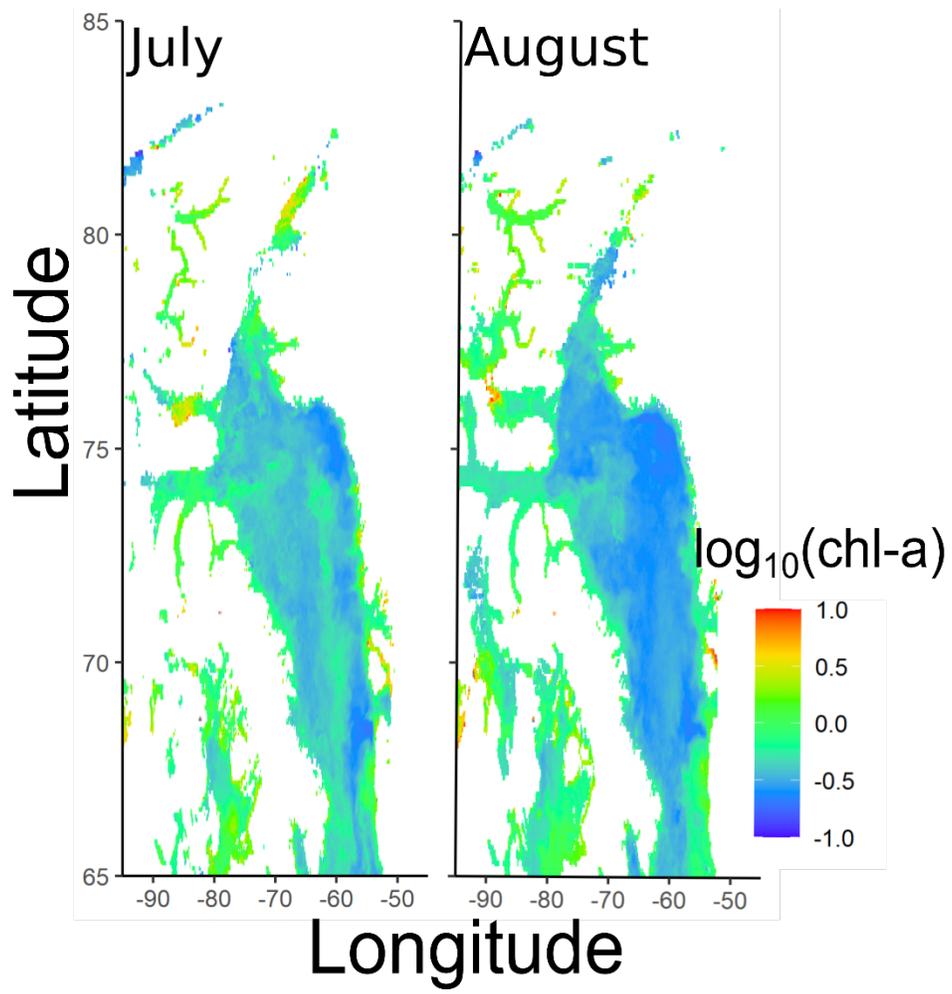
**Figure S1:** Salinity cross sections of A. Transect 1 and B. Transect 2



**Figure S2:** Nitrate depth profiles for A. Western Baffin Bay B. Eastern Baffin Bay C. Nares Strait and D. Jones and Lancaster Sounds



**Figure S3:** Jones Sound Salinity (A) and Temperature (B) Cross Section.



**Figure S4:** Chlorophyll-a in the study region via satellite observations