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2 **Status of Glaciers in the Western United States based on Sentinel-2A Images and Machine**
3 **Learning Algorithm**

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7 **Key Points:**

- 8 • This research evaluates the present condition of glaciers in the western United States.
9 • Over the study period, 930 glaciers have vanished entirely.
10 • There is a 35.43% reduction in glacier area in the western U.S.
11 • The volume of glaciers has decreased by 4.9 km³, roughly equivalent to 4.7 gigatons of
12 water.

13 **Abstract**

14 In this study, we employed random forest machine learning classification to assess the current
15 state of glaciers in the western United States using Sentinel-2A satellite imagery. By analyzing
16 Sentinel-2A imagery from September 2020 and comparing it to the RGI inventory, the study
17 determined the current conditions of the glaciers. Our findings unveiled a significant reduction in
18 both glacier area and volume in the western United States since the mid-20th century. Currently,
19 the region hosts 4091 glaciers spanning seven states, covering a total area of 432.01 km² with a
20 corresponding volume of 9.02 km³. During the study period, a loss of 237.07 km² in glacier area
21 was observed, representing a 35.43% decrease when contrasted with the RGI boundaries. The
22 volume lost during this period amounted to 4.9 km³, roughly equivalent to 4.7 gigatons of water.
23 Among the states, Washington experienced the most significant glacier area reduction, with a loss
24 of 130.06 km². Notably, glaciers in the North Cascade Range of Washington, such as those in Mt.
25 Baker and Mt. Shuksan, now cover, on average, only 85% of their original glacier boundaries
26 with ice and snow at the conclusion of the 2020 hydrological year. Major glaciers, including the
27 White River glacier, West Nooksack glacier, and White Chuck glacier, have lost more than 50
28 percent of their original area.

29 **Plain Language Summary**

30 We used advanced computer programs to study glaciers in the western United States. We
31 analyzed Sentinel-2A satellite images collected in September 2020. By comparing these images
32 with a database of glaciers, we found present conditions of glaciers. Our research found that the
33 glaciers in the western United States have been shrinking since the middle of the 1900s.
34 Currently, there are 4091 glaciers in seven states, covering a total area of 432.01 square
35 kilometers and holding 9.02 cubic kilometers of ice. During our study time, we saw that the
36 glaciers lost 237.07 square kilometers of ice, which is a 35.43% decrease compared to before.
37 The amount of ice that melted during this time is about 4.9 cubic kilometers, roughly the same as
38 4.7 gigatons of water. Among the states, Washington saw the biggest reduction in glacier size,
39 losing 130.06 square kilometers. Specifically, glaciers in the North Cascade Range of
40 Washington, like those in Mt. Baker and Mt. Shuksan, now only cover 85% of their original
41 boundaries with ice and snow by the end of the 2020 hydrological year. Some major glaciers,
42 including the White River glacier, West Nooksack glacier, and White Chuck glacier, have lost
43 more than 50% of their original size.

44 **1 Introduction**

45 Glaciers in the Western United States have shrunk and lost mass in the last century
46 (McCabe & Fountain, 2013; Inamdar and Ambinakudige 2016; Riedel et al. 2015). An inventory
47 of the current status of glaciers in the US is essential to monitor rapid changes in glacier area and
48 volume due to climate change. The GLIMS (Global Land Ice Measurements from Space) and
49 RGI (Randolph Glacier Inventory) database provides the boundaries of glaciers in the world
50 including the U.S. The boundaries of the U.S. glaciers in GLIMS and RGI database were
51 developed based on old Landsat satellite images and topographic maps. In this study, we present
52 the status of glaciers in the western United States (excluding Alaska) as of 2020. This study
53 assesses glacier parameters (area and volume) and their changes within the study period using
54 Sentinel-SA images and a Random Forest algorithm.

55 The cryosphere is composed of snow and ice in the form of snow cover, sea ice,
56 freshwater ice, permafrost, and continental ice masses such as glaciers and ice sheets
57 (Ambinakudige & Joshi, 2012). Changes in global climate have had a significant impact on the
58 world's glaciers, causing them to shrink rapidly in size and mass, subsequently raising global sea
59 levels, altering hydrology, and increasing natural hazards (Ambinakudige & Intsiful, 2022;
60 Berthier et al., 2010; Dixon & Ambinakudige, 2015; Hugonnet et al., 2021; Huss & Hock, 2018;
61 Maloof et al., 2014; Rounce et al., 2023; Wouters et al., 2019; Zemp et al., 2019; Fountain et al.,
62 2017; Inamdar & Ambinakudige, 2016; Singh et al., 2018). Glacier retreat and volume loss are
63 directly linked to a warming climate (Intsiful & Ambinakudige, 2021). According to the IPCC
64 (2023), global surface temperatures have increased by 1.09°C due to anthropogenic activities.
65 The IPCC (2023) report also predicts that many low-elevation and small glaciers would lose most
66 of their mass or disappear within decades to centuries. Between 1901 and 2018, the global mean
67 sea level increased by 0.20 m (IPCC, 2023). Glacier mass loss of all glaciers outside Antarctica
68 and Greenland glacier accelerated significantly between 2000 and 2019 with an average loss of
69 267 ± 16 gigatons per year, equivalent to $21 \pm 3\%$ of the observed sea-level rise (Hugonnet et al.,
70 2021). In the mid-19th century, glaciers have retreated substantially due to increased temperature
71 (Moore et al., 2009) and this has resulted in major variations in streamflow throughout the globe.
72 Continuous warming is causing glaciers to melt rapidly, resulting in the loss of small glaciers,
73 and faster shrinking of large glaciers (Moore et al., 2009). These changes in cryosphere
74 subsequently led to raising global sea levels, altering hydrology, and increasing natural hazards
75 (Berthier et al., 2010; Maloof et al., 2014; Rounce et al., 2023).

76 All parts of the world have witnessed significant decrease in glacier area and volume
77 (Hugonnet et al., 2021; Huss & Hock, 2018; Rounce et al., 2023; Wouters et al., 2019; Zemp et
78 al., 2015, 2019). Hugonnet et al., (2021) assessed the global mass loss of glaciers during the early
79 21st century using satellite images and airborne elevation datasets as from the year 2000 to 2019
80 of all glaciers outside Antarctica and Greenland and reported an estimated glacier mass loss has
81 accelerated significantly during this period with average loss of 267 ± 16 gigatons per year,
82 equivalent to $21 \pm 3\%$ of the observed sea-level rise.

83 In the western United States, glaciers continue to lose ice, which is consistent with
84 broader trends of glacier retreat and shrinkage worldwide. Using historic photographs and maps,
85 McCabe & Fountain, (2013) provided an assessment of glaciers in the conterminous United
86 States during the twentieth century. According to this study, 24% of the glacier area has been lost
87 in Mt. Rainer located in Washington State. In addition, McCabe & Fountain (2013) also reported
88 an extensive recession of 66% in the Lewis Range in Montana (located in the Glacier National
89 Park) and the Sierra Nevada in California. Riedel et al., (2015) investigated the status of glaciers
90 in the Olympic Mountains and their contribution to streamflow in the Hoh River basin using a
91 hydrological model and they reported that all glaciers in this area have experienced significant ice
92 loss from 1986 to 2011, with the Anderson and Blue glaciers losing more than 50% of their ice
93 volume. The study also indicated that glacier shrinkage in the past 30 years within the Olympics
94 is greater than in the Cascades and southern Coast Mountains.

95 In another study, Fountain et al., (2022) examined how glaciers in the Olympic Mountains
96 have changed in area and volume from early 1980 to 2015. They discovered that glaciers
97 decreased at a rate of 0.59 km² per year, resulting in the disappearance of 35 glaciers and 16
98 perpetual snowfields. They also noted that warming winter temperatures are particularly

99 contributing to glacier shrinkage. In the Sierra Nevada in California, Basagic & Fountain, (2011)
100 conducted a comprehensive study of glaciers using a combination of historical aerial
101 photographs, satellite imagery, and field measurements to assess glacier area and volume changes
102 over time. This study revealed that the glaciers in the Sierra Nevada have decreased in area by
103 about 39.86 ± 9.59 km² (56%) in 2004 (a loss since 1903). Declining snowfall and rising
104 temperatures are the main causes of glacier retreat in the Sierra Nevada (Basagic & Fountain,
105 2011).

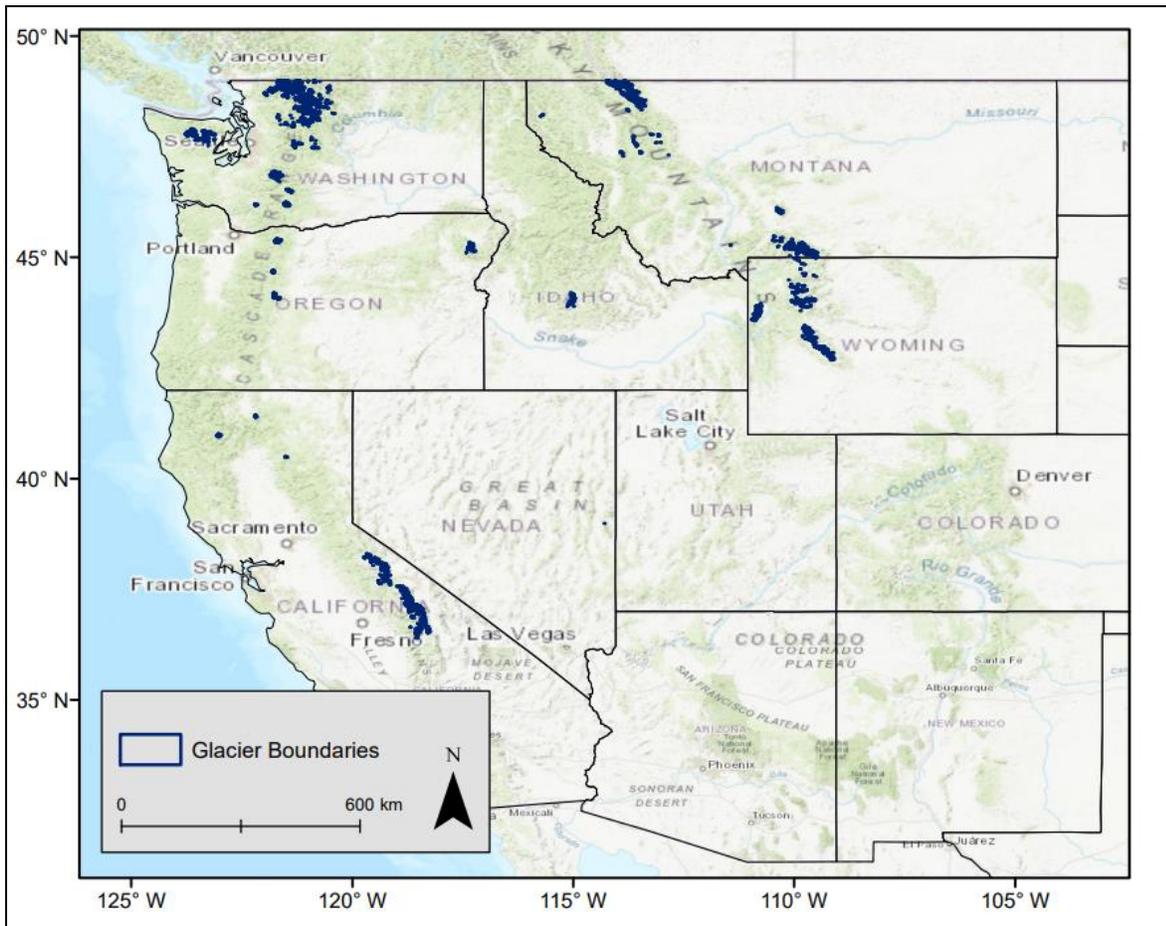
106 Melting of glaciers consistent with the global warming trend observed in the area. O’Neel
107 et al., (2019) conducted a study on the United States Geological Survey Benchmark Glaciers
108 (Gulkana, Wolverine, Lemon Creek, South Cascade and Sperry glaciers) to understand the long-
109 term insight into climate forcing of glacier mass balance from 1953 to 2015. To do this, they
110 reanalyzed all available data from the United States Geological Survey Benchmark Glaciers by
111 combining all available data from each glacier into a single dataset. The results of the reanalysis
112 showed that all five glaciers have experienced negative mass balances since the mid-20th
113 century, with average rates ranging from -0.58 to -0.30 m w.e. a⁻¹.

114 Glacier inventories have been very valuable in assessing glacier changes and their
115 contribution to the rise of sea level (Gardner et al., 2013; Pfeffer et al., 2014). The first complete
116 inventory of glaciers to assess the status, magnitude, and rate of change of glaciers area in the
117 western United States in the mid-20th century was assembled by (Fountain et al., 2007; Fountain
118 et al., 2017) which is part of the GLIMS and RGI inventory of the United States. They utilized
119 1:100,000 (100K) and 1:24,000 (24K)-scale topographic maps published by the U.S. Geological
120 Survey (USGS) and U.S. Forest Service (USFS) based on aerial photos to provide the most
121 comprehensive glacier inventory in the western United States. Since the completion of the first
122 glacier inventory of the United States, no comprehensive assessment of the glaciers in the
123 Western United States, particularly the Lower 48 conterminous States, has been carried out.
124 However, there have been studies that have provided information on the status of individual
125 glaciers as well as the status of glaciers on a regional or state level (Basagic & Fountain, 2011;
126 Granshaw & Fountain, 2006; Martin-Mikle & Fagre, 2019). Here, we present the current status of
127 glaciers in the western United States (excluding Alaska) spanning the mid-20th century and 2020.
128 This study assesses glacier parameters (area and volume) and their changes within the study
129 period using Sentinel-SA images and a Random Forest algorithm.

130 **2 Study Area**

131 The Randolph Glacier Inventory (RGI, 2017) identifies 5021 glaciers in the western
132 conterminous United States which are found in California, Washington, Montana, Idaho,
133 Wyoming, Nevada, and Oregon (Figure.1). It includes significant mountain ranges such as the
134 Rocky Mountains spanning Wyoming, Montana, and Idaho; the Cascade Range extending
135 through Washington, Oregon, and California; the Sierra Nevada located in California; and the
136 Olympic Mountains situated in Washington. Most of these glaciers are found in Washington
137 state, specifically in the north Cascade Range. The RGI inventory indicates that the Cascades
138 region hosted ~ 454 km² of glaciers. Among these, around 408km² of the total glacier area within
139 the Cascades was observed in Washington state. Roughly 41km² of these glaciers were located in
140 cascade of Oregon, while the remaining 5 km² was distributed in the northern part of California.
141 Majority of large glaciers (>2 km²) are located in the Cascade and the Olympic Mountains which

142 make up a significant portion of total area of glaciers in the western U.S. The Sierra Nevada of
 143 California and the South of the Cascade Range is composed of small glaciers and snow patches
 144 with a total area of $\sim 42 \text{ km}^2$. In the Rocky Mountains of Wyoming and Montana, the RGI
 145 inventory maps 139 km^2 of glaciers. Most of the glaciers are found in the Wind River Range, the
 146 Lewis Range and the Beartooth Mountains. Several of the glaciers are $>1 \text{ km}^2$, including Gannett
 147 Glacier, Grasshopper Glacier and the Sacagawea Glacier, the largest glaciers in the continental
 148 U.S. outside of Washington.



149
 150 Figure. 1. Distribution of Glaciers in the western United States defined by RGI inventory (Base
 151 map Source: ESRI Topographic map)

152 Typically, the Cascades and Sierra Nevada areas exhibit maritime snow climates, which
 153 involve considerable precipitation and temperature around 0°C for a significant portion of the
 154 winter. However, the Rocky Mountains tend to showcase more continental conditions. This is
 155 evident through comparatively reduced precipitation levels and winter temperatures considerably
 156 below 0°C at elevated altitudes (Selkowitz & Forster, 2016).

157 **3 Data and Methods**

158 3.1 Satellite Data

159 Sentinel-2A satellite images for the September month of 2020 were used to identify the
 160 current conditions of all the glaciers covering the glacierized areas of the western United States.
 161 Images were obtained from the Sentinel Open Hub (<https://scihub.copernicus.eu/>). Sentinel-2A
 162 Multispectral Instrument (MSI), launched in 2015, is a multispectral sensor with 13 bands
 163 covering the visible, near infrared (VNIR) and short-wave infrared (SWIR) wavelength regions.
 164 These bands comprise spatial resolutions of 10m, 20m, and 60m. The 10m resolution contains
 165 four bands: blue, green, red, and near-infrared-1. The 20m resolution contains six bands: red edge
 166 1-3, near-infrared-2, short-wave infrared 1 and 2. The 60m resolution contains: Band 1, Band 9,
 167 and Band 10 that are 60m resolution. The three 60m resolution bands specifically dedicated for
 168 atmospheric correction and cloud screening (Drusch et al., 2012). Table 1 summarizes Sentinel-
 169 2A satellite data by band characteristics, wavelength, and spatial resolution.

170 Table 1: Sentinel-2A Bands

Sentinel Bands	Characteristic	Central Wavelength(μm)	Resolution (m)
Band 1	Coastal Aerosol	0.443	60
Band 2	Blue	0.490	10
Band 3	Green	0.560	10
Band 4	Red	0.665	10
Band 5	Near Infrared	0.705	20
Band 6	Near Infrared	0.740	20
Band 7	Near Infrared	0.783	20
Band 8	Near Infrared	0.842	10
Band 8A	Near Infrared	0.865	20
Band 9	Water Vapor	0.945	60
Band 10	Cirrus	1.375	60
Band 11	Shortwave Infrared	1.610	20
Band 12	Shortwave Infrared	2.190	20

171

172 3.2 Reference Data

173 The RGI 6.0 glacier outline was used as reference data. Google Earth images and
 174 topographic maps downloaded from the National Geologic Map Database developed by the
 175 USGS were used as secondary reference data to aid in manual digitization of glacier boundaries
 176 and verification of the glacier names.

177 3.3 Methods

178 Normalized Difference Vegetation Index (NDVI), Normalized Difference Glacier Index
 179 (NDGI), and Normalized Difference Snow/Ice Index (NDSII) were calculated from Sentinel
 180 bands (Table 2). NDVI is a normalized ratio of the near-infrared and red bands used to quantify

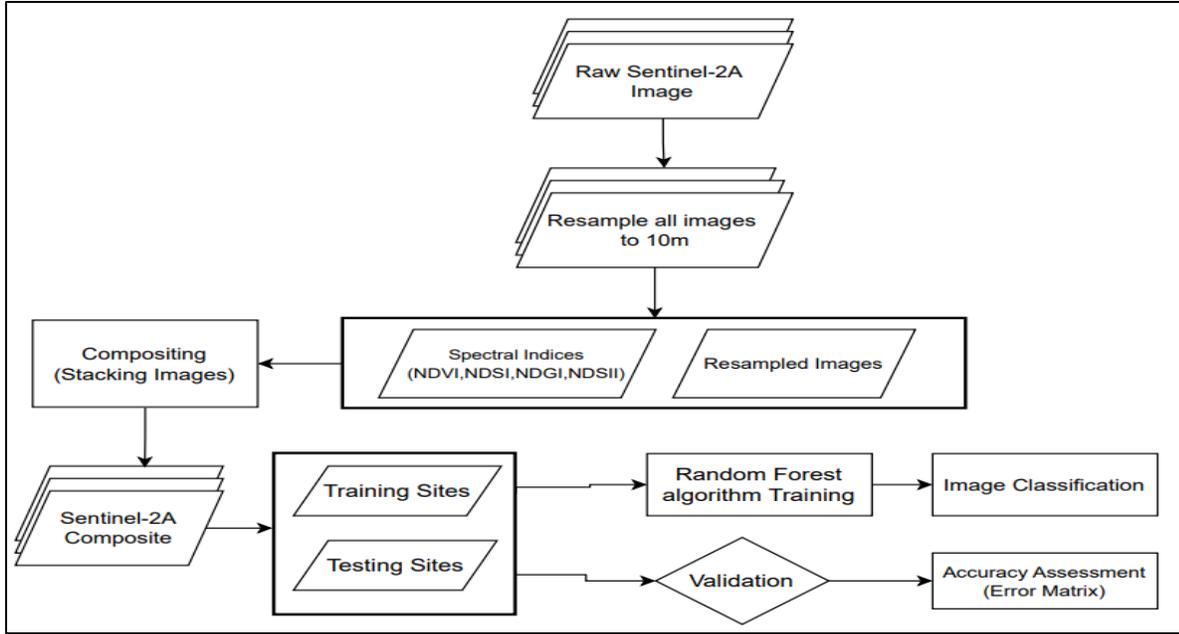
181 vegetation cover and vegetation health (Huete et al., 2002). NDSI takes advantage of the
 182 reflectance of the visible and short-wave infrared regions of the electromagnetic spectrum. NDGI
 183 uses the ratio between the red and green regions of the EMS to detect and monitor glaciers.
 184 NDSII detects snow and ice by band ratioing the green band and the near infrared. The fact that
 185 ice reflectance declines towards the near infrared (NIR) indicates that this spectral region has
 186 potential for distinguishing between snow and ice (Keshri et al., 2009). All these band ratios also
 187 reduce noise such as illumination differences, cloud shadows, and atmospheric attenuation.
 188 Composite bands for all the images in the study area were created using all 13 bands and the four
 189 spectral indices to help in classification of the land cover classes.

190 Table 2: Spectral indices used in the study.

Index	Name	Formula
NDSI	Normalized difference snow index	$\frac{B11 - B3}{B11 + B3}$
NDVI	Normalized Difference Vegetation Index	$\frac{B8A - B4}{B8A + B4}$
NDGI	Normalized Difference Glacier Index	$\frac{B3 - B4}{B3 + B4}$
NDSII	Normalized difference snow/Ice index	$\frac{B3 - B8A}{B3 + B8A}$

191

192 A flowchart of the classification process is provided in Figure.2. All bands were
 193 resampled to 10 meters using the nearest neighbor method. A machine learning supervised image
 194 classification called Random Forest (RF) algorithm is used to classify the Sentinel-2A satellite
 195 images. Random forest is a type of non-parametric ensemble model where multiple decision trees
 196 are constructed and the final output is determined based on the combined results of these trees
 197 (Horning, 2010). Random forest aggregates multiple decision trees' predictions to obtain a more
 198 accurate and stable result. The trees produce a class prediction for a pixel, and the model predicts
 199 the pixel as the class with the most votes. In this study, classification of the images was done
 200 based on RF machine learning classification (Ambinakudige & Intsiful, 2022). During the
 201 classification, over 26,500 pixels were selected in each image as training site pixels and each
 202 pixel was classified as water, barren, snow, ice, vegetation, shadow, or debris. Shadowed areas
 203 were manually classified separately for ice or snow after the main classification, so it was
 204 imperative to separate the glacierized areas (Including shadowed areas with ice and snow) from
 205 the non-glacierized areas. While 70 percent of the pixels were used for the training, 30 percent of
 206 pixels were used for testing. The glacial boundary shapefiles from the RGI 6.0 were used as
 207 references. RGI glacier boundaries were overlaid on different color composite bands, to ensure
 208 accurate identification of the new glacier boundaries. Classified images, topographic maps and
 209 Google Earth images also served as references for detailed manual digitization. Digitization of
 210 the new glacier boundaries was done in ArcGIS 10.8.2. As part of the digitization process, a
 211 separate polygon feature class was generated for glaciers in the study area, and these polygons
 212 served as the new boundaries of the glaciers. The difference between the areas of the RGI
 213 boundaries and the newly created boundary for each glacier is computed to assess how much area
 214 has been lost or gained within the study period.



215
216 Figure. 2. Schematic diagram of the classification process

217 Glacier volume change is an important glacier parameter because it can be applied to
218 estimate glacier mass change and to estimate meltwater runoff (Granshaw & Fountain, 2006).
219 Area measurements were used to estimate glacier volume using a simple area-volume scaling
220 relationship (Ambinakudige & Intsiful, 2022; Chen & Ohmura, 1990). The volume of an alpine
221 glacier is related to its surface area as follows:

222 $V = cA^e$ 1

223 where V is the glacier volume (km^3), A is glacier area (km^2) and c and e are coefficients and can
224 be empirically or theoretically derived (Basagic & Fountain, 2011; Martin-Mikle & Fagre, 2019).
225 There are several coefficients available for different study areas. Chen & Ohmura, (1990) derived
226 nine coefficients for different study areas with four applicable to alpine glaciers in the western
227 United States. They include “Cascade, small glaciers” ($c = 0.021346$, $e = 1.145$), “Cascade &
228 other areas” ($c = 0.03834$, $e = 1.405$), “Alps, Cascade and other areas” ($c = 0.027551$, $e = 1.358$),
229 and “Alps, Cascade and Svalbard etc.” ($c = 0.028524$, $e = 1.357$). Fountain et al., (2017) reported
230 that the volume of glaciers in the western United States is better estimated using the empirical
231 “Cascade, small glaciers” approach where $c = 0.021346$ and $e = 1.145$. Total volume change is
232 then estimated by calculating the difference between the total volume of the RGI boundary and
233 the total volume of the new glacier boundaries.

234 3.4 Accuracy Assessment

235 Assessment of the classifier's accuracy of Random Forest classification was done using
236 the overall accuracy, and Kappa coefficient from the confusion matrix (error matrix). The
237 confusion matrix summarizes the classifier's performance by comparing the classified data with
238 reference data (Banko, 1998).

239 Overall accuracy is the proportion of correctly classified pixels to the total pixels. It is computed
240 as:

$$241 \quad \textbf{Overall Accuracy} = \frac{\textit{Number correctly classified pixels}}{\textit{Total Number of pixels}} \quad \mathbf{2}$$

242 Kappa Coefficient provides a measure of the classifier's performance that is less affected by
243 chance agreement (Banko, 1998). It is computed as:

$$244 \quad \mathbf{K} = \frac{N \sum_{i=1}^r x_{ii} - \sum_{i=1}^r (x_{i+} \cdot x_{+i})}{N^2 - \sum_{i=1}^r (x_{i+} \cdot x_{+i})} \quad \mathbf{3}$$

245 Where, r = number of rows and columns in the confusion matrix

246 x_{ii} = total number of observations in row i column i

247 x_{i+} = Total number of observations in row i

248 x_{+i} = Total number of observations in column i

249 N = Total number of observations in the matrix

250 3.5 Error Estimation

251 This study used satellite imagery to extract glacier parameters (area and volume), hence it
252 is subject to uncertainties arising from image quality, sensor characteristics, interpretation of
253 glacial features, and post-processing techniques (Paul et al., 2017). To estimate glacier area
254 uncertainty, we employed the buffer method (Bolch et al., 2010; Granshaw & Fountain, 2006).
255 The buffer method which expands and shrinks the outline of each glacier by an uncertainty value
256 provides minimum and maximum estimates of uncertainty for each glacier that can be converted
257 to a standard deviation (STD) when a normal distribution is assumed for the differences. The
258 standard deviation is then used as a component of the precision of the outline, and it is
259 recommended when there is no reliable reference data available (Paul et al., 2017). A buffer of
260 10m, equivalent to one pixel size was created around each glacier. This resulted in an average
261 mapping uncertainty for a single glacier of 0.032 km² and total uncertainty in area change
262 measurement of ± 5.3 km².

263 4 Results

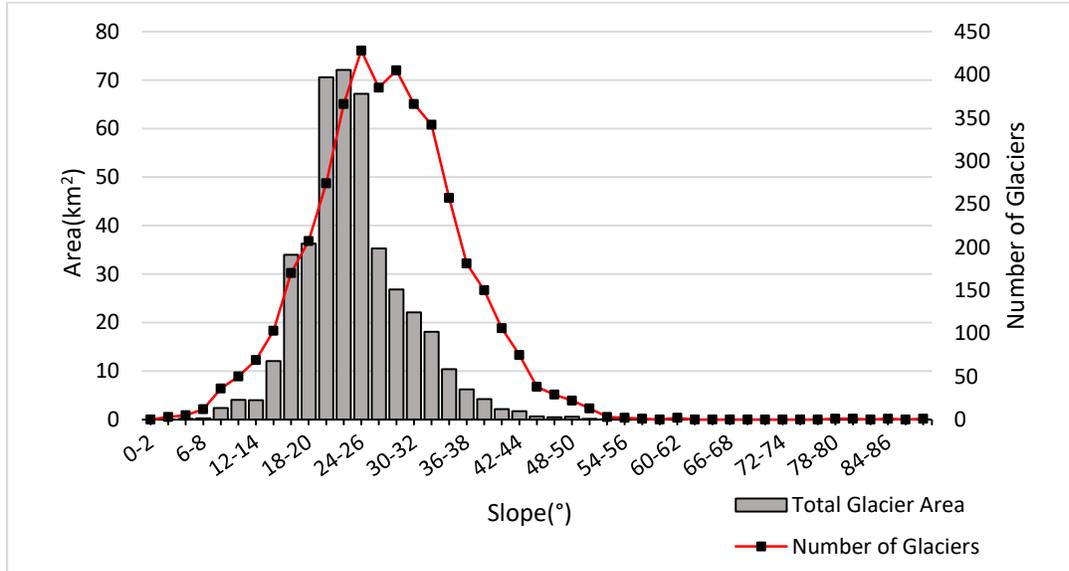
264 The Random Forest algorithm applied to classify Sentinel images produced a high level
265 of accuracy, as evidenced by the performance measures obtained from the analysis. Random
266 forest machine learning image classification method provided over 98 percent classification
267 accuracy with a kappa coefficient of 0.989.

268 4.1 Glacier Distribution and Characteristics

269 The RGI inventory had 5021 with a total area of 669.08 km² in the study area. The
270 inventory created in this study found only 4091 glaciers covering an area of 432.01 km². This
271 study identified 1790 glaciers in Washington, 776 in Wyoming, 733 in Montana, 604 in
272 California, 153 in Oregon, 34 in Idaho, and 1 in Nevada (Table 3). That is 930 less glaciers, and
273 about 18.52% decrease in number of glaciers.

274 The study region is dominated by small glaciers (size <1km²). Out of 4091 glaciers, 4005
275 are smaller than 1 km², which is ~98% of the total number of glaciers. These form 49.2% of the

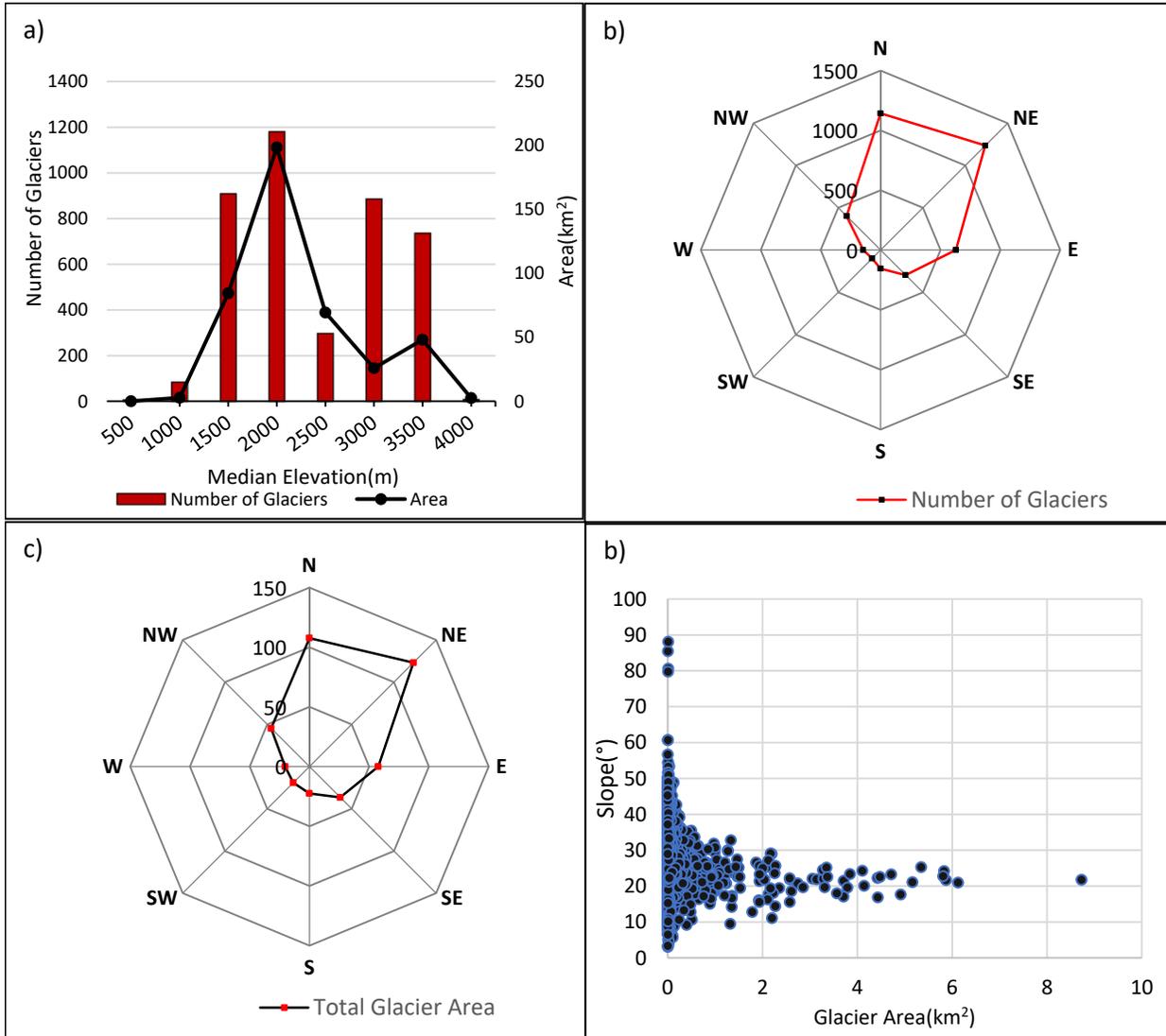
276 total area in this study, covering 212.68 km². Conversely, the remaining 86 glaciers which are
 277 larger than 1 km², make up ~ 51% (219.4 km²). There is a strong unevenness in the number of
 278 glaciers towards smaller glaciers (<0.05 km²). Glaciers less than 0.05 km² account for 77% of the
 279 number of glaciers in the study area, but they represent only 11% of the total area in this study.



280

281 Figure. 3. Slope - Area/Number distribution of all glaciers

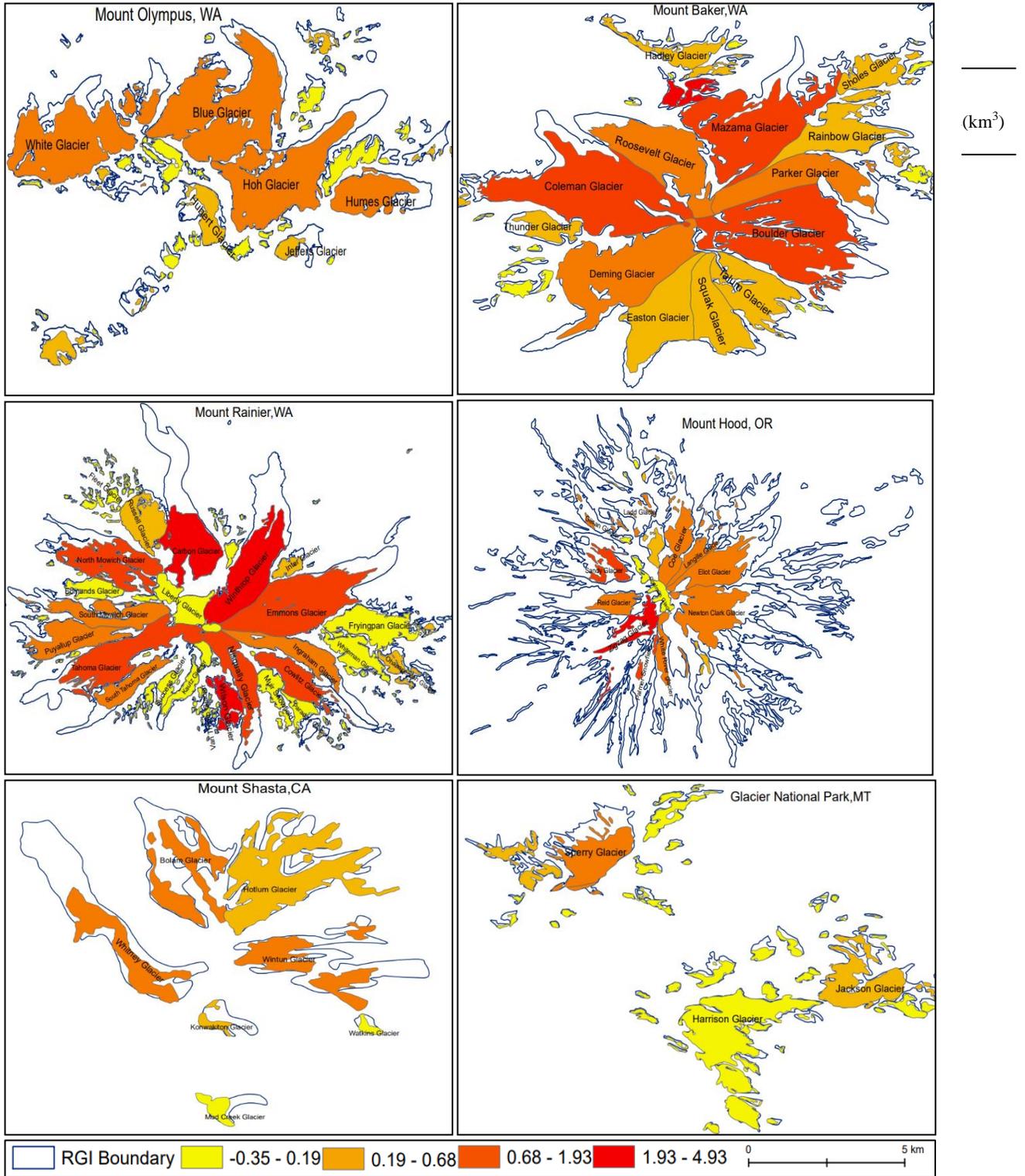
282 Generally, most of the glaciers in this study have a surface slope between 22°-32°.
 283 Glaciers with slopes of 20°-26° make up ~49% (209.85 km²) of the total glacierized area.
 284 (Figure.3). Figure 4a shows the elevation-area/number distribution with 1000m elevation
 285 intervals in the seven glacierized states. The distribution shows that the majority (352.5 km²,
 286 81.8%) of the total area in the study region is found between 1000m a.s.l. and 2500m a.s.l. The
 287 lowest and highest elevations are 636 and 4340m a.s.l. both in the Cascade Range of Washington
 288 while only 0.66% and 17.8% are distributed below and above 1000m a.s.l. and 2500m a.s.l.
 289 respectively. The north (n = 1143) and northeast (n = 1236) are the two glacier aspects that occur
 290 most frequently (Figure. 4b), as well as having the largest surface areas, 107.7 km² and 123.1 km²
 291 respectively (Figure. 4c). Also, large glaciers tend to have gentle slopes (Figure. 4d).



292 Figure. 4. Characteristics of glacier aspect, elevation, and slope. (a) Elevation – Area/Number
 293 distribution in 1000m bins. (b) Distribution of number of glaciers at different orientation (c) Total
 294 glacierized area at different orientation. (d). Slope-Area distribution of all glaciers.

295 4.2 Glacier Changes

296 The area and volume of glaciers and perennial snow fields in this study decreased
 297 significantly (Table 3, Figure. 5). The total area of glaciers in the RGI inventory is 669.08 km²
 298 with an overall volume of 13.92 km³. In our study, there are only 4091 glaciers that were found
 299 with a total area of 432.01 km² and volume of 9.02 km³. Hence, the total area loss during the
 300 study period is 237.07 km² (Table 3) which represents 35.43% of the area lost compared to the
 301 RGI glacier boundaries. The volume lost during this period was 4.9 km³.



302 Figure. 5. Glacier area changes of major glaciers in the western United States

California	-0.019	604	42.04	14.891	-27.15	-64.58	0.224
Cascade Range	-0.037	10	5.04	2.991	-2.05	-40.67	0.060
Sierra Nevada	-0.018	591	35.12	11.86	-23.26	-66.23	0.163
Trinity Alps	-0.050	3	1.88	0.039	-1.84	-97.87	0.001
Idaho	-0.023	34	2.02	0.195	-1.83	-90.59	0.002
Sawtooth Range	-0.023	34	2.02	0.195	-1.83	-90.59	0.002
Montana	-0.026	733	67.81	40.716	-27.09	-39.95	0.678
Beartooth-Absaroka	-0.034	254	21.85	10.235	-11.62	-53.18	0.155
Cabinet Mountains	-0.150	3	0.71	0.081	-0.63	-88.73	0.001
Crazy Mountains	-0.021	40	1.89	0.839	-1.05	-55.56	0.011
Lewis Range	-0.026	389	39.83	27.54	-12.29	-30.86	0.481
Madison Range	-0.018	1	0.04	0.003	-0.04	-100	0.000
Mission-Swan-Flathead	-0.009	46	3.49	2.018	-1.47	-42.12	0.03
Nevada	—	1	0.10	0.023	-0.08	-80	0
Snake Range	—	1	0.10	0.023	-0.08	-80	0
Oregon	-0.043	153	41.62	13.856	-27.76	-66.7	0.263
Cascade Range	-0.039	116	40.52	13.146	-27.37	-67.55	0.254
Wallowa Mountains	-0.006	37	1.1	0.71	-0.39	-35.45	0.009
Washington State	-0.051	1790	444.79	314.734	-130.06	-29.24	6.985
Cascade Range- North	-0.050	1306	285.41	201.786	-83.62	-29.3	4.291
Cascade Range- South	-0.038	244	122.81	88.682	-34.13	-27.79	2.168
Olympic Mountain	-0.040	240	36.57	24.266	-12.3	-33.63	0.526
Wyoming	-0.019	776	70.70	47.59	-23.12	-32.7	0.870
Absaroka Range	-0.011	216	8.38	5.589	-2.79	-33.29	0.077
Teton Range	-0.010	134	6.39	3.854	-2.54	-39.75	0.053
Wind River Range	-0.025	426	55.93	38.147	-17.78	-31.79	0.741
Total	-0.028	4091	669.08	432.01	-237.07	-35.43	9.023

303 Table 3. Glacier Area Changes across western United States including major mountain ranges

304 Individual glaciers lost area from 0 km² to -3.57 km². Small glaciers with area <0.5 km² lost
305 109.25 km² (41%) while glaciers larger than >0.5 km² lost 127.79 km² (31%). However, very
306 small glaciers (<0.01 km²) gained 6.05 km² of area (Figure. 6, Table 4).

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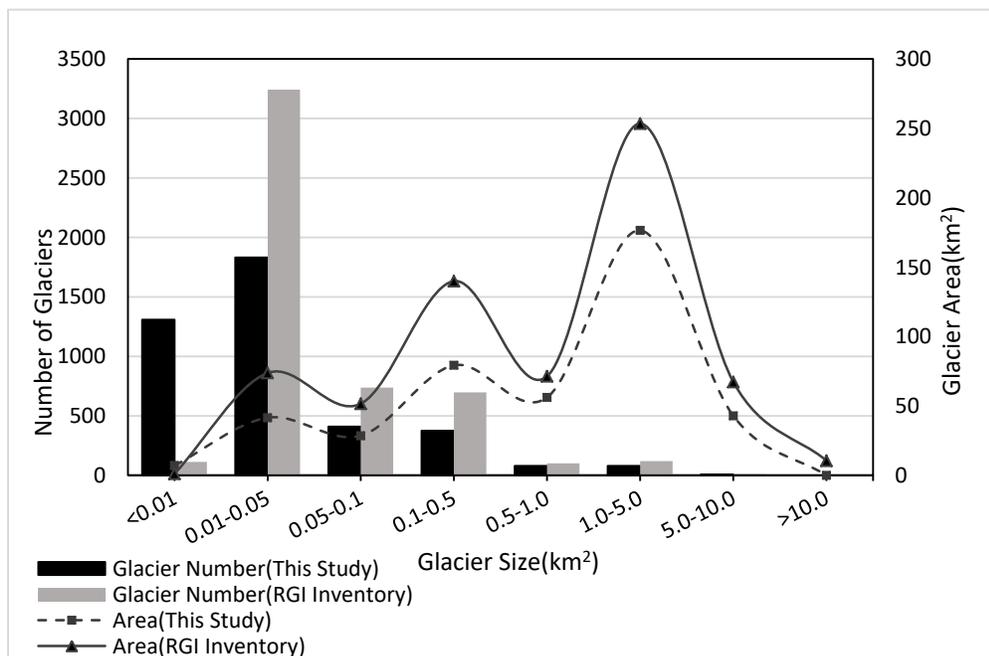
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319 Table 3. Glacier area changes across different size classes

Size	Count	Area	Area	Change	%
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interval(km ²)	(This study)	(This Study) km ²	Count (RGI)	(RGI) km ²	(km ²)	Change
<0.01	1308	7.15	113	1.10	6.05	548.04
0.01-0.05	1832	41.53	3242	73.78	-32.25	-43.45
0.05-0.1	410	28.52	738	51.63	-23.11	-44.62
0.1-0.5	375	79.30	697	139.92	-60.62	-43.02
0.5-1.0	80	56.14	101	71.36	-15.22	-22.23
1.0-5.0	79	176.50	119	253.42	-76.92	-30.35
5.0-10.0	7	42.87	10	67.29	-24.42	-36.29
>10.0	0	0	1	10.58	-10.58	-100
Total	4091	432.01	5021	669.08	-237.07	-35.43

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322 Figure. 6. Glacier distributions and changes according to area-size classes

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324 Among the states that contains glaciers, Washington state is the most glacierized state
 325 with an ice coverage of 314.73 km², which accounts for ~73% of the total glacierized regions in
 326 the western United States. Glacier area reduction is also most prominent in Washington State
 327 (Table 3). This study identified 1790 glaciers in Washington, with a mean area of 0.176 km².
 328 Washington State glacier coverage has decreased by 29.24% of the RGI value, from 444.79 km²
 329 to 314.73 km². This region has the highest mean area of the seven states (Table 3).
 330 Approximately 96% of the glaciers in Washington state are less than 1 km² with their total area
 331 contributing to ~39% (123.3 km²) of the total area in the region. Only 69(~4%) glaciers had an
 332 area between 1-10 km²; they had a combined area of 191.4 km², representing ~61% of the total
 333 area in Washington state. The total number of glaciers has declined from 1968 in the RGI
 334 inventory to 1790 in this current study. Glacier area reduction from Washington state alone
 335 represent ~55% of glaciers lost in the entire western United States during the study period. The

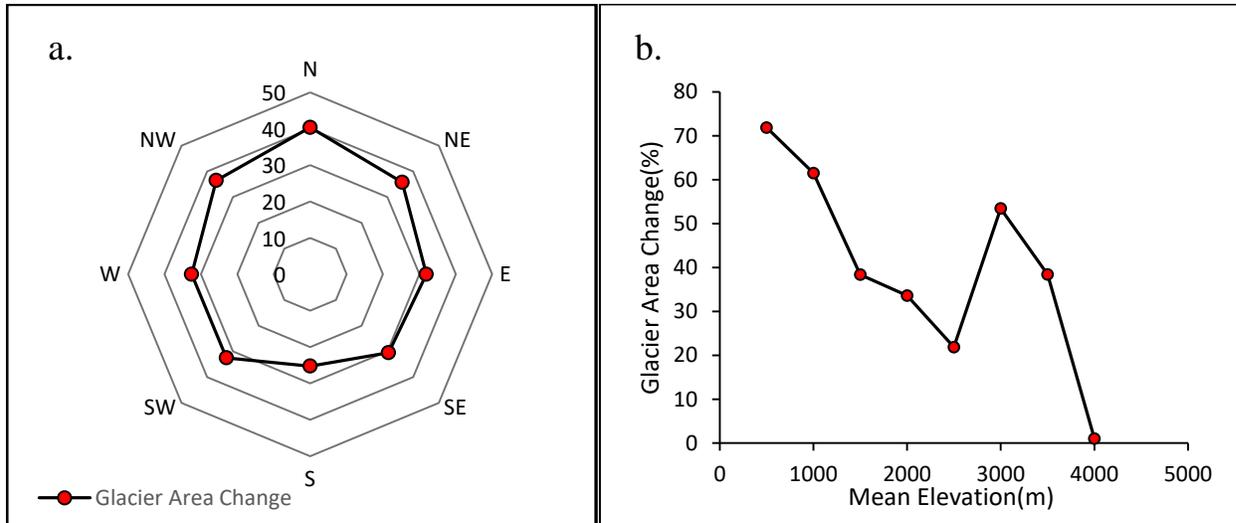
336 area of glaciers $<0.5 \text{ km}^2$ in this state decreased in area extent by 35.84% (-44.59 km^2) while the
337 total area of glaciers $>0.5 \text{ km}^2$ declined by 26.67% (85.47 km^2). These findings underscore that
338 larger glaciers in Washington state have a more profound impact on local water resources due to
339 their greater contribution to ice loss.

340 Glaciers across Oregon, California, Idaho, Oregon, Montana, and Wyoming make up
341 only 27% (117.27 km^2) of the total area of glaciers in this study with a combined mean area of
342 0.051 km^2 . The number of glaciers less than 0.5 km^2 accounted for 99% (2266 glaciers) of the
343 glacier population in these states. Glaciers in California and Oregon also experienced substantial
344 ice loss with both states with both states seeing glacier areas decline by over 60%. In contrast,
345 Montana and Wyoming recorded comparatively less shrinkage, with reductions of around 40%
346 (27.09 km^2) and 33% (23.12 km^2), respectively. The most drastic shrinkage occurred in Idaho
347 which shrunk over 90% of its glacier coverage. This can be attributed to the few numbers of
348 glaciers in Idaho, coupled with their smaller size, rendering them exceptionally vulnerable to
349 accelerated melting.

350 The study area has experienced changes in the number of glaciers, either through the
351 fragmentation of larger glaciers or the disappearance of smaller glaciers. This study found
352 complete disappearance of 930 glaciers. All states experienced a decrease in glaciers except
353 Nevada with an unchanged number of glaciers. The number of glaciers $<0.01 \text{ km}^2$ increased
354 from 113 in the RGI inventory to 1308 in this study, with a corresponding increase in glacier
355 area from 1.1 km^2 to 7.2 km^2 . This is directly related to a 43.15% decrease in the number of
356 glaciers between $0.01\text{-}10 \text{ km}^2$ from the RGI inventory. The total number of glaciers that
357 disappeared accounts for 18.5% of the total number of glaciers in the RGI inventory. The highest
358 rate of glacier disappearance occurred in California, which lost 356 glaciers. It could be
359 explained by the fact that this area mainly consists of smaller glaciers, which are more detached
360 than larger glaciers, making them more prone to melting (Intsiful & Ambinakudige, 2021). We
361 discovered 408 glaciers that grew, most of which are small glaciers. The area gain of individual
362 glaciers varied from 0.00 km^2 to 0.163 km^2 .

363 Glacier variability in the western United States shows that glaciers at lower elevations
364 lose more area than those at higher elevations (Figure. 7). The rapid decline in glacier size at
365 lower elevations can be linked to temperature rise. This is because glaciers at lower elevations
366 (ablation zone), are highly sensitive to temperature variations while glaciers at higher elevations
367 (accumulation zone), are more sensitive to trends in precipitation (Tian et al., 2014).

368 Also, as shown in Figure 7, glaciers in the north, northeast and northwest facing aspects
369 experienced the most shrinkage with a reduction of -72.9 km^2 (-40.4%), -68.5 km^2 (-35.8%) and
370 -26 km^2 (-36.5%) respectively. A factor that may have contributed to the faster shrinkage in the
371 north, northeast and northwest facing aspect glaciers could be the size of the glaciers. Majority of
372 glaciers with these aspects (N, NE, NW) tend to be smaller (<0.1). However, other factors such
373 as elevation, relative humidity, and temperature may also contribute to the shrinkage, but a
374 comprehensive analysis of these factors would require further research and investigation.



375 Figure 7. Glacier Area Changes at different (a.) orientation and (b.) elevation. All percentage
376 changes are negative.

377 Shrinkage patterns varied across various mountain ranges (Table 3). There was an
378 indication of a drastic decrease in glacier extent. In particular, the Trinity Alps, Sawtooth Range
379 and Maddison Range lost more area than other mountain ranges. In the north cascades national
380 park complex in the Cascade Range, glaciers and perennial snowfields area have declined by
381 29.3% (83.62 km²) of its coverage from the RGI glacier boundaries, with major glaciers such as
382 White River glacier, West Nooksack glacier, and White Chuck glacier losing more than 50
383 percent of their area. A large area reduction was found in glaciers on Mount Baker (-12.32 km², -
384 24.24%), Mount Shuksan (-4.05 km², -23.72%), and Mount Challenger (-3.38 km², -25.12%) in
385 the North Cascade Range, Washington. The glaciers in the Olympic mountains and Glacier
386 National Park of the Lewis range in the Rocky Mountains only have 66% and 69% of their
387 original glacier boundary covered by snow/ice in 2020. Glaciers such as Blue glacier (-0.747
388 km², 13.19%), White glacier (-0.719 km², -16.78%), Hoh glacier (-0.826 km², -18.22%) and
389 Sperry glacier (-0.491 km², -38.09%) declined at a rate of 2.26%a⁻¹, -2.19%a⁻¹, -2.5%a⁻¹ and -
390 0.91%a⁻¹ respectively. The Wind River Range and the Teton Range of Wyoming have lost -
391 31.79% (-17.78 km²) and -39.65% (-2.533 km²) respectively, with an average rate of a single
392 glacier being -0.076%a⁻¹ and -0.027%a⁻¹ respectively. The Gannett Glacier in the Wind River
393 Range, which is the largest glacier in the contiguous United States excluding Washington State,
394 currently retains only 85% of its original coverage. It is diminishing at an approximate rate of
395 1.002% a⁻¹.

396 5 Discussion and Conclusion

397 From the Randolph Glacier Inventory (RGI), the Western United States (excluding
398 Alaska) had 5021 glaciers with a combined area of 669.08 km² with a corresponding volume of
399 13.92 km³. By 2020 the number of glaciers had decreased to 4091, with the total area shrinking
400 to 432.01 km² and the volume declining to 9.02 km³ as well. Over the study period, combined
401 glacier area decreased by 237.07 km² (35.43%) and volume loss of 4.9 km³, indicating a general
402 decrease in glacier extent in the study area during the study period.

403 Evidence from various studies monitoring changes in glacier area over time in the western U.S.,
404 reveals notable reductions in glacier area extent (Fountain et al., 2017; Riedel et al., 2015;
405 Selkowitz & Forster, 2016). Generally, results from this study indicate a substantial decrease in
406 glacier extent in the American West from the mid-20th century to 2020. The total area of glaciers
407 and snowfields reduction by 35.43% (237.07 km²) in our study is comparable with the findings
408 from Selkowitz & Forster, (2016) who reported a total area change of 163.6 km² with a
409 corresponding area change of 28% from 2010 to 2014. Similarly, while the study conducted by
410 (Fountain et al., 2017) reported a 39% decrease in glacier and perennial snowfield coverage, our
411 study reveals a slightly lower ice coverage decrease of 35.43%. However, it is crucial to consider
412 the specific findings and methodologies of both studies when interpreting these results.

413 Examination of glacier extent changes across distinct regions in the western U.S. reveals
414 consistent patterns of glacier area reduction. For instance, the glacier area within North Cascades
415 National Park has decreased by 29.3% according to the present study. This finding corresponds
416 to prior research by (Granshaw & Fountain, 2006), which reported a 7% reduction in glacier area
417 in the North Cascades region between 1958 and 1998. Additionally, Selkowitz & Forster, (2016)
418 documented a 21% loss of area in the North Cascade Range between 2010 and 2014. In a similar
419 vein, the Wind River range in Wyoming has experienced a 31.79% reduction in glacier area.
420 DeVisser & Fountain, (2015) observed comparable trends, noting a 26.9% and 47% reduction in
421 glacier extent from 1963 to 2006 and 1900 to 2006, respectively. Furthermore, Maloof et al.,
422 (2014) reported a 39% reduction in glacier area across 44 studied glaciers in this region. These
423 findings collectively highlight the consistent trajectory of glacier area decline across different
424 regions in the western U.S.

425 As expected from the area loss, the glaciers also experienced a decrease in volume.
426 Volume estimates based on area-volume scaling indicate a volume loss of about 4.9 km³. This is
427 equivalent to about 4.7 gigatons of water.

428 Despite the fact that a number of factors, including wind speed, aspect, elevation, cloud
429 cover, relative humidity, avalanche, glacier size, and glacier type, can affect glacier recession or
430 growth, temperature and precipitation are the primary drivers of glacier variability. Increasing
431 temperatures and snowpack decline are the main factors influencing glacier loss in the United
432 States. Vose et al., (2017) observed a similar trend of increasing temperatures in the United
433 States since the start of the 20th century. Similarly, (Basagic & Fountain, 2011; Mote et al.,
434 2005, 2016, 2018) reported a similar pattern of decreasing snowpack in the western United
435 States. They emphasized that anthropogenic activities and sea surface temperature (SST) are
436 factors that brought about snow drought across the Pacific states, with SST contributing twice
437 the amount of anthropogenic influence. Therefore, if climate models continue to predict higher
438 temperatures and reduced precipitation in the western United States, the ongoing trend of glacier
439 retreat is expected to persist.

440 A number of issues arise as a result of glacier shrinkage, including glacier hazards such
441 as ice avalanches and glacial lake outburst floods (GLOFs), the rise of sea level and changes in
442 hydrology (Intsiful & Ambinakudige, 2021). From the satellite imagery used in this study as well
443 as images from Google Earth, we noticed that several proglacial lakes have been formed and
444 developed. As glaciers shrink, GLOFs may increase. Therefore, considering all these factors

445 mentioned above, continuous monitoring of glaciers to gather valuable data on their changing
 446 conditions holds great significance in understanding the impact of global climate change, the rise
 447 in sea levels, and ensuring the sustenance of the communities downstream from the glaciers.

448 The new glacier inventory created in this study (a complete list of glaciers will be
 449 published online) can be used as a baseline for future glacier change assessment study in the
 450 western U.S. Its accuracy would equip researchers and policymakers with reliable data,
 451 facilitating the monitoring and comprehensive understanding of the dynamic condition of these
 452 glaciers over time. However, there are several uncertainties which we have faced in the area
 453 change assessment that potentially have an impact on the quality of the inventory. Uncertainties
 454 could be caused due to the resolution of the satellite images, amount of seasonal snow remains,
 455 the time of satellite image acquisition, image classification and digitization accuracies.

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457 Reference

- 458 1. Ambinakudige, S., & Intsiful, A. (2022). Estimation of area and volume change in the
 459 glaciers of the Columbia Icefield, Canada using machine learning algorithms and Landsat
 460 images. *Remote Sensing Applications: Society and Environment*, 26.
 461 <https://doi.org/10.1016/j.rsase.2022.100732>
- 462 2. Ambinakudige, S., & Joshi, K. (2012). Remote sensing of cryosphere, *Remote Sensing-*
 463 *Applications*, Dr. Boris Escalante (Ed).
- 464 3. Ambinakudige, S., & Joshi, K. (2015). Multi-Decadal Changes in Glacial Parameters of
 465 the Fedchenko Glacier in Tajikistan. *International Journal of Advanced Remote Sensing*
 466 *and GIS*, 4(1), 911–919. <https://doi.org/10.23953/cloud.ijarsg.86>
- 467 4. Banko, G. (1998). A Review of Assessing the Accuracy of Classifications of Remotely
 468 Sensed Data and of Methods Including Remote Sensing Data in Forest Inventory.
 469 www.iiasa.ac.at
- 470 5. Basagic, H., & Fountain, A. (2011). Quantifying 20th century glacier change in the Sierra
 471 Nevada, California. *Arctic, Antarctic, and Alpine Research*, 43(3), 317–330.
 472 <https://doi.org/10.1657/1938-4246-43.3.317>
- 473 6. Berthier, E., Schiefer, E., Clarke, G. K. C., Menounos, B., & Rémy, F. (2010).
 474 Contribution of Alaskan glaciers to sea-level rise derived from satellite imagery. *Nature*
 475 *Geoscience*, 3(2), 92–95. <https://doi.org/10.1038/ngeo737>
- 476 7. Bolch, T., Menounos, B., & Wheate, R. (2010). Landsat-based inventory of glaciers in
 477 western Canada, 1985-2005. *Remote Sensing of Environment*, 114(1), 127–137.
 478 <https://doi.org/10.1016/j.rse.2009.08.015>
- 479 8. Chen, J., & Ohmura, A. (1990). Hydrology in Mountainous Regions. I-Uydrological
 480 Measurements; the Water Cycle (Proceedings of two Lausanne Symposia (Issue 193)).
 481 IAHS Publ. <https://www.researchgate.net/publication/241663882>

- 482 9. DeVisser, M. H., & Fountain, A. G. (2015). A century of glacier change in the Wind
483 River Range, WY. *Geomorphology*, 232, 103–116.
484 <https://doi.org/10.1016/j.geomorph.2014.10.017>
- 485 10. Dixon, L., & Ambinakudige, S. (2015). Remote Sensing Study of Glacial Change in the
486 Northern Patagonian Icefield. *Advances in Remote Sensing*, 04(04), 270–279.
487 <https://doi.org/10.4236/ars.2015.44022>
- 488 11. Drusch, M., Del Bello, U., Carlier, S., Colin, O., Fernandez, V., Gascon, F., Hoersch, B.,
489 Isola, C., Laberinti, P., Martimort, P., Meygret, A., Spoto, F., Sy, O., Marchese, F., &
490 Bargellini, P. (2012). Sentinel-2: ESA's Optical High-Resolution Mission for GMES
491 Operational Services. *Remote Sensing of Environment*, 120, 25–36.
492 <https://doi.org/10.1016/j.rse.2011.11.026>
- 493 12. Fountain, A. G., Glenn, B., & Basagic, H. J. (2017). The Geography of Glaciers and
494 Perennial Snowfields in the American West. *Arctic, Antarctic, and Alpine Research*,
495 49(3), 391–410. <https://doi.org/10.1657/AAAR0017-003>
- 496 13. Fountain, A. G., Gray, C., Glenn, B., Menounos, B., Pflug, J., & Riedel, J. L. (2022).
497 Glaciers of the Olympic Mountains, Washington—The Past and Future 100 Years.
498 *Journal of Geophysical Research: Earth Surface*, 127(4).
499 <https://doi.org/10.1029/2022JF006670>
- 500 14. Fountain, Hoffman, M., Jackson, K., Hassan, B., Thomas, N., & David, P. (2007). Digital
501 Outlines and Topography of the Glaciers of the American West. US Geological Survey.
502 <https://doi.org/10.3133/ofr20061340>
- 503 15. Gardner, A. S., Moholdt, G., Cogley, J. G., Wouters, B., Arendt, A. A., Wahr, J.,
504 Berthier, E., Hock, R., Pfeffer, W. T., Kaser, G., Ligtenberg, S. R. M., Bolch, T., Sharp,
505 M. J., Hagen, J. O., Van Den Broeke, M. R., & Paul, F. (2013). A reconciled estimate of
506 glacier contributions to sea level rise: 2003 to 2009. *Science*, 340(6134), 852–857.
507 <https://doi.org/10.1126/science.1234532>
- 508 16. Granshaw, F. D., & Fountain, A. G. (2006). Glacier change (1958-1998) in the North
509 Cascades National Park Complex, Washington, USA. *Journal of Glaciology*, 52(177),
510 251–256. <https://doi.org/10.3189/172756506781828782>
- 511 17. Horning, N. (2010). Random Forests : An algorithm for image classification and
512 generation of continuous fields data sets.
- 513 18. Huete, A., Didan, K., Miura, T., Rodriguez, E. P., Gao, X., & Ferreira, L. G. (2002).
514 Overview of the radiometric and biophysical performance of the MODIS vegetation
515 indices. *Remote Sensing of Environment*, 83, 195–213. www.elsevier.com/locate/rse
- 516 19. Hugonnet, R., McNabb, R., Berthier, E., Menounos, B., Nuth, C., Girod, L., Farinotti, D.,
517 Huss, M., Dussailant, I., Brun, F., & Käab, A. (2021). Accelerated global glacier mass
518 loss in the early twenty-first century. *Nature*, 592(7856), 726–731.
519 <https://doi.org/10.1038/s41586-021-03436-z>
- 520 20. Huss, M., & Hock, R. (2018). Global-scale hydrological response to future glacier mass
521 loss. *Nature Climate Change*, 8(2), 135–140. <https://doi.org/10.1038/s41558-017-0049-x>

- 522 21. Inamdar, P., & Ambinakudige, S. (2016). Spatial patterns of glacier mass change in the
523 Southern Andes. In *Photogrammetric Engineering and Remote Sensing* (Vol. 82, Issue
524 10, pp. 811–818). American Society for Photogrammetry and Remote Sensing.
525 <https://doi.org/10.14358/PERS.82.10.811>
- 526 22. Intsiful, A., & Ambinakudige, S. (2021). Glacier cover change assessment of the
527 Columbia Icefield in the Canadian rocky mountains, Canada (1985-2018). *Geosciences*
528 (Switzerland), 11(1), 1–9. <https://doi.org/10.3390/geosciences11010019>
- 529 23. IPCC, 2023: Climate Change 2023: Synthesis Report. Contribution of Working Groups I,
530 II and III to the Sixth Assessment Report of the Intergovernmental Panel on Climate
531 Change [Core Writing Team, H. Lee and J. Romero (eds.)]. IPCC, Geneva, Switzerland,
532 184 pp., doi: <https://doi.org/10.59327/IPCC/AR6-9789291691647>.
- 533 24. Keshri, A. K., Shukla, A., & Gupta, R. P. (2009). ASTER ratio indices for supraglacial
534 terrain mapping. *International Journal of Remote Sensing*, 30(2), 519–524.
535 <https://doi.org/10.1080/01431160802385459>
- 536 25. Maloof, A., Piburn, J., Tootle, G., & Kerr, G. (2014). Recent Alpine Glacier Variability:
537 Wind River Range, Wyoming, USA. *Geosciences*, 4(3), 191–201.
538 <https://doi.org/10.3390/geosciences4030191>
- 539 26. Martin-Mikle, C. J., & Fagre, D. B. (2019). Glacier recession since the Little Ice Age:
540 Implications for water storage in a Rocky Mountain landscape. *Arctic, Antarctic, and*
541 *Alpine Research*, 51(1), 280–289. <https://doi.org/10.1080/15230430.2019.1634443>
- 542 27. McCabe, G. J., & Fountain, A. G. (2013). Glacier variability in the conterminous United
543 States during the twentieth century. *Climatic Change*, 116(3–4), 565–577.
544 <https://doi.org/10.1007/s10584-012-0502-9>
- 545 28. Meier, M. F., 1961: Distribution and variations of glaciers in the United States exclusive
546 of Alaska. International Association of Hydrological Sciences Publications, 54: 420–429.
- 547 29. Moore, R. D., Fleming, S. W., Menounos, B., Wheate, R., Fountain, A., Stahl, K., Holm,
548 K., & Jakob, M. (2009). Glacier change in western North America: Influences on
549 hydrology, geomorphic hazards and water quality. In *Hydrological Processes* (Vol. 23,
550 Issue 1, pp. 42–61). <https://doi.org/10.1002/hyp.7162>
- 551 30. Mote, P. W., Hamlet, A. F., Clark, M. P., & Lettenmaier, D. P. (2005). Declining
552 mountain snowpack in western north America. *Bulletin of the American Meteorological*
553 *Society*, 86(1), 39–49. <https://doi.org/10.1175/BAMS-86-1-39>
- 554 31. Mote, P. W., Rupp, D. E., Li, S., Sharp, D. J., Otto, F., Uhe, P. F., Xiao, M., Lettenmaier,
555 D. P., Cullen, H., & Allen, M. R. (2016). Perspectives on the causes of exceptionally low
556 2015 snowpack in the western United States. *Geophysical Research Letters*, 43(20),
557 10,980-10,988. <https://doi.org/10.1002/2016GL069965>
- 558 32. Mote, P. W., Li, S., Lettenmaier, D. P., Xiao, M., & Engel, R. (2018). Dramatic declines
559 in snowpack in the western US. *Npj Climate and Atmospheric Science*, 1(1).
560 <https://doi.org/10.1038/s41612-018-0012-1>

- 561 33. O’Neel, S., McNeil, C., Sass, L. C., Florentine, C., Baker, E. H., Peitzsch, E., McGrath,
562 D., Fountain, A. G., & Fagre, D. (2019). Reanalysis of the US Geological Survey
563 Benchmark Glaciers: Long-term insight into climate forcing of glacier mass balance.
564 *Journal of Glaciology*, 65(253), 850–866. <https://doi.org/10.1017/jog.2019.66>
- 565 34. Paul, F., Bolch, T., Briggs, K., Kääb, A., McMillan, M., McNabb, R., Nagler, T., Nuth,
566 C., Rastner, P., Strozzi, T., & Wuite, J. (2017). Error sources and guidelines for quality
567 assessment of glacier area, elevation change, and velocity products derived from satellite
568 data in the Glaciers_cci project. *Remote Sensing of Environment*, 203, 256–275.
569 <https://doi.org/10.1016/j.rse.2017.08.038>
- 570 35. Pelto, M. S., Dryak, M., Pelto, J., Matthews, T., & Perry, L. B. (2022). Contribution of
571 Glacier Runoff during Heat Waves in the Nooksack River Basin USA. *Water*
572 (Switzerland), 14(7). <https://doi.org/10.3390/w14071145>
- 573 36. Pfeffer, W. T., Arendt, A. A., Bliss, A., Bolch, T., Cogley, J. G., Gardner, A. S., Hagen,
574 J. O., Hock, R., Kaser, G., Kienholz, C., Miles, E. S., Moholdt, G., Mölg, N., Paul, F.,
575 Radić, V., Rastner, P., Raup, B. H., Rich, J., Sharp, M. J., ... Wyatt, F. R. (2014). The
576 randolph glacier inventory: A globally complete inventory of glaciers. *Journal of*
577 *Glaciology*, 60(221), 537–552. <https://doi.org/10.3189/2014JoG13J176>
- 578 37. RGI Consortium, 2017. Randolph Glacier Inventory - A Dataset of Global Glacier
579 Outlines, Version 6. [Indicate subset used]. Boulder, Colorado USA. NSIDC: National
580 Snow and Ice Data Center. doi: <https://doi.org/10.7265/4m1f-gd79>
- 581 38. Riedel, J. L., & Larrabee, M. A. (2016). Impact of Recent Glacial Recession on Summer
582 Streamflow in the Skagit River. *Northwest Science*, 90(1), 5–22.
583 <https://doi.org/10.3955/046.090.0103>
- 584 39. Riedel, J. L., Wilson, S., Baccus, W., Larrabee, M., Fudge, T. J., & Fountain, A. (2015).
585 Glacier status and contribution to streamflow in the Olympic Mountains, Washington,
586 USA. *Journal of Glaciology*, 61(225), 8–16. <https://doi.org/10.3189/2015JoG14J138>
- 587 40. Rounce, D. R., Hock, R., Maussion, F., Hugonnet, R., Kochtitzky, W., Huss, M.,
588 Berthier, E., Brinkerhoff, D., Compagno, L., Copland, L., Farinotti, D., Menounos, B., &
589 McNabb, R. W. (2023). Global glacier change in the 21st century: Every increase in
590 temperature matters. In *Science* (Vol. 379). <https://www.science.org>
- 591 41. Russell, I. C., 1897: *Glaciers of North America*. Boston: Ginn & Company, Athenaeum
592 Press, 210 pp.
- 593 42. Selkowitz, D. J., & Forster, R. R. (2016). Automated mapping of persistent ice and snow
594 cover across the western U.S. with Landsat. *ISPRS Journal of Photogrammetry and*
595 *Remote Sensing*, 117, 126–140. <https://doi.org/10.1016/j.isprsjprs.2016.04.001>
- 596 43. Singh, S., Kumar, R., & Dimri, A. P. (2018). Mass balance status of Indian Himalayan
597 glaciers: A brief review. In *Frontiers in Environmental Science* (Vol. 6, Issue AUG).
598 Frontiers Media S.A. <https://doi.org/10.3389/fenvs.2018.00030>

- 599 44. Tennant, C., & Menounos, B. (2013). Glacier change of the Columbia Icefield, Canadian
600 Rocky Mountains, 1919-2009. *Journal of Glaciology*, 59(216), 671–686.
601 <https://doi.org/10.3189/2013JoG12J135>
- 602 45. Tian, H., Yang, T., & Liu, Q. (2014). Climate change and glacier area shrinkage in the
603 Qilian mountains, China, from 1956 to 2010. *Annals of Glaciology*, 55(66), 187–197.
604 <https://doi.org/10.3189/2014AoG66A045>
- 605 46. Wouters, B., Gardner, A. S., & Moholdt, G. (2019). Global glacier mass loss during the
606 GRACE satellite mission (2002-2016). *Frontiers in Earth Science*, 7.
607 <https://doi.org/10.3389/feart.2019.00096>
- 608 47. Zemp, M., Huss, M., Thibert, E., Eckert, N., McNabb, R., Huber, J., Barandun, M.,
609 Machguth, H., Nussbaumer, S. U., Gärtner-Roer, I., Thomson, L., Paul, F., Maussion, F.,
610 Kutuzov, S., & Cogley, J. G. (2019). Global glacier mass changes and their contributions
611 to sea-level rise from 1961 to 2016. *Nature*, 568(7752), 382–386.
612 <https://doi.org/10.1038/s41586-019-1071-0>