

1 **Regression forest approaches to gravity wave**
2 **parameterization for climate projection**

3 **David S. Connelly¹, Edwin P. Gerber¹**

4 ¹Center for Atmosphere Ocean Science, Courant Institute, New York University

5 **Key Points:**

- 6 • Two kinds of regression forest emulate a gravity wave parameterization offline and
7 online, with boosted forests outperforming random forests
- 8 • Relative to a neural network benchmark, the boosted forest exhibits similar online
9 skill and ability to generalize to new climates
- 10 • Feature importance analysis informs a retraining procedure to improve online behavior
11 of data-driven parameterizations

Corresponding author: David S. Connelly, dsconnelly@nyu.edu

Abstract

We train random and boosted forests, two machine learning architectures based on regression trees, to emulate a physics-based parameterization of atmospheric gravity wave momentum transport. We compare the forests to a neural network benchmark, evaluating both offline errors and online performance when coupled to an atmospheric model under the present day climate and in 800 and 1200 ppm CO₂ global warming scenarios. Offline, the boosted forest exhibits similar skill to the neural network, while the random forest scores significantly lower. Both forest models couple stably to the atmospheric model, and control climate integrations with the boosted forest exhibit lower biases than those with the neural network. Integrations with all three data-driven emulators successfully capture the Quasi-Biennial Oscillation (QBO) and sudden stratospheric warmings, key modes of stratospheric variability, with the boosted forest more accurate than the random forest in replicating their statistics across our range of carbon dioxide perturbations. The boosted forest and neural network capture the sign of the QBO period response to increased CO₂, though both struggle with the magnitude of this response under the more extreme 1200 ppm scenario. To investigate the connection between performance in the control climate and the ability to generalize, we use techniques from interpretable machine learning to understand how the data-driven methods use physical information. We leverage this understanding to develop a retraining procedure that improves the coupled performance of the boosted forest in the control climate and under the 800 ppm CO₂ scenario.

Plain Language Summary

Parameterizations are reduced-complexity models that estimate the effects of physical processes smaller than what can be resolved by the grid of a weather or climate model. While necessary for realistic simulations, they are a source of uncertainty in climate projections. Recently, machine learning has been used to augment or replace conventional parameterizations of atmospheric gravity waves, a type of motion by which disturbances near the Earth's surface can affect the wind higher up. We compare several machine learning approaches to the gravity wave parameterization problem. In particular, we test neural networks against random and boosted forests, which are built around flowchart-like models called regression trees. We find that boosted forests, though not widely used for climate model parameterization, are especially successful, scoring as well as or better than neural networks on various performance metrics. We then provide proof-of-concept of a novel method to retrain the

44 boosted forest so that it uses its input data more in line with the physics of the system,
45 and show that this technique improves the forest’s behavior when used together with an
46 atmospheric model.

47 **1 Introduction**

48 Momentum transport by atmospheric gravity waves is a key driver of several features
49 of the large-scale circulation, such as the Quasi-Biennial Oscillation (QBO) in the tropical
50 stratosphere (Fritts & Alexander, 2003), and influences others, including the tropospheric
51 jet structure (Palmer et al., 1986). Waves transporting appreciable momentum can have
52 horizontal and vertical wavelengths as small as 100 m. Because of the small scales at play,
53 atmospheric models must rely on parameterizations of gravity wave momentum flux to
54 represent these climate features (Anstey et al., 2022).

55 Gravity wave parameterizations (GWPs), however, must make simplifying assumptions
56 about wave propagation to remain computationally feasible. The dynamics may be derived
57 from the linearized equations of motion, for example, assuming planar waves that do not
58 interact with each other. In addition, schemes are in general computationally constrained
59 to operate on single atmospheric columns, so that they cannot model lateral propagation.
60 Due to these limitations, GWPs are significant sources of model uncertainty. For exam-
61 ple, parameterizations tuned to match present-day QBO statistics can produce different
62 projections of QBO behavior in warmer climate simulations (Richter et al., 2022).

63 Recently, machine learning has been used to build new parameterizations of subgrid-
64 scale phenomena including ocean eddy momentum forcings (Bolton & Zanna, 2019), radiative
65 and microphysical tendencies in atmospheric moisture variables (Yuval & O’Gorman,
66 2020), and gravity wave momentum transport (Chantry et al., 2021; Espinosa et al., 2022).
67 Such approaches attempt to learn a mapping from resolved-scale variables to subgrid quan-
68 tities from data. Machine learning can be employed both to learn these relationships from
69 increasingly available high-resolution data (Brenowitz & Bretherton, 2018; O’Gorman &
70 Dwyer, 2018; Bolton & Zanna, 2019; Yuval & O’Gorman, 2020) and to accelerate existing
71 parameterizations by replacing them with emulators (Chevallier et al., 1998; Chantry et al.,
72 2021).

73 Building on work by Espinosa et al. (2022), we explore the use of machine learning
74 models to emulate the behavior of an existing, physics-based reference parameterization,

75 the Alexander and Dunkerton (1999) scheme (hereafter AD99) as it is implemented in
76 the Model of an idealized Moist Atmosphere (MiMA), an intermediate-complexity climate
77 model (Garfinkel et al., 2020). Emulation serves as an intermediate step towards training
78 fully data-driven GWPs using a combination of observations of gravity waves and high-
79 resolution simulations capable of resolving them.

80 In particular, emulation allows us to address challenges inherent to the design of data-
81 driven GWPs — e.g. comparing machine learning architectures, achieving stable coupling
82 with atmospheric models, and testing the ability of schemes to generalize — separately
83 from issues of data availability and processing that arise when working with more realistic
84 data sources. In the emulation problem, data is cheap to generate by running the reference
85 parameterization. Moreover, AD99 produces a reference climate when coupled to MiMA,
86 which can be perturbed by increasing the carbon dioxide concentration. As a result, there are
87 straightforward performance metrics when coupling data-driven emulators and evaluating
88 their performance in a warmer climate.

89 Espinosa et al. (2022) used neural networks to emulate AD99, finding that they could
90 reproduce both the QBO and its AD99-predicted response to an increase in CO₂. The
91 principal contributions of this work are the application of tree-based machine learning ar-
92 chitectures to the same problem, the comparison of their offline and online performance with
93 that of neural networks, and the use of offline feature importance analyses to develop a re-
94 training procedure that improves online behavior. Section 2 introduces the parameterization
95 to be emulated and characterizes the datasets used. Section 3 describes the machine learn-
96 ing architectures along with the interpretability techniques used to analyze them. Section 4
97 presents the performance of the emulators, both offline and in coupled integrations under a
98 series of CO₂ perturbations. Section 5 analyzes the emulators’ preferential use of particular
99 input features to explain their offline and online behavior using so-called interpretable ma-
100 chine learning techniques. Section 6 reviews the main results and discusses future research
101 directions.

102 **2 Models and data**

103 We begin with descriptions of the atmospheric model MiMA and the reference param-
104 eterization AD99. These are relevant to our work in that they are used to generate the

105 training and offline test datasets, the main features of which are outlined in Section 2.2, and
106 in that they are necessary for the coupled integration experiments summarized in Section 2.3.

107 **2.1 MiMA and the AD99 parameterization**

108 The Model of an idealized Moist Atmosphere (MiMA) is an atmospheric circulation
109 model coupled with a purely thermodynamic, or slab, ocean model. The atmosphere com-
110 ponent includes interactive moisture, full radiation with the Rapid Radiative Transfer Model
111 scheme (Mlawer et al., 1997; Iacono et al., 2000), and Betts-Miller convection (Betts, 1986;
112 Betts & Miller, 1986). The carbon dioxide concentration is set globally at 390 ppm, though
113 we change this value in experiments presented later in this work. Ozone is distributed
114 according to the monthly- and zonal-mean climatologies used in CMIP6 pre-industrial sim-
115 ulations (Checa-Garcia et al., 2018; Checa-Garcia, 2018). See Jucker and Gerber (2017)
116 and Garfinkel et al. (2020) for a complete description of MiMA.

117 The AD99 gravity wave parameterization takes in resolved-scale variables from a single
118 column and time step of an atmospheric model and returns the velocity tendencies from
119 parameterized gravity waves at each vertical level. The scheme computes the forcing by
120 propagating a collection of independent, monochromatic waves of varying phase speed from
121 the tropopause. Each wave has an associated momentum flux determined by a parameter-
122 ized source spectrum. That momentum flux spectrum, Gaussian in magnitude with width
123 35 m s^{-1} , is positive for phase speeds greater than the source level velocity and negative
124 elsewhere. The amplitude peak is at phase speed zero in the extratropics, where gravity
125 waves are assumed to be mainly orographic, and at phase speed equal to the source level
126 velocity in the tropics, where waves are assumed to be generated by convection. The source
127 spectrum and launch level vary only with latitude, a key simplification of the scheme.

128 Waves are propagated upwards until they reach a level where one of several breaking
129 criteria is met, where they deposit their momentum. Breaking occurs either when the
130 density is sufficiently low to permit overturning or at a *critical level*, where the mean flow
131 speed is very close to the wave’s phase speed and the vertical group velocity is accordingly
132 very small. Waves that reach the model top without breaking have their momentum flux
133 evenly distributed over the highest three vertical levels. We refer the reader to Alexander
134 and Dunkerton (1999) for a detailed discussion of the parameterization.

2.2 Training inputs and outputs

Machine learning parameterizations use data to learn a mapping from resolved scale variables to subgrid quantities — or, in this case, to learn the mapping encoded in AD99. To generate training data, we first run MiMA for 20 years, using AD99 as the gravity wave parameterization, to ensure the climate system has reached statistical equilibrium. We then integrate for a further 60 years, outputting data every six hours. At T42 resolution, this control run yields just under 12 million vertical profiles per year of model time. We sample 10 million profiles from the first 4 years of this 60-year run to use as training data, and another 10 million profiles from years 5-8 to use as an offline test set.

We explore the use of various resolved-scale variables as *input features*, the data passed as input to our machine learning emulators. These may include vertical profiles of wind u or v , temperature T , or buoyancy frequency N . We provide some models with a profile of the differences between winds at adjacent levels; this feature has units of m s^{-1} and, for brevity, we call it the *shear*. Every emulator takes in the latitude ϑ and surface pressure p_s of each training sample, except when we explicitly test the effects of withholding these features. We expect latitude to be an important feature because, in AD99, the peak of the gravity wave momentum flux source spectrum transitions from phase speed zero in the extratropics to phase speed moving with the tropopause flow in the tropics. The source level also varies with latitude to follow the tropopause.

The *targets*, the outputs associated with particular inputs that the machine learning model tries to predict, are the AD99 gravity wave accelerations at each of the 40 vertical levels in MiMA. Because AD99 handles zonal and meridional accelerations identically, our emulators are trained on both zonal and meridional data. When predicting zonal acceleration, they take u as an input feature, and likewise they take v when predicting meridional acceleration. Both the training and test datasets are evenly split between zonal and meridional samples.

The left panel of Figure 1 shows example zonal wind and temperature profiles, and the black curve in the right panel is the resulting gravity wave acceleration profile as parameterized by AD99. Note that the target accelerations vary by two to three orders of magnitude over the vertical column. We use the mean and variance of the accelerations at each vertical level, computed from the training samples, to standardize the targets to zero mean and unit variance. As a result, the emulators weight errors at each level equally, as

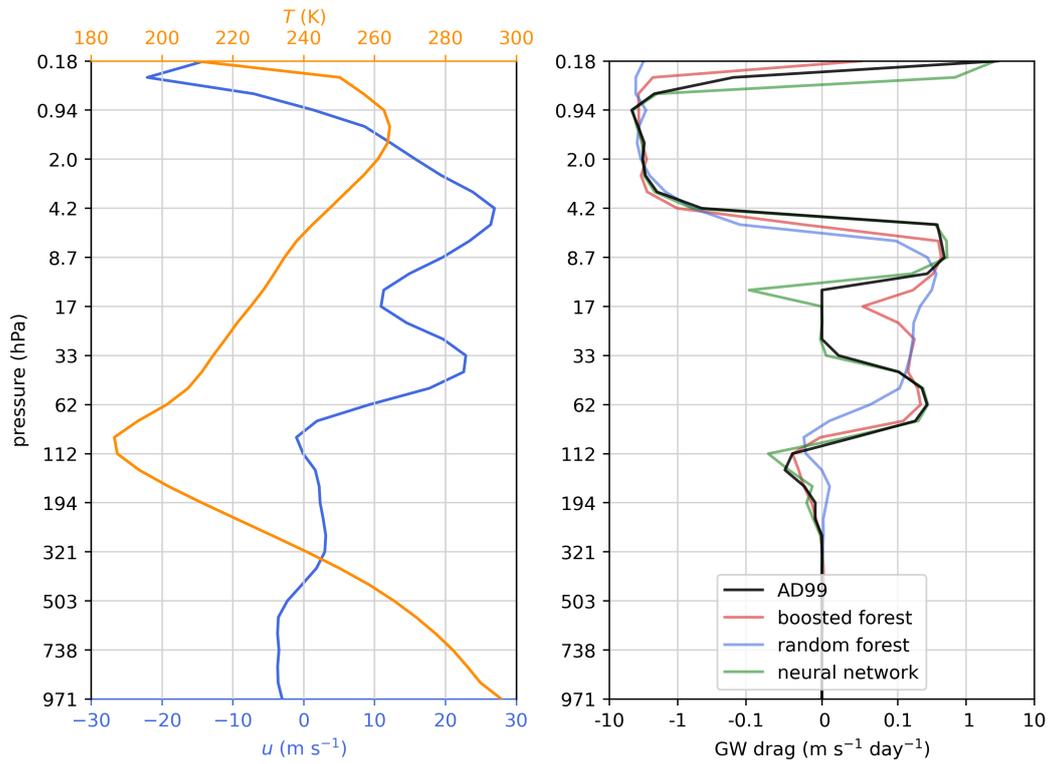


Figure 1. Example wind and temperature profiles (left) and the resulting parameterized accelerations predicted by AD99 and several emulators (right).

167 opposed to prioritizing performance near the model top at the expense of the lower levels.
 168 We rescale emulator predictions to their original units before using them to compute offline
 169 performance metrics like R^2 or passing them to MiMA in coupled runs.

170 **2.3 Coupled integrations**

171 After training and offline error analysis, we couple each trained emulator to MiMA,
 172 initialized with the final state of the spinup used to initialize the AD99 control run, and
 173 integrate for 60 years. The configuration is identical to that of the control run except that
 174 the emulators are used in place of AD99.

175 To assess the emulators' response to climate perturbations, we also run MiMA with
 176 the CO₂ concentration set at 800 ppm and at 1200 ppm. As with the 390 ppm CO₂ control
 177 runs, for each concentration value, we first integrate with AD99 for twenty years of spinup
 178 followed by sixty further years. We then couple each emulator (without retraining) and
 179 integrate for sixty years starting from the final state of the same spinup period.

180 For all coupled integrations, we retain only the last 56 years for analysis. Section 4
 181 discusses the output of these integrations.

182 **3 Machine learning architectures and interpretation**

183 In this section, we first review tree- and forest-based machine learning architectures,
 184 distinguishing between random and boosted forests, and specify the neural network bench-
 185 mark. Random forests have been used in atmospheric modeling before (O'Gorman & Dwyer,
 186 2018); however, we believe this work is the first use in this context of boosting, which is well-
 187 known in the broader machine learning literature. Next, we summarize the interpretability
 188 metric we use to analyze the behavior of our emulators. Finally, we indicate the existing
 189 libraries we used and briefly describe new software written for this study.

190 **3.1 Regression trees and forests**

191 The way humans solve problems is often well-approximated by asking a series of yes-no
 192 questions about the available data and predicting accordingly. For example, if asked to
 193 guess the price of a house, one might first ask if it is located in New York, then whether
 194 it has more than two bedrooms, and so on, perhaps selecting later questions based on the
 195 answers to earlier ones.

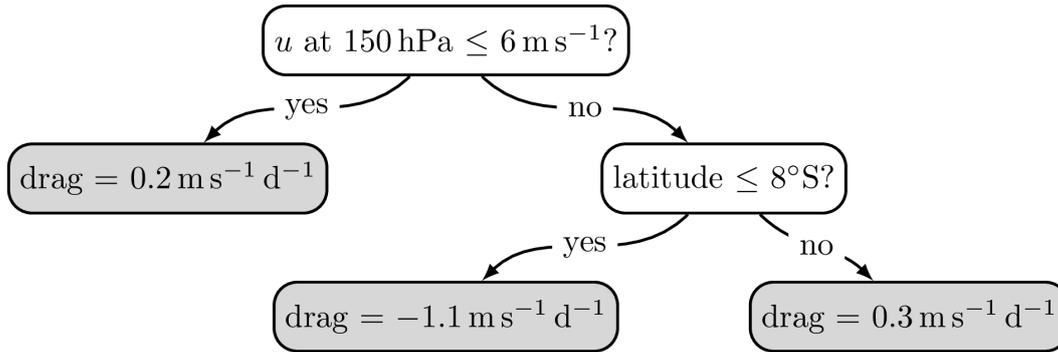


Figure 2. A simple regression tree. Leaf nodes are shaded. Note that the predictions of this example tree are scalars (e.g. accelerations at a particular level), while the trees used in this work yield vector-valued predictions (acceleration profiles).

196 Regression trees (Breiman et al., 1984) are machine learning models that attempt to
 197 make predictions in an analogous manner. Once trained, they make predictions by traversing
 198 a binary tree according to the given input features. The traversal repeatedly proceeds to
 199 one of the current node’s two children based on whether a specified input feature exceeds
 200 a set threshold. The tree returns as its prediction the value stored at whichever leaf node
 201 the traversal terminates at. Figure 2 shows a simple schematic of a regression tree with
 202 features relevant to the gravity wave parameterization problem. See Text S1 for a detailed
 203 explanation of how regression trees are constructed from training data.

204 If their depth is unlimited, regression trees can achieve zero error on the training data.
 205 However, such trees typically generalize very poorly to unseen samples because they have
 206 learned the noise in the dataset. Instead, a lower-variance model can be constructed from
 207 an ensemble, or *forest*, of regression trees of fixed depth. In this study we consider two kinds
 208 of ensembles: random forests and boosted forests.

209 A random forest (Breiman, 2001) is a collection of regression trees, each of which is
 210 trained independently on a bootstrapped subsample of the training dataset. The prediction
 211 of the forest is simply the mean of the predictions of each constituent tree. Figure 3(a)
 212 shows this concept: individual ensemble members have high error, but their ensemble mean
 213 matches the target much more closely.

214 The term boosting (Schapire, 1990) encompasses a wide class of machine learning algo-
 215 rithms that train individual ensemble members sequentially and combine them into a more

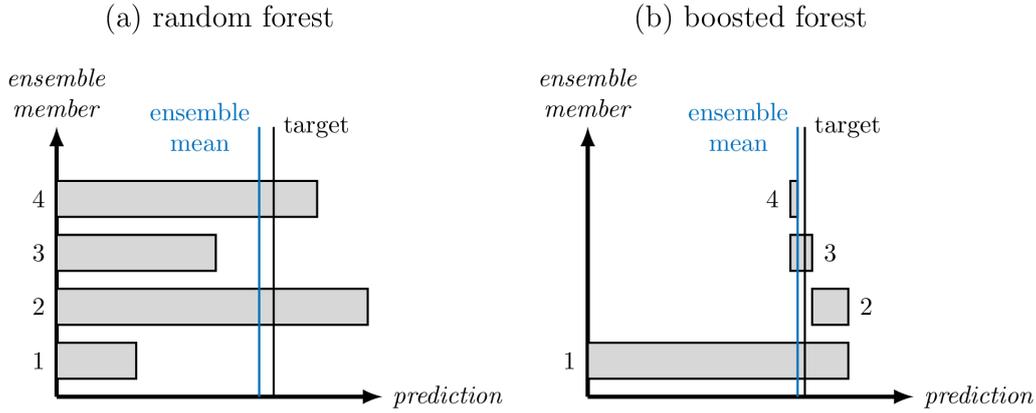


Figure 3. The types of regression forests considered in this work, shown conceptually. Each bar represents an individual tree in the ensemble, and the black “target” line is the true output associated with a particular sample. In (b), the bars for each tree are positioned relative to the sum of the preceding members. The blue line is the aggregate prediction of the ensemble — the mean in (a), and the sum in (b).

216 powerful model. In a boosted forest (Friedman, 2001), each tree is trained to reduce the
 217 residual errors accrued by its predecessors. Figure 3(b) illustrates how the trees in a boosted
 218 forest work to correct the under- or overshoot of the sum of the previous trees’ predictions.

219 3.2 Forest hyperparameters

220 The training of both random and boosted forests is governed by several *hyperparameters*,
 221 tunable values chosen by the user. Most of these fall into one of two broad categories: those
 222 which set the size of the ensemble, and those which inject randomness into the training
 223 process. The former includes the number of trees in the forest and the maximum allowed
 224 depth of each tree. The latter includes the size of the subsampled dataset used to train each
 225 tree and the fraction of input features considered as potential splits at each node. Boosted
 226 forests also have a *learning rate*, a scalar less than unity multiplying the prediction of each
 227 constituent tree, which limits the rate at which the forest can “zero in” on the targets and
 228 helps prevent overfitting to the training data.

229 Table S3 summarizes the hyperparameter values used in this study. They were chosen
 230 using cross-validation: for each candidate hyperparameter set, a model is trained several
 231 times with different subsets of the training data held out, and the parameter set that mini-

232 mizes the average out-of-set error is chosen. Hyperparameters were then kept constant for
 233 consistency across all experiments. We find that performance is robust to moderate changes
 234 in these hyperparameters.

235 **3.3 WaveNet as a benchmark**

236 We use the WaveNet neural network architecture, developed by Espinosa et al. (2022) to
 237 emulate AD99, as a benchmark against which to compare our forest emulators. The neural
 238 network features four shared fully-connected hidden layers followed by a branching structure
 239 with two independent fully-connected layers for each output pressure level. Our network
 240 has approximately 385,000 trainable parameters (the exact number depends on the input
 241 features chosen), roughly the minimum size found to be necessary by Espinosa et al. (2022)
 242 for successful coupled integrations. It is challenging to meaningfully compare parameter
 243 counts of neural networks with those of regression forests because their architectures differ
 244 so dramatically. The neural networks and regression forests we use have roughly comparable
 245 runtimes in coupled integrations, which is perhaps the more practically important metric of
 246 model complexity. Neither architecture was optimized for performance on our machine.

247 Our approach differs from that of Espinosa et al. (2022) in that we use mean-square
 248 error as our loss function (instead of the log cosh loss) and we predict gravity wave drags at
 249 all 40 output levels (instead of predicting at the highest 33 levels and padding with zeros).
 250 We made these changes for the sake of simplicity and did not observe significant effects on
 251 network behavior. As with our forest emulators, we train one neural network to predict
 252 both zonal and meridional accelerations.

253 **3.4 Feature importance and SHAP values**

254 Because machine learning architectures can be fairly opaque, it is often difficult to
 255 assess whether a model has learned to use its input data in a physically plausible way.
 256 Moreover, even data-driven schemes that achieve low error on their training and test sets
 257 can behave unpredictably when coupled back to the atmospheric model. This suggests that
 258 parameterization design might be well-served by studying not just how highly a particular
 259 model scores, but also how it uses its data.

260 Feature importance metrics attempt to understand the behavior of a machine learning
 261 model by quantifying how individual features affect model predictions. The *SHAP* (*SHapley*

262 *Additive exPlanation) value* (Lundberg & Lee, 2017), an adaptation of game-theoretic Shap-
 263 ley values (Shapley, 1953), is a metric of feature importance defined for arbitrary machine
 264 learning models. Given any function φ and a sample \mathbf{x} , first consider

$$265 \quad a_k(\mathbf{x}) = \mathbb{E}_{z_k} \left[\varphi(\mathbf{z}) \mid z_i = x_i \text{ for all } i \neq k \right] \quad (1)$$

266 the average output of φ on inputs matching \mathbf{x} except in the k^{th} component. The expectation
 267 in (1) is the interventional expectation (Pearl, 2000; Janzing et al., 2020) which, to avoid
 268 mistaking correlations between components for patterns in the behavior of φ , breaks the
 269 dependence of the k^{th} component on the others by averaging over the full distribution of
 270 z_k found in the training dataset. The SHAP value of feature k is then defined as $s_k(\mathbf{x}) \equiv$
 271 $\varphi(\mathbf{x}) - a_k(\mathbf{x})$, the change in φ that results when x_k is known exactly. Efficient algorithms
 272 exist for approximating $a_k(\mathbf{x})$, and by extension $s_k(\mathbf{x})$, for both regression trees (Lundberg
 273 et al., 2020) and neural networks (Shrikumar et al., 2017; Lundberg & Lee, 2017).

274 SHAP values are local, in the sense that they are calculated for each input sample.
 275 In Section 5 we compute dataset-averaged absolute values of the SHAP values for a global
 276 measure of feature importance. Moreover, the SHAP value as described is defined for scalar-
 277 valued functions, but the parameterizations we consider in this work have vector-valued
 278 outputs; we calculate SHAP values for each output channel separately and examine how
 279 features vary in their importance to predictions of gravity wave drag at different vertical
 280 levels.

281 3.5 Software implementation

282 The random and boosted forests we use are built using the decision tree interface in the
 283 Python library scikit-learn (Pedregosa et al., 2011), and our neural networks use PyTorch
 284 (Paszke et al., 2019). To couple our emulators with MiMA, we use Forpy (Rabel, 2020),
 285 which allows Python functions to be called from the Fortran numerical solver.

286 Support for multioutput regression, however, is limited in existing tree boosting li-
 287 braries. For scalar problems, boosting admits an optimization known as gradient boosting,
 288 but that formulation involves Taylor expansions in the output space and so becomes im-
 289 practical for multi-output target data. As such, we created Mubofa (*muti-output boosted*
 290 *forests*), a tree boosting library extending Scikit-learn and based on the train-on-the-residuals
 291 perspective of boosting schematized in Figure 3 (Connelly, 2023). Mubofa also implements
 292 the vector-valued Gini importance calculations described in Text S2.

4 Offline and online evaluation

We first evaluate our emulators *offline*; that is, we examine their skill on the training and test datasets described in Section 2.2, uncoupled from the atmospheric model that generated that data. We then discuss the emulators' *online* performance — how the atmospheric model behaves when coupled to a data-driven emulator instead of to AD99, as outlined in Section 2.3. Because parameterizations are developed with the goal of enhancing atmospheric models, online performance is of greater importance than offline error. However, it is more difficult both to improve directly, because emulator training occurs offline, and to evaluate, because doing so requires computationally demanding integrations of the atmospheric model.

4.1 Offline R^2 scores

The left panel of Figure 1 shows sample tropical zonal wind and temperature profiles passed by MiMA to the gravity wave scheme, and the right panel shows the gravity wave acceleration profiles as parameterized by AD99, one of each kind of regression forest emulator, and a neural network emulator. The AD99 profile exhibits local maxima just below the two maxima in the input wind profile, reflecting the deposition of momentum just below critical levels. The boosted forest and neural network emulators reproduce the shape of this profile, including the two maxima, although the neural network appears somewhat closer to the AD99 profile overall. The random forest, on the other hand, seems to smooth out many of the features of the target profile; this is unsurprising, since random forests have an intrinsic tendency to average.

These anecdotal impressions of performance are borne out by a more global assessment of error relative to AD99. Figure 4 shows the three emulators' coefficients of determination R^2 , the fraction of target variance accounted for by emulator output, as a function of vertical level (on the left) and latitude (on the right). The R^2 score is unity for an emulator that explains the data exactly, zero for constant predictions of the target mean, and arbitrarily negative as emulator performance degrades. Dashed lines indicate performance on the training data, solid lines on the test data unseen during training. All three emulators perform slightly better on the training data than the held-out test data, as is expected, but no emulator exhibits the large train-test gap characteristic of overfitting. Figure 4 shows

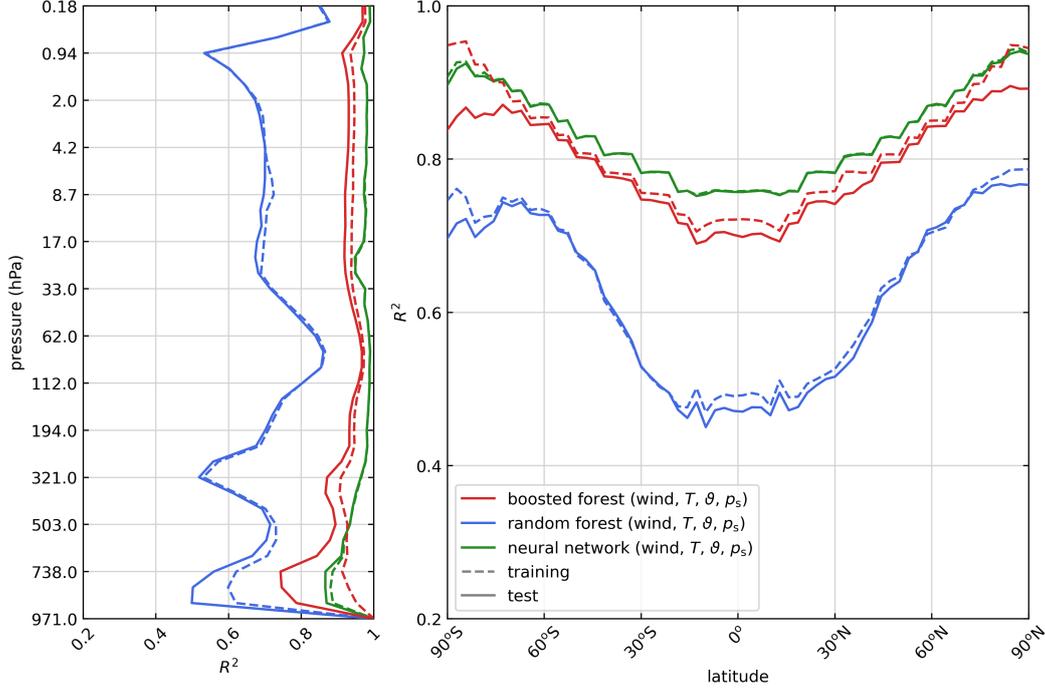


Figure 4. Offline R^2 scores of three data-driven emulators on the training and test datasets. Parentheses indicate the input features used by each model.

323 that the neural network has the best offline performance, closely followed by the boosted
 324 forest, with the random forest much worse than either.

325 Performance tends to be better aloft than near the surface. In AD99, layers below
 326 321 hPa are sometimes below the tropopause-following source level and sometimes above it,
 327 depending on latitude, and emulator predictions are worse at these lower levels. Emulator
 328 error is also larger in the tropics, likely because the peak of the AD99 source spectrum
 329 switches there to following the mean flow at the source level, making the emulation task
 330 more complex. This degraded performance is not an artifact of sampling, because training
 331 samples were sampled weighted by grid box area, so that the tropics are well-represented in
 332 the training data.

333 Following Espinosa et al. (2022), we train boosted forests on different combinations
 334 of input features to assess the utility of various physical variables. All forests used the
 335 hyperparameter choices described in Table S3. Figure 5 shows the R^2 scores for these models.
 336 We observe that changing the input features does not result in significant performance
 337 increases or decreases, except that the forest that does not see latitude ϑ performs very

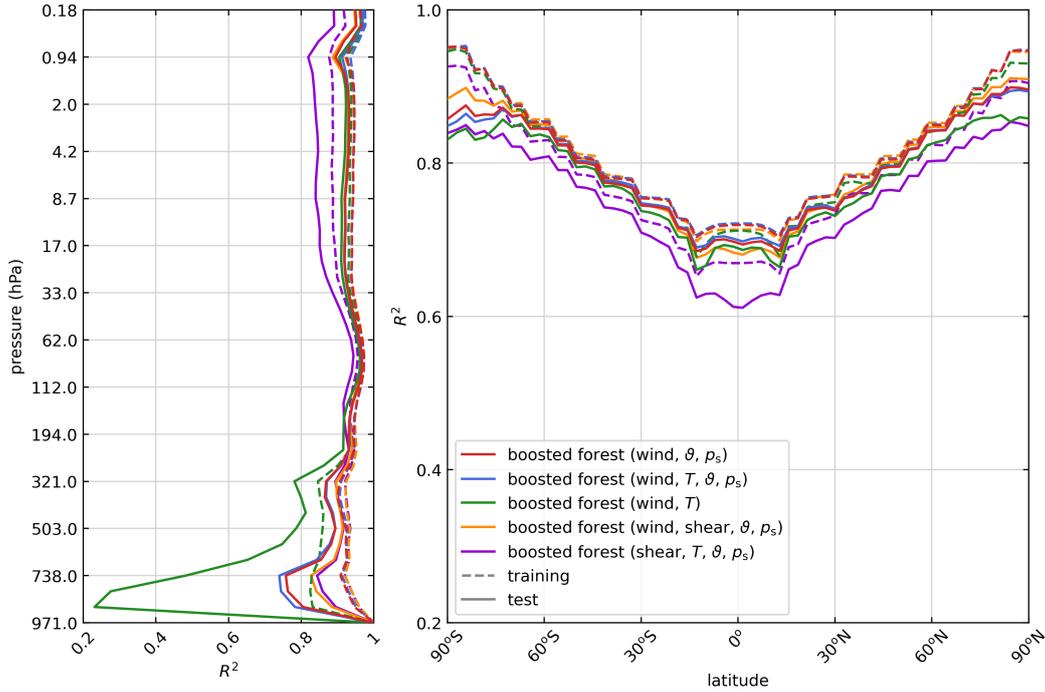


Figure 5. As in Figure 4, but for boosted forests using different combinations of input features.

338 poorly at vertical levels which can be below the source level. Otherwise, the differences
 339 between particular boosted forests are smaller than those between boosted forests and the
 340 other architectures considered in this work. For comparisons between architectures, we
 341 use the boosted forest trained on wind, temperature, latitude, and surface pressure, as it
 342 outperforms the others by a slight margin in the tropics, where the GWP-driven QBO
 343 occurs. This is also the same feature set used by AD99.

344 We performed a similar set of experiments with random forests. Again, their perfor-
 345 mance was largely unaffected by changing the input features, and all the random forests
 346 remained substantially worse than the boosted forests, as Figure 4 suggests. For this reason,
 347 the bulk of the analysis in the remainder of this work is focused on boosted forest emulators,
 348 with comparisons to random forests only when appropriate.

349 4.2 Climatological biases

350 Figure 6 shows the biases in zonal mean u and T relative to the AD99 control run of the
 351 coupled runs described in Section 2.3. The boosted forest shows the least bias overall. The
 352 random forest and neural network produce large biases mainly in the tropical stratosphere,

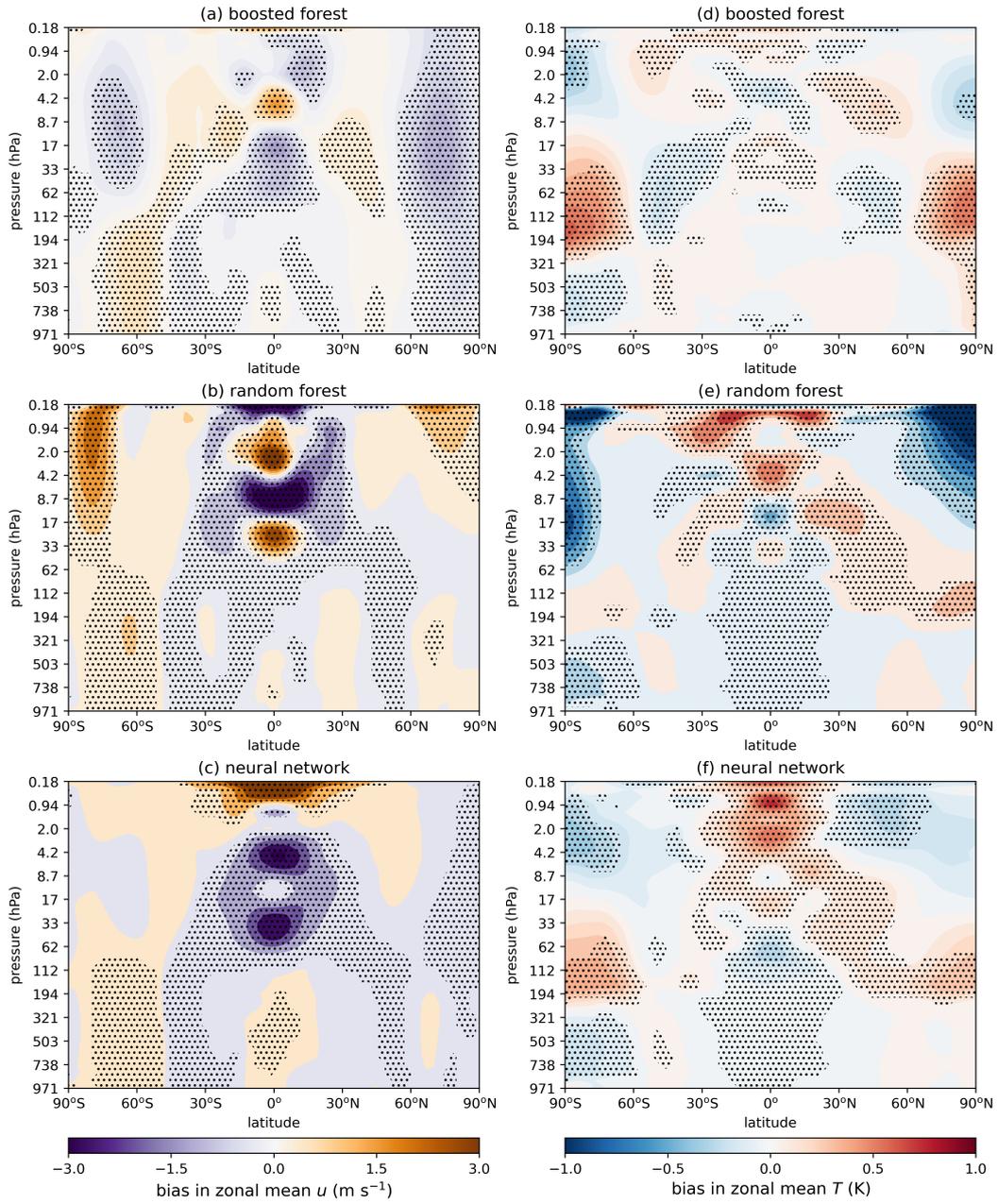


Figure 6. Biases with respect to AD99 integrations in zonal mean u (a-c, left column) and T (d-f, right column) from coupled integrations of the three emulators shown in Figure (4). Stippling indicates regions where bias is significant at the 95% level.

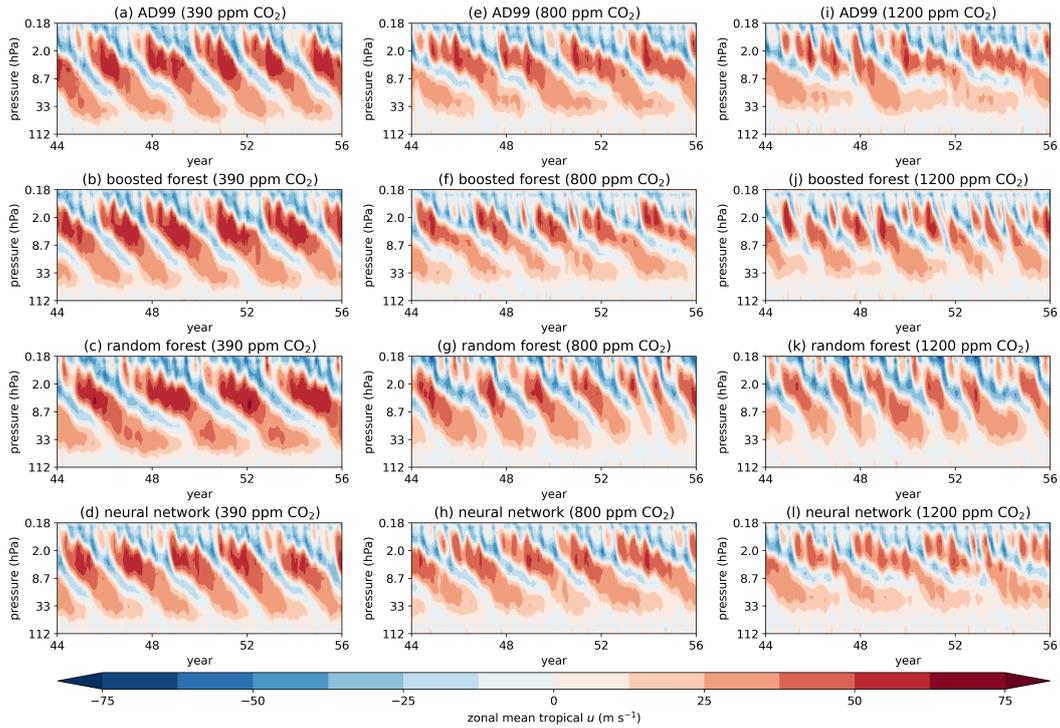


Figure 7. QBOs from the final twelve years of integrations of MiMA coupled to AD99 (top row) and the three data-driven emulators from Figure 4 (next three rows). The control climate integrations are on the left, the 800 ppm CO₂ runs in the middle, and the 1200 ppm CO₂ runs on the right.

353 where variability from the QBO dominates, and to a lesser extent near the poles. In particu-
 354 lar, the tropospheric jet structure is emulated well and free of systematic bias. Remarkably,
 355 the random forest, which exhibited much poorer offline performance in Figure 4, runs stably
 356 online without significantly altering the zonal mean climate in the troposphere.

357 The zonal-mean biases were similar in the coupled runs with increased atmospheric
 358 carbon dioxide. The remainder of this section will focus on climate phenomena that are
 359 particularly dependent on the particular gravity wave parameterization: the QBO and the
 360 occurrence of sudden stratospheric warmings (SSWs).

361 4.3 The Quasi-Biennial Oscillation

362 Atmospheric models generally need parameterized gravity wave drag to represent the
 363 QBO in the tropical stratosphere (Anstey et al., 2022), and so the emulators' skill at re-

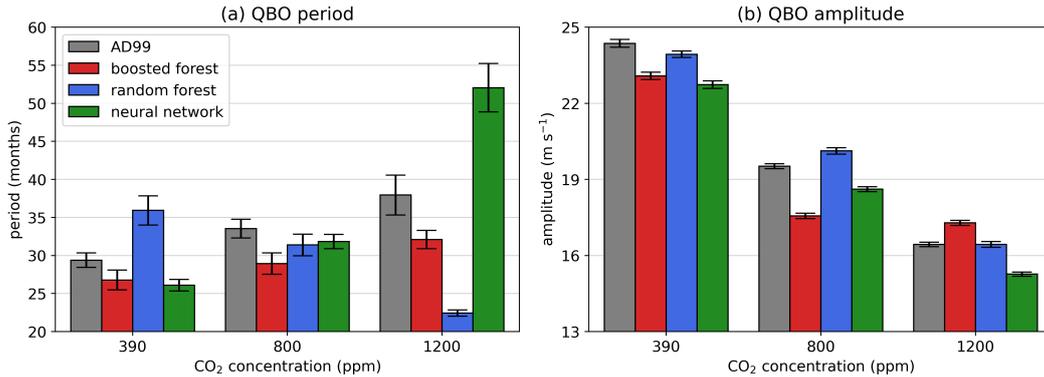


Figure 8. Periods (left) and amplitudes (right) of QBOs driven by AD99 and three emulators in integrations with three values of CO₂ concentration. Periods are determined by the dominant Fourier mode at 10 hPa, with uncertainty given by the half-width of the spectral peak around that mode. Amplitudes are the standard deviation at 10 hPa, with uncertainty determined from the 95% confidence intervals as calculated by bootstrapping.

364 producing the statistics of the QBO simulated by AD99 is a key metric of their online
 365 performance. For simplicity, we will refer to any simulated oscillation in the tropically- and
 366 zonally-averaged zonal wind as a QBO, even when the period is no longer “quasi-biennial”.
 367 Figure 7 shows the QBO time series from MiMA integrations coupled to AD99 and the three
 368 emulator architectures. In all three carbon scenarios, AD99 and each of the three emulators
 369 produce a QBO. Especially in the high-carbon integrations, though, these oscillations vary
 370 both qualitatively and quantitatively. The period and amplitude response of the QBO for
 371 each parameterization is shown in Figure 8.

372 The period of the QBO driven by AD99 increases monotonically with CO₂ concen-
 373 tration. The sign of this response is captured by the boosted forest and neural network
 374 emulators, though the boosted forest periods are biased low and the neural network signif-
 375 icantly overshoots the AD99 period in the 1200 ppm case. (Note, however, that the QBO
 376 in Figure 7i is not nearly as divergent from the AD99 oscillation as this calculation might
 377 suggest.) The random forest, by contrast, drives a QBO with a significantly longer period in
 378 the control 390 ppm CO₂ scenario, and the period decreases with increasing carbon dioxide.
 379 This failure to replicate the behavior of AD99 is perhaps unsurprising, given the random
 380 forest’s relatively poor offline performance (Figure 4).

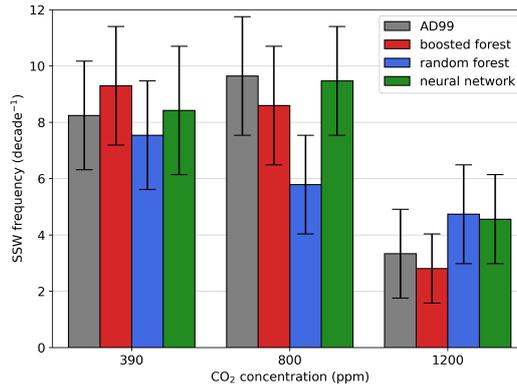


Figure 9. Sudden stratospheric warming frequencies from the MiMA runs described in Section 2.3. SSWs are identified using the criterion of Butler et al. (2017), based on sign changes in the zonal mean zonal wind at 60°N and 10 hPa. The uncertainties are the 95% confidence intervals, as calculated by bootstrapping winters and counting SSWs in the subsampled populations. This approach assumes that SSWs in different winters are independent.

381 The amplitude of the reference QBO decreases in response to increased CO₂. All three
 382 emulators show at least broadly similar behavior (Figure 8(b)), though the boosted forest
 383 significantly overestimates the response in the 800 ppm CO₂ case, and the neural network
 384 amplitudes are biased low. It is of note that the random forest, heretofore the worst of the
 385 three emulators, is the most able to reproduce the correct QBO amplitudes across all three
 386 integrations.

387 4.4 Sudden stratospheric warmings

388 Sudden stratospheric warmings (SSWs) are abrupt increases in the temperature of
 389 the wintertime polar stratosphere accompanied by a reversal in the polar vortex. Their
 390 occurrence is governed in part by gravity wave propagation — either directly through wave-
 391 driven momentum flux (Song et al., 2020) or indirectly through moderation by the QBO
 392 (Butler et al., 2017) — and their frequency thus provides an additional statistic with which
 393 to assess the emulators’ online performance. Figure 9 shows the SSW frequencies for AD99
 394 and the three emulators at three CO₂ concentrations.

395 In the 390 ppm CO₂ integrations, the random forest appears to best reproduce the SSW
 396 frequency, but the uncertainties are large and the confidence intervals for all three emulators
 397 overlap considerably with that of AD99. When the CO₂ concentration is raised to 800 ppm,

AD99 produces SSWs more frequently. This response is not well captured by any emulator: none of them demonstrates a large response relative to the uncertainties. However, in the more extreme 1200 ppm CO₂ runs, all three emulators capture the large decrease in SSW frequency exhibited by AD99.

5 Feature importance analysis

The offline and online results presented in Section 4 show that the neural network slightly outperforms the boosted forest offline, performs comparably online in the control climate, and is somewhat better at reproducing the AD99 response to the 800 ppm CO₂ scenario. The final goal of this study is to use the interpretability tools laid out in Section 3.4 to calibrate the online behavior of regression forest models towards more desirable outcomes.

5.1 Computed SHAP values

To further analyze the behavior of our emulators, we calculate SHAP values (Section 3.4), a measure of feature importance, the relative importance a model ascribes to various input features in making its predictions. Figure 10 shows dataset-averaged SHAP values for the three emulators in Figure 4 and for the AD99 parameterization itself. Each panel shows the importance of input features from all levels (wind and temperature importances are summed per level) to predictions at a single level. Note that SHAP values have the same units as the parameterization outputs, which in this case are standard deviations of the gravity wave acceleration at the prediction level.

The SHAP profiles for AD99, the boosted forest, and the neural network show preferential use of input information from at and just below the prediction level. This pattern matches our physical intuition, given that AD99 simulates strictly upward propagation of waves, and indicates that the boosted forest and neural network have learned to make predictions according to the “physics” encoded in AD99. (The importance maxima at levels immediately above the prediction level do not contradict this characterization: AD99 calculates drags at level interfaces before interpolating to full levels, so that information from one level higher than the prediction level is used.)

The random forest profiles, however, are much more muted. This suggests that the random forest does not identify the set of features on which the drag at each level depends, reflecting a failure to learn the physical structure of the problem and at least partially

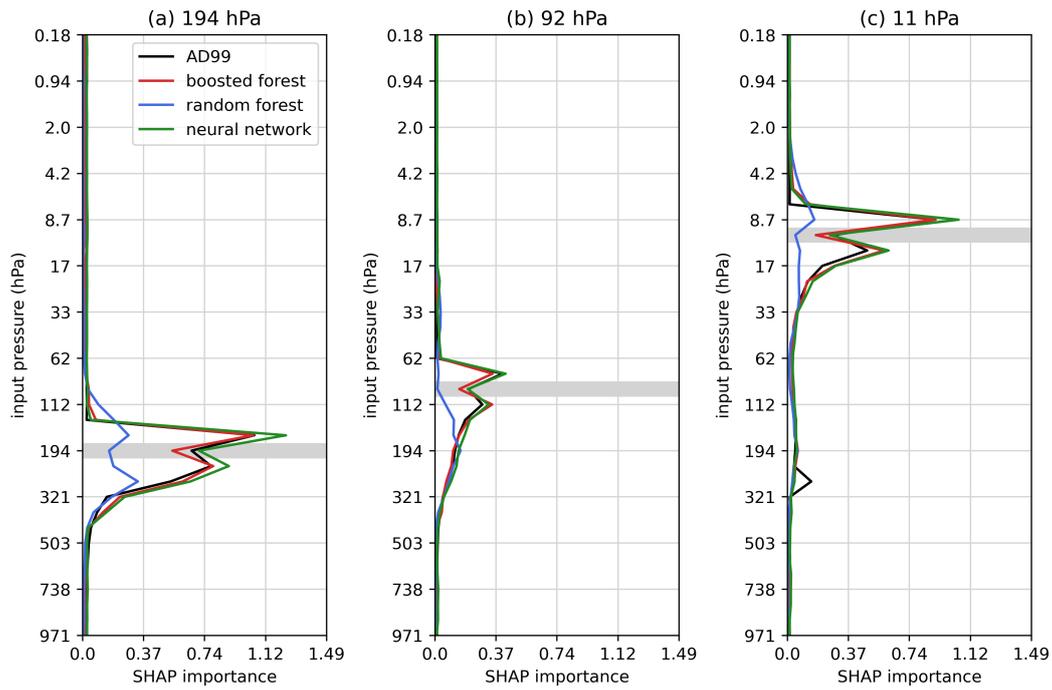


Figure 10. Test dataset-averaged absolute SHAP values for predictions at several vertical levels for AD99 and the three emulators in Figure 4. Each input pressure level includes the importances of both wind and T features. The prediction level is highlighted in gray.

428 explaining the poor R^2 scores in Figure 4. The emulators with high R^2 scores, the boosted
 429 forest and neural network, use the input features almost identically to AD99 — as opposed
 430 to, say, achieving good performance by relying on spurious correlations between features.

431 Of additional interest is the peak in the AD99 importance profile near 200 hPa in
 432 Figure 10(c), showing that even for predictions aloft, AD99 makes use of layers near the
 433 tropopause. Indeed, in AD99 the tropical phase speed spectrum is set by the source level
 434 winds, and many waves are immediately filtered at the tropopause. All three emulators
 435 appear to under-emphasize input information near the source level.

436 For the random and boosted forests, we computed *Gini importances*, an additional
 437 metric of feature importance defined only for regression tree-based architectures, and found
 438 them to be in good qualitative agreement with the SHAP values shown here. See Text S2
 439 and Figure S4 for details. We therefore believe the conclusions about the models considered
 440 here to be reasonably robust to feature importance method.

441 5.2 SHAP-informed retraining

442 Although Figure 10 indicates that the boosted forest and neural network use wind and
 443 temperature information in a physically plausible way, it does not explain the differences in
 444 online behavior observed in Figures 7 and 8. The left panel of Figure 11 shows the SHAP
 445 importance of latitude to all prediction levels for AD99 and the three emulators. AD99 has a
 446 large maximum in latitude SHAP value near the tropopause. Strikingly, the neural network
 447 values match this maximum quite closely, while the boosted forest considerably under-
 448 emphasizes latitude in this region of the atmosphere. (Further aloft, all three emulators
 449 have latitude SHAP values less than those of AD99.)

450 This result suggests that underuse of latitude might be a key factor differentiating the
 451 boosted forest from the neural network. Moreover, one might wonder if a boosted forest
 452 forced to pay closer attention to latitude might have better online behavior. In particular,
 453 the distribution of input latitudes is necessarily fixed, while the distribution of input flow
 454 variables is subject to shift under climate perturbations.

455 To test this, we trained a boosted forest with identical hyperparameters to the one
 456 considered thus far; however, we added the latitude of each training sample as an additional
 457 target. The idea is that, since latitude is now both an input and an output, the input

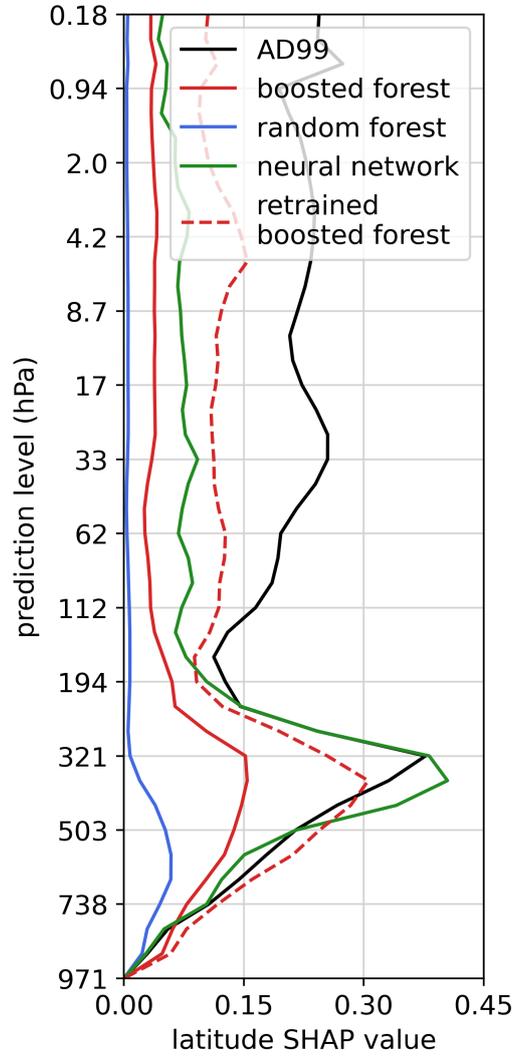


Figure 11. Test dataset-averaged absolute SHAP importance of latitude to parameterization outputs at each vertical level for AD99, the three emulators from Figure 4 (solid lines), and the boosted forest retrained as described in Section 5.2.

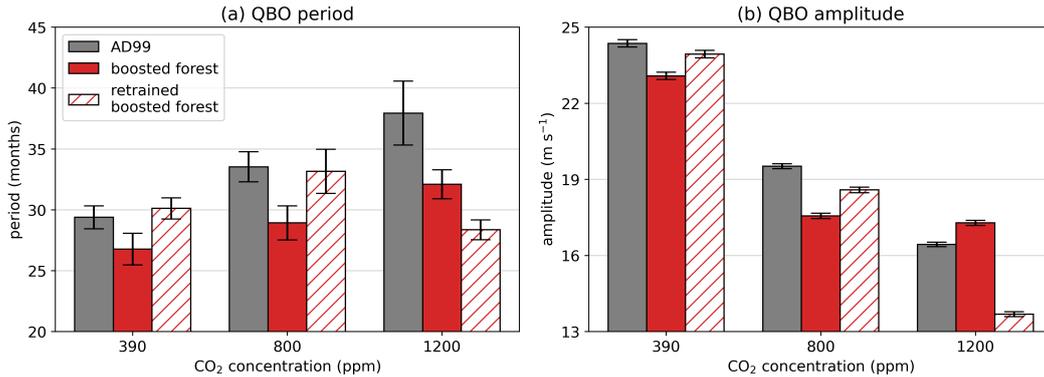


Figure 12. As in Figure 8, but including the boosted forest trained as described in Section 5.2.

458 latitude should be more useful in partitioning the training output vectors into low-impurity
 459 subsets (see Text S1). Latitude should therefore be selected at more important nodes in
 460 the constituent trees. We can multiply the latitude values added as targets by a constant
 461 (introducing one new hyperparameter) to control the strength of this effect. At prediction
 462 time, we simply discard the predicted latitude and retain only the predicted drags.

463 The boosted forest trained in this manner achieves R^2 scores (not shown) essentially
 464 indistinguishable from the boosted forest in Figure 4. Nor was the use of level-specific infor-
 465 mation (as in Figure 10) significantly different from the original boosted forest. However,
 466 the dashed line in Figure 11 demonstrates that this training procedure produces a forest
 467 for which latitude is two to three times as important throughout the vertical extent of the
 468 atmosphere. More significantly, Figure 12 shows that when coupled to MiMA, the retrained
 469 boosted forest drives a QBO with period and amplitude closer to those of AD99 than the
 470 original forest in both the 390 ppm and 800 ppm CO₂ scenarios. The QBO statistics in
 471 the 1200 ppm integration remain poor, suggesting that this extreme carbon perturbation is
 472 simply beyond the ability of this boosted forest architecture to generalize.

473 6 Discussion

474 In this study, we used boosted and random forests to emulate a gravity wave parameter-
 475 ization, performed offline and online evaluations of emulator performance, and investigated
 476 whether emulator use of input features respected known physical properties of the data. To
 477 our knowledge, this work represents the first use of boosted forests in the design of param-
 478 eterizations for climate models. Boosted forests are not uncommon in the wider machine

479 learning literature; for example, they are often used in winning submissions to competitions
480 (Bojer & Meldgaard, 2021). But while random forests have been employed in several param-
481 eterization studies (Belochitski et al., 2011; O’Gorman & Dwyer, 2018; Yuval & O’Gorman,
482 2020), boosted forests have not, perhaps because of the absence of native support in boosting
483 libraries for the multioutput problems that abound in climate modeling.

484 We have shown that boosted forests significantly outperform random forests accord-
485 ing to almost every metric, even with a relatively simple implementation of multioutput
486 boosting that lacks much of the flexibility of the optimized boosting libraries used for scalar
487 problems. Boosted forests may therefore be a valuable and yet-underused tool as climate
488 models continue to move towards incorporating data-driven parameterizations, especially
489 since they were as skillful, or nearly so, as neural networks. In particular, even though a
490 boosted forest makes inferences based exclusively on outputs in the training data, ours was
491 able to generalize to out-of-sample conditions equally as well as the neural network bench-
492 mark: both schemes performed well under moderately enhanced CO₂, but struggled under
493 our most extreme scenario. Out-of-sample generalization is often challenging for data-driven
494 methods, but recent work by Sun et al. (submitted) suggests that transfer learning may be
495 a solution if one can provide a small amount of high-quality data from the new regime —
496 for example, data from a high-resolution simulation under increased CO₂.

497 We further interrogated our data-driven schemes with methods from interpretable ma-
498 chine learning to quantify how they used input features to make predictions. While the
499 Gini importance is a natural interpretability metric for regression forests (Text S2), we
500 found that it provided nearly the same information (Figure S4) as SHapley Additive ex-
501 Planation (SHAP) analysis (Lundberg & Lee, 2017), a method-agnostic approach that can
502 be used on any ML scheme and even on the original physics-based parameterization. The
503 skillful boosted forest and neural network emulators exhibited SHAP values nearly matching
504 those of AD99, much more closely than did the under-performing random forest (Figure 10).
505 For the machine learning architectures considered in this work, then, emulation skill appears
506 to go hand in hand with capacity to learn elements of the spatial structure of the problem.
507 This observation suggests that these kinds of data-driven models may be able to infer similar
508 structures from more realistic data sources, for which the true SHAP values will of course
509 be unavailable.

510 Moreover, the analysis in Section 5.2 demonstrates the utility of SHAP analysis for
511 understanding and improving online behavior beyond more common error metrics like R^2
512 scores. The offline errors in Figure 4 showed the boosted forest performing worse than the
513 neural network, and the QBO statistics in Figure 8 confirmed the forest to be less successful
514 online under certain scenarios. Only the SHAP analysis culminating in Figure 11, though,
515 provided actionable information, informing us that the boosted forest was not sufficiently
516 taking into account latitudinal variation at prediction time. This informed a training proce-
517 dure to improve the online behavior. Our approach remains somewhat *ad hoc*: the decision
518 to focus on the input latitude was made by eye, and there may be superior ways to constrain
519 forests to emphasize given input features. Nonetheless, we believe this result to be a useful
520 preliminary towards calibrating the online behavior of data-driven parameterizations that
521 may lack the explicit parameters used to tune physics-based schemes.

522 Finally, from a practical standpoint, multioutput boosted forest libraries like the one
523 implemented here will need to be made more efficient and self-contained if they are to provide
524 a competitive alternative to neural networks in climate research. We found anecdotally that
525 boosted forests constructed with only a few deep trees followed by many much shallower ones
526 could perform nearly as well as, and considerably faster than, the constant-depth forests
527 used here, though more investigation is required to fully explore the interplay between forest
528 size and skill.

529 **7 Data Availability Statement**

530 The code used to run the experiments in this work is available, with documentation, at
531 <https://github.com/dsconnelly/willow>. The random and boosted forests were trained
532 using Mubofo (Connelly, 2023), a Python package maintained by author D. S. Connelly and
533 available through Python Package Index (PyPI) at <https://pypi.org/project/mubofo>.
534 The source code may be found at <https://github.com/dsconnelly/mubofo> and is dis-
535 tributed under the BSD-3-Clause license.

536 Mubofo is built around scikit-learn (Pedregosa et al., 2011), and the neural network was
537 trained using PyTorch (Paszke et al., 2019). The SHAP values were computed with the shap
538 Python package (Lundberg & Lee, 2017) available at <https://github.com/shap/shap>. The
539 idealized atmospheric model MiMA is described in Jucker and Gerber (2017) and Garfinkel
540 et al. (2020) and is available at <https://github.com/mjucker/MiMA>. The coupling interface

541 Forpy (Rabel, 2020) is available at <https://github.com/ylikx/forpy>. The authors are
542 not involved with the maintenance of any of these software packages.

543 **Acknowledgments**

544

545 This work was supported by Schmidt Futures, a philanthropic initiative founded by
546 Eric and Wendy Schmidt, as part of the Virtual Earth System Research Institute (VESRI),
547 the U.S. National Science Foundation through award OAC-2004572, and the U.S.-Israel
548 Binational Science Foundation through award 2020316.

549 **References**

- 550 Alexander, M. J., & Dunkerton, T. J. (1999). A Spectral Parameterization of Mean-Flow
551 Forcing due to Breaking Gravity Waves. *Journal of the Atmospheric Sciences*, *56*(24),
552 4167-4182.
- 553 Anstey, J. A., Butchart, N., Hamilton, K., & Osprey, S. M. (2022). The sparc quasi-biennial
554 oscillation initiative. *Quarterly Journal of the Royal Meteorological Society*, *148*(744),
555 1455-1458.
- 556 Belochitski, A., Binev, P., DeVore, R., Fox-Rabinovitz, M., Krasnopolsky, V., & Lamby, P.
557 (2011). Tree approximation of the long wave radiation parameterization in the NCAR
558 CAM global climate model. *Journal of Computation and Applied Mathematics*, *236*,
559 447-460.
- 560 Betts, A. K. (1986). A new convective adjustment scheme. Part I: Observational and
561 theoretical basis. *Quarterly Journal of the Royal Meteorological Society*, *112*(473),
562 677-691.
- 563 Betts, A. K., & Miller, M. J. (1986). A new convective adjustment scheme. Part II:
564 Single column tests using GATE wave, BOMEX, ATEX and arctic air-mass data sets.
565 *Quarterly Journal of the Royal Meteorological Society*, *112*(473), 693-709.
- 566 Bojer, C. S., & Meldgaard, J. P. (2021). Kaggle forecasting competitions: An overlooked
567 learning opportunity. *International Journal of Forecasting*, *37*, 587-603.
- 568 Bolton, T., & Zanna, L. (2019). Applications of Deep Learning to Ocean Data Inference
569 and Subgrid Parameterization. *Journal of Advances in Modeling Earth Systems*, *11*,
570 379-399.
- 571 Breiman, L. (2001). Random Forests. *Machine Learning*, *45*(1), 5-32.

- 572 Breiman, L., Friedman, J. H., Olshen, R. A., & Stone, C. J. (1984). *Classification and*
573 *Regression Trees*. Chapman and Hall/CRC.
- 574 Brenowitz, N. D., & Bretherton, C. S. (2018). Prognostic Validation of a Neural Network
575 Unified Physics Parameterization. *Geophysical Research Letters*, *45*, 6289-6298.
- 576 Butler, A. H., Sjoberg, J. P., Seidel, D. J., & Rosenlof, K. H. (2017). A sudden stratospheric
577 warming compendium. *Earth System Science Data*, *9*, 63-76.
- 578 Chantry, M., Hatfield, S., Dueben, P., Polichtchouk, I., & Palmer, T. (2021). Machine
579 Learning Emulation of Gravity Wave Drag in Numerical Weather Forecasting. *Journal*
580 *of Advances in Modeling Earth Systems*, *13*.
- 581 Checa-Garcia, R. (2018). *CMIP6 Ozone forcing dataset: supporting information*. Zenodo.
582 doi: 10.5281/zenodo.1135127
- 583 Checa-Garcia, R., Hegglin, M. I., Kinnison, D., Plummer, D. A., & Shine, K. P. (2018).
584 Historical Tropospheric and Stratospheric Ozone Radiative Forcing Using the CMIP6
585 Database. *Geophysical Research Letters*, *45*(7), 3264-3273.
- 586 Chevallier, F., Ch eruy, F., Scott, N. A., & Ch edin, A. (1998). A Neural Network Approach
587 for a Fast and Accurate Computation of a Longwave Radiation Budget. *Journal of*
588 *Applied Meteorology*, *37*, 1385-1397.
- 589 Connelly, D. S. (2023). *Mubofo*. Github repository. Retrieved from [https://github.com/](https://github.com/dsconnelly/mubofo)
590 [dsconnelly/mubofo](https://github.com/dsconnelly/mubofo)
- 591 Espinosa, Z. I., Sheshadri, A., Cain, G. R., Gerber, E. P., & DallaSanta, K. J. (2022).
592 Machine Learning Gravity Wave Parameterization Generalizes to Capture the QBO
593 and Response to Increased CO₂. *Geophysical Research Letters*, *49*(8).
- 594 Friedman, J. H. (2001). Greedy Function Approximation: A Gradient Boosting Machine.
595 *The Annals of Statistics*, *29*(5), 1181-1232.
- 596 Fritts, D. C., & Alexander, M. J. (2003). Gravity wave dynamics and effects in the middle
597 atmosphere. *Reviews of Geophysics*, *41*(1).
- 598 Garfinkel, C. I., White, I., Gerber, E. P., Jucker, M., & Erez, M. (2020). The Building
599 Blocks of Northern Hemisphere Wintertime Stationary Waves. *Journal of Climate*,
600 *33*(13), 5611-5633.
- 601 Iacono, M. J., Mlawer, E. J., Clough, S. A., & Morcrette, J.-J. (2000). Impact of an
602 improved longwave radiation model, RRTM, on the energy budget and thermodynamic
603 properties of the NCAR community climate model, CCM3. *Journal of Geophysical*
604 *Research*, *105*(D11), 14873-14890.

- 605 Janzing, D., Minorics, L., & Bloebaum, P. (2020). Feature relevance quantification in
 606 explainable AI: A causal problem. In *Proceedings of the Twenty Third International*
 607 *Conference on Artificial Intelligence and Statistics* (Vol. 108, p. 2907-2916). PMLR.
- 608 Jucker, M., & Gerber, E. P. (2017). Untangling the Annual Cycle of the Tropical Tropopause
 609 Layer with an Idealized Moist Model. *Journal of Climate*, *30*(18), 7339-7358.
- 610 Lundberg, S. M., Erion, G., Chen, H., DeGrave, A., Prutkin, J. M., Nair, B., ... Lee, S.-I.
 611 (2020). From local explanations to global understanding with explainable AI for trees.
 612 *Nature Machine Intelligence*, *2*(1), 56–67.
- 613 Lundberg, S. M., & Lee, S.-I. (2017). A Unified Approach to Interpreting Model Predictions.
 614 In *Advances in Neural Information Processing Systems* (Vol. 30). Curran Associates,
 615 Inc.
- 616 Mlawer, E. J., Taubman, S. J., Brown, P. D., Iacono, M. J., & Clough, S. A. (1997).
 617 Radiative transfer for inhomogeneous atmospheres: RRTM, a validated correlated-k
 618 model for the longwave. *Journal of Geophysical Research*, *102*(D14), 16663-16682.
- 619 O’Gorman, P. A., & Dwyer, J. G. (2018). Using Machine Learning to Parameterize Moist
 620 Convection: Potential for Modeling of Climate, Climate Change, and Extreme Events.
 621 *Journal of Advances in Modeling Earth Systems*, *10*, 2548-2563.
- 622 Palmer, T. N., Shutts, G. J., & Swinbank, R. (1986). Alleviation of a systematic westerly bias
 623 in general circulation and numerical weather prediction models through an orographic
 624 gravity wave drag parametrization. *Quarterly Journal of the Royal Meteorological*
 625 *Society*, *112*, 1001-1039.
- 626 Paszke, A., Gross, S., Massa, F., Lerer, A., Bradbury, J., Chanan, G., ... Chintala, S.
 627 (2019). PyTorch: An Imperative Style, High-Performance Deep Learning Library.
 628 In *Advances in neural information processing systems 32* (pp. 8024–8035). Curran
 629 Associates, Inc.
- 630 Pearl, J. (2000). *Causality*. Cambridge University Press.
- 631 Pedregosa, F., Varoquaux, G., Gramfort, A., Michel, V., Thirion, B., Grisel, O., ... Duch-
 632 esnay, E. (2011). Scikit-learn: Machine learning in Python. *Journal of Machine*
 633 *Learning Research*, *12*, 2825-2830.
- 634 Rabel, E. (2020). *Forpy*. Github repository. Retrieved from [https://github.com/ylikx/](https://github.com/ylikx/forpy)
 635 [forpy](https://github.com/ylikx/forpy)
- 636 Richter, J. H., Butchart, N., Kawatani, Y., Bushell, A. C., Holt, L., Serva, F., ... Yukimoto,
 637 S. (2022). Response of the Quasi-Biennial Oscillation to a warming climate in global

- 638 climate models. *Quarterly Journal of the Royal Meteorological Society*, *148*(744),
639 1490-1518.
- 640 Schapire, R. E. (1990). The strength of weak learnability. *Machine Learning*, *5*, 197-227.
- 641 Shapley, L. S. (1953). A value for n-person games. In H. W. Kuhn & A. W. Tucker (Eds.),
642 *Contributions to the Theory of Games* (p. 307-317).
- 643 Shrikumar, A., Greenside, P., & Kundaje, A. (2017). Learning Important Features Through
644 Propagating Activation Differences. In *Proceedings of the 34th International Confer-*
645 *ence on Machine Learning* (pp. 3145–3153). PMLR.
- 646 Song, I.-S., Lee, C., Chun, H.-Y., Kim, J.-H., Jee, G., Song, B.-G., & Bacmeister, J. T.
647 (2020). Propagation of gravity waves and its effects on pseudomomentum flux in a
648 sudden stratospheric warming event. *Atmospheric Chemistry and Physics*, *20*, 7617-
649 7644.
- 650 Sun, Y. Q., Pahlavan, H. A., Chattopadhyay, A., Hassanzadeh, P., Lubis, S. W., Alexan-
651 der, M. J., ... Guan, Y. (submitted). Data Imbalance, Uncertainty Quantification,
652 and Generalization via Transfer Learning in Data-driven Parameterizations: Lessons
653 from the Emulation of Gravity Wave Momentum Transport in WACCM. *Journal of*
654 *Advances in Modeling the Earth System*.
- 655 Yuval, J., & O’Gorman, P. (2020). Stable machine-learning parameterization of subgrid
656 processes for climate modeling at a range of resolutions. *Nature Communications*, *11*.